

**Flow-dependent reliability:** A practical route to more skilful ensemble forecasts Mark Rodwell With **David Richardson** Heini Wernli **Dave Parsons** Simon Lang ... and many others

5<sup>th</sup> WGNE workshop on systematic errors in weather and climate models

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Shinfield Park, Reading, RG2 9AX, UK

### Forecast reliability

#### Ensemble forecast of storm location



In a "**reliable**" forecast system, the truth can be considered as another ensemble member

Reliability is very useful: if we predict an event with probability 70%, it will happen with frequency 70%

A testable consequence of reliability is that:

average Error = average Spread (averaged over many forecast start dates)

Given we had a reliable system, progress would be ...

Predicting "**sharper**" (tighter) distributions **while retaining reliability** 

Longitude



#### Ensemble spread and error



In 1996, Spread ≈ RMSE to D+2 due to Singular Vector (SV) initial perturbations. Since then, Error and Spread have converged and reduced throughout forecast range due to improved observations, model, representation of observation and model uncertainty, and introduction of Ensemble of Data Assimilations (EDA); allowing weaker SV perturbations

... but uncertainty varies from day-to-day. The real reason we make ensemble forecasts. What causes this, and how can we evaluate it in our forecasts?

500 hPa geopotential height (Z500). "Error" is RMS of ensemble-mean error Spread = ensemble standard deviation (scaled to account for finite ensemble size)



Potential Vorticity on the Potential Temperature = 320K surface. 20110410 0 UTC. Step (days, hours) = 0 00.0





Potential Vorticity on the Potential Temperature = 320K surface. 20110410 0 UTC. Step (days, hours) = 6 00.0





Potential Vorticity on the Potential Temperature = 320K surface. 20110404 0 UTC. Step (days, hours) = 0 00.0





Potential Vorticity on the Potential Temperature = 320K surface\_20110404 0 UTC\_Step (days, hours) = 6 00.0





# Animation of v<sub>850</sub>, Total precip., PV<sub>315</sub>=2 (tropopause), $\hat{\sigma}_{PV_t}/\hat{\sigma}_{PV}$ (tendency uncertainty)

20161127 18Z



Ensembles represent the integral of instantaneous uncertainty growth-rate; would like to diagnose deficiencies in the instantaneous growth-rate Here consider 1h tendencies in EDA background forecasts. Filter to T21, 1d to emphasise growth-rate within synoptic-scale systems  $\hat{\sigma}_{PV_t}$  used to highlight dynamical (*e.g.* baroclinic) and physical (*e.g.* stochastic) sources of uncertainty. 315K  $\Rightarrow$  strong propagation in Jetstream Results emphasise role of moist processes (MCS, WCB); intrinsic to real system or artefact of *e.g.* deficiencies in model uncertainty? Key question



#### Average initial conditions of 584 single forecast "busts" over Europe at day 6

#### a Z500 anomaly

Rodwell et al, 2013, BAMS



Trough over the Rocky mountains, with high convective potential ahead Conducive to the formation of mesoscale convection

Interesting flow regime to evaluate 'instantaneous' uncertainty growthrate

(Using different set of dates to avoid misleading results)



'CAPE' = Convective Available Potential Energy



## Short-range variance assessment for u200 in "trough/CAPE" situations using EDA



#### ECIVITE European C

## Short-range mean error assessment for T300 in "trough/CAPE" situations using DA

Model-space process tendency and analysis increment budget





<sup>54</sup> cases

### The Jetstream and mesoscale convection: "The piano string and hammer"

#### 54 cases

Met3D: Marc Rautenhaus



If we don't hit the string hard enough, the wave in the string will be too weak

If we hit the string at the wrong time, the wave will arrive over Europe at the wrong time

We do not know when to press the key (mesoscale convection itself involves chaotic uncertainty)

What we want is that the ensemble members generate such convection with the "right" uncertainty





### Short-range variance assessment for mid-tropospheric temperatures: Control



Residual suggests EnsVar underestimates (overestimates) uncertainty in convective (clear-sky) regions (and/or deficiencies in ObsUnc<sup>2</sup>)



## Short-range variance assessment for mid-tropospheric temperatures: No stoch. phys.



EnsVar is highly sensitive to representation of model uncertainty. Turning it off deteriorates (improves) Residual in convective (clear-sky) regions



### Top 50 Warm Conveyor Belt inflow events in box indicated from Nov 15 – Oct 16

Inflow D+0 ( > 800 hPa )



From Heini Werni. Based on trajectories ascending by more than 600 hPa in 2d



Outflow D+1 (< 400 hPa )



## EDA variance assessment with ASCAT surface v wind: Non-WCB composite

87 cases

Bias<sup>2</sup> and Residual are not significant in absence of WCBs.

ObsUnc<sup>2</sup> a large component





Observation density (O80, 12h)

Unit: 0.1 cell<sup>-1</sup> cycle<sup>-1</sup> Mean: 16 Sig: 99% 0 4 8 12 16 20 24 36





## EDA variance assessment with ASCAT surface v wind: WCB (inflow) composite





## EDA variance assessment with MHS "all sky" mid-tropospheric humidity: Non-WCB

Bias and residual are not significant in

absence of WCBs

87 cases



Unit: (K)<sup>2</sup> Mean: 16.8 Sig: 95% 5 10 15 20 25 30 730





Unit: (K)<sup>2</sup> Mean: -12.7 Sig: 61%

-610 -50 -30 -10 10 30 50 70



Unit: (K)<sup>2</sup> Mean: 5.83 Sig: 93%

10 15 20 25 30 150

EnsVar

5

Observation density (O80, 12h) Unit: cell<sup>-1</sup>cycle<sup>-1</sup> Mean: 4.05 Sig: 96% 2 4 6 8 10 12 14



Microwave channel 5



## EDA variance assessment with MHS "all sky" mid-tropospheric humidity: WCB events

#### 50 cases

Increased Depar<sup>2</sup> and EnsVar for inflow composite

Difficult to say if EnsVar is reasonable since negative residual is dominated by the overestimation of ObsUnc<sup>2</sup> in WCB region (it is larger than the departures!)

The aim of the diagnostic is to indicate deficiencies and chart our progress towards the ultimate target, but it is not practical to get there immediately

Microwave channel 5



Unit: (K)<sup>2</sup> Mean: -16.5 Sig: 69% -450 -50 -30 -10 10 30 50 210





Observation density (O80, 12h)

Unit: cell<sup>-1</sup>cycle<sup>-1</sup> Mean: 3.95 Sig: 94% 2 4 6 8 10 12 14



Unit: (K)<sup>2</sup> Mean: 21.5 Sig: 92%

10 15 20 25 30 435

## Short-range flow-dependent reliability<sup>†</sup> – a practical route to forecast improvement<sup>‡</sup>

#### Summary, implications and future directions

- Forecast system improvement ≡ Minimisation of a proper score
- Proper Score = Reliability Resolution + Uncertainty
- Red curve (constant forecast) = Black curve (climatological distribution)
  Perfectly reliable (Reliability = 0), but no resolution (Resolution = 0)
- Green and blue forecasts based on (local) partition of phase-space
  If we work to make these reliable (Residual = 0 in EDA budget) then:
  Model uncertainty is likely to be well-represented and ...

Shaded regions  $\Rightarrow$  Resolution > 0  $\Rightarrow$  **More skill** (into medium-range?)

<sup>‡</sup> Practical but not trivial. Need to improve model and representation of model and observation uncertainty

<sup>†</sup> Better (and more numerous) observations will help increase sharpness; Can view as a separate important task

The instantaneous uncertainty growth-rate can be written as:

$$\frac{1}{\sigma_X}\frac{\partial\sigma_X}{\partial t} = \rho_{XX_t}\frac{\sigma_Y}{\sigma_X}$$



My animation showed the  $\hat{\sigma}_{X_t} / \hat{\sigma}_X$  part. Do the random numbers used in some model uncertainty formulations deteriorate the correlation aspect?





Thank you



#### Extra slides



### Reliability in ensemble forecasting



(Cross-terms on squaring have zero expectation. EnsVar is scaled variance to account for finite ensemble-size)



## Reliability in ensemble data assimilation



(Cross-terms on squaring have zero expectation. EnsVar is scaled variance to account for finite ensemble-size)

### Trend in probabilistic forecast performance & Summary



