
ASPECTS OF STRATOSPHERIC ENSEMBLE DATA ASSIMILATION

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Context and Objectives

Context

Fast Chemistry-Climate Model (CCM) tuned for the stratosphere.

First implementation of ensemble data assimilation with a CCM.

perfect twin → imperfect twin → real observations

Possible applications: stratospheric reanalysis, guidelines for operational systems.

Objectives

Test the applicability and possible benefits of ensemble data assimilation to a sparsely-observed, multivariate, nonlinear system like the stratosphere.

Improve the unobserved stratospheric winds (Daley, 1995; Riishojgaard, 1996), through multivariate ensemble data assimilation.

INTRODUCTION

● Context and Objectives

● Outline

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

Outline

INTRODUCTION

● Context and Objectives

● Outline

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

- EnKF-CCM system
- Chemical-dynamical interaction
- Ensemble Kalman smoother

INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: \mathbf{O}_x Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

ENSEMBLE KALMAN FILTERING

with a

CHEMISTRY CLIMATE MODEL

Experimental setup : EnKF

EnKF with perturbed obs (Evensen, 1994; Burgers, 1998)

$$\delta\mathbf{x} = \mathbf{K}_e \mathbf{d}$$

$$\delta\mathbf{x} = \mathbf{x}^a - \mathbf{x}^f = \text{analysis increments}$$

$$\mathbf{d} = \mathbf{y} - \mathcal{H}(\mathbf{x}^f) = \text{observation innovations}$$

$$\mathbf{K}_e = \mathbf{P}_e^f \mathbf{H}^T (\mathbf{H} \mathbf{P}_e^f \mathbf{H}^T + \mathbf{R})^{-1} = \text{Kalman Gain}$$

$$\mathbf{P}_e^f = \frac{1}{M} \sum_{m=1}^M (\mathbf{x}_m^f - \overline{\mathbf{x}^f})(\mathbf{x}_m^f - \overline{\mathbf{x}^f})^T$$

= sample background error-covariance matrix

\mathbf{R} = observations error-covariance matrix (prescribed)

\mathcal{H} = measurement operator

INTRODUCTION

EnKF-CCM

● EnKF Theory

- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: \mathbf{O}_x Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

Experimental Setup : CCM

CHEMISTRY-CLIMATE MODEL (CCM)

IGCM (Forster et al, 2000):

- Multilayer spectral GCM run at T21L26, lid at 0.1 hPa
- Intermediate-complexity physics parametrization
- Prescribed surface temperatures

FASTOC (Taylor and Bourqui, 2005):

- Fast surrogate chemistry scheme
- Based upon comprehensive box model by Fish and Burton (1997), with JPL02 rates.
- Timestep: 24 hrs (diurnal-averaged chemistry)
- Represented catalytic cycles: O_x , HO_x , NO_x .
- Advected species: O_x , N_2O_5 , NO_x , HNO_3

INTRODUCTION

EnKF-CCM

- EnKF Theory
- **Chemistry-Climate Model**
- Filter Configurations
- Observations
- Localization
- Optimal Localization: T Assimilation
- Optimal Localization: O_x Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

Experimental Setup : Filter Configurations

INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: T Assimilation
- Optimal Localization: O_x Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

model state
vector

$$\begin{pmatrix} \mathbf{u} \\ \mathbf{v} \\ T \\ P_s \\ Q \\ O_x \\ N_2O_5 \\ NO_x \\ HNO_3 \end{pmatrix}$$

- Perfect-twin experiment
- Initial ensemble is climatological with 128 members (Jan 1st of each year)
- Sequential Double-EnKF assimilation of observations by batches (Houtekamer & Mitchell, 2001)
- Separate horizontal and vertical covariance localization parameters for ozone and temperature covariances
- No covariance inflation
- Analysis performed every 24 hours

Experimental Setup : Observations

INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations

● Observations

- Localization
- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: \mathbf{O}_x Assimilation
- Optimal Simulations

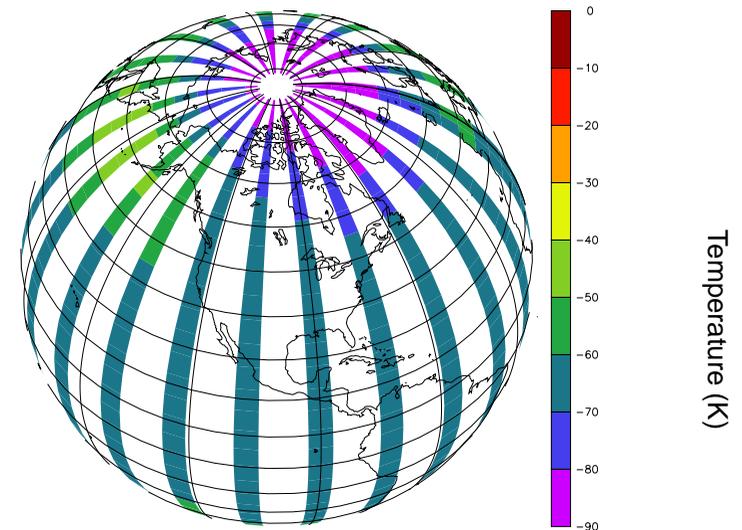
CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

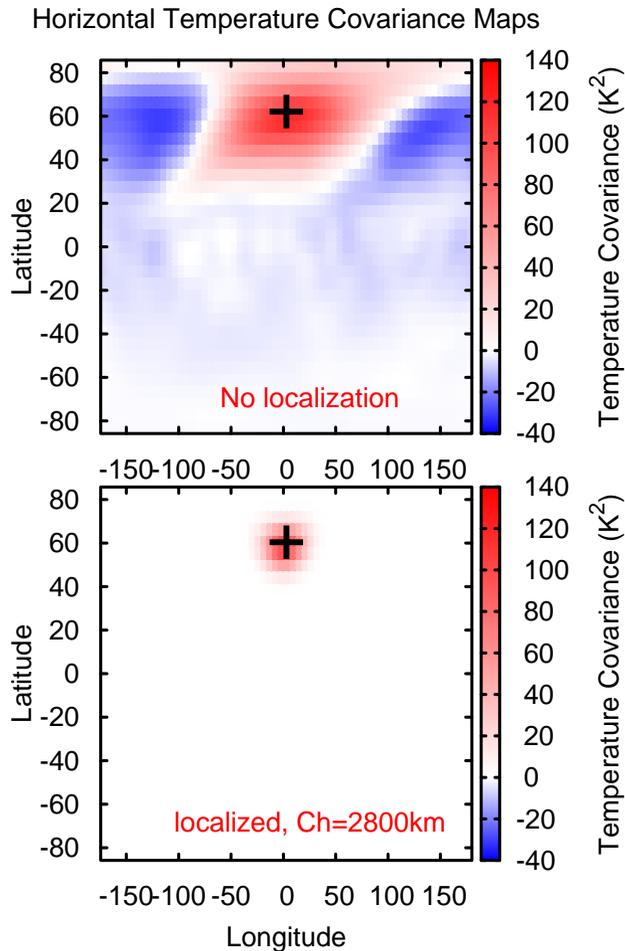
- Synthetic MIPAS-like **temperature** retrievals with 2K error
- Synthetic MIPAS-like **ozone** retrievals with 10% error
- Diagonal \mathbf{R} matrix
- Obs instantaneous at 00UTC
- Vertical coverage between 4hPa and 200hPa on pressure levels

- Horizontal coverage on model grid points :



Localization

Localization: $\rho_v \circ \rho_h \circ \mathbf{P}^f$



+ Increase rank of background error-covariance matrix.

+ Remove (far-away) sampling noise.

- Lose the natural anisotropy : risk of introducing imbalance.

→ Ideally, find optimal decorrelation length

INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: \mathbf{O}_x Assimilation
- Optimal Simulations

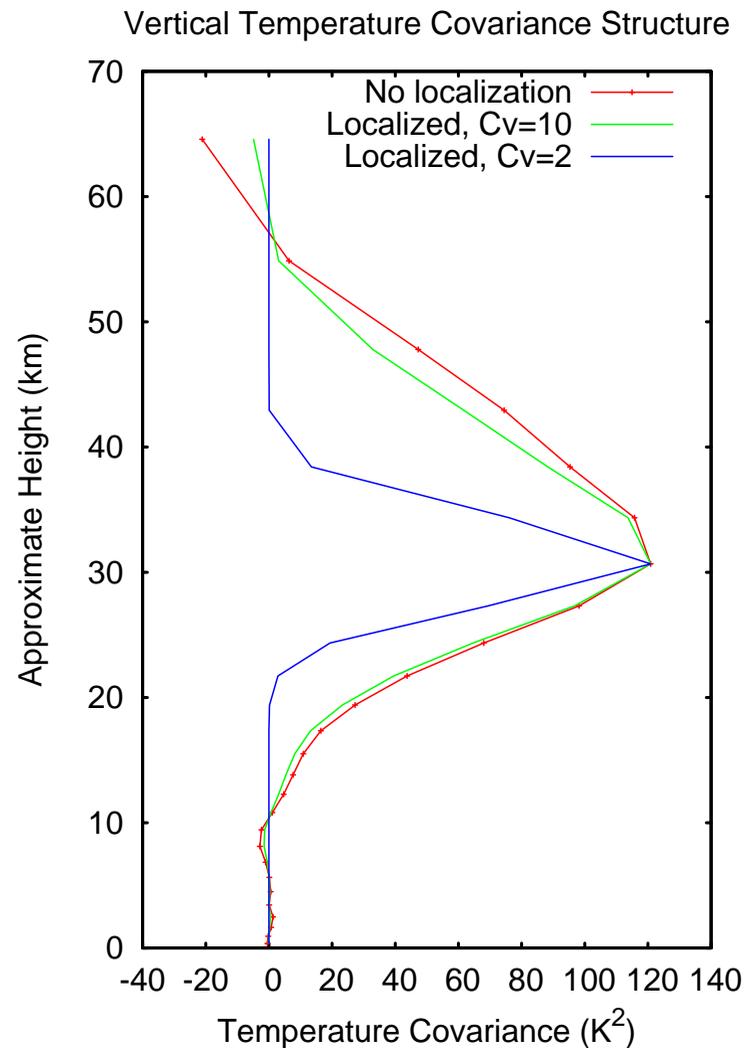
CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

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INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations

● Localization

- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: \mathbf{O}_x Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

Diagnostics: RMSE and SPREAD

INTRODUCTION

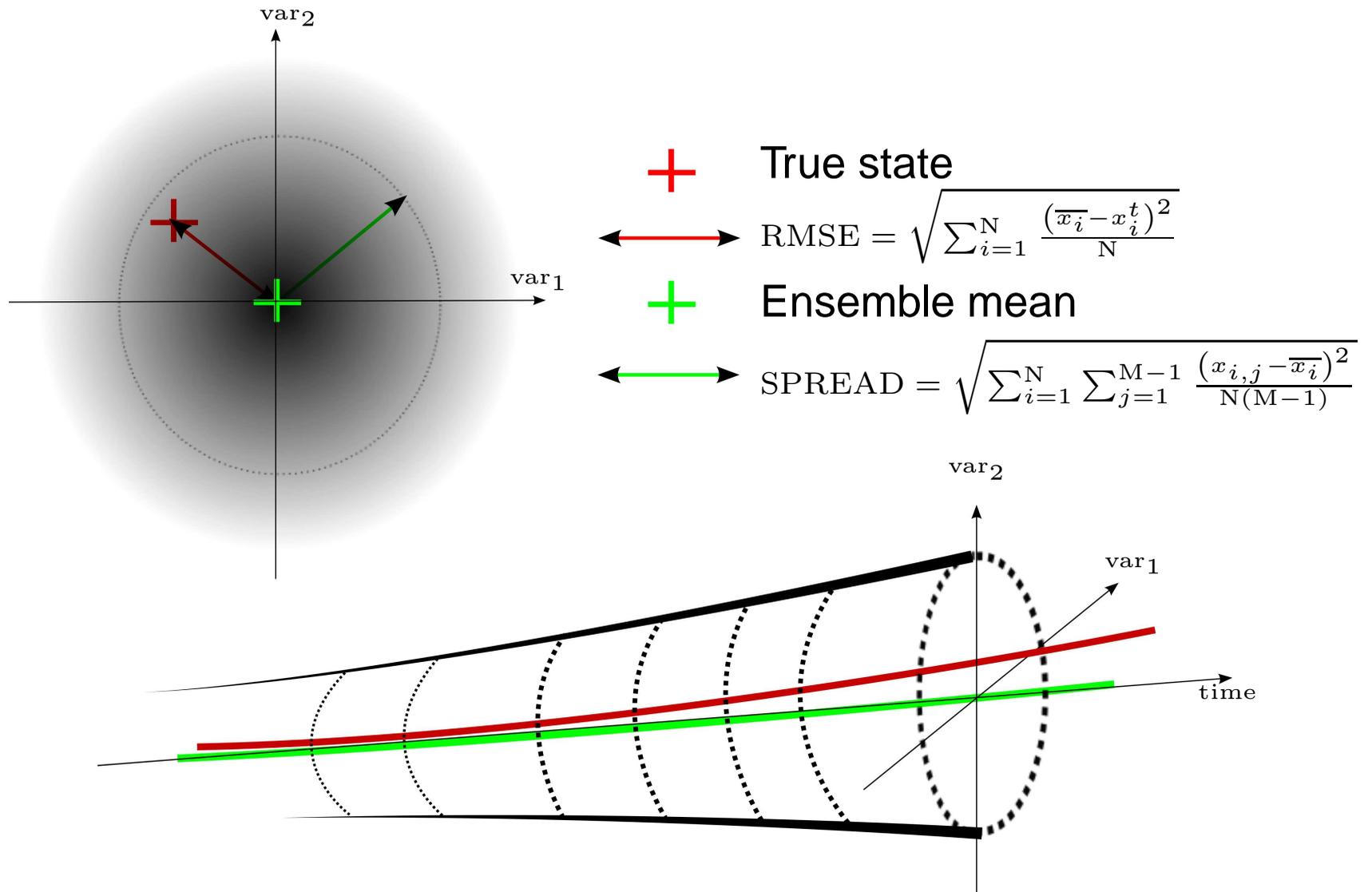
EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T}
- Assimilation
- Optimal Localization: \mathbf{O}_x
- Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

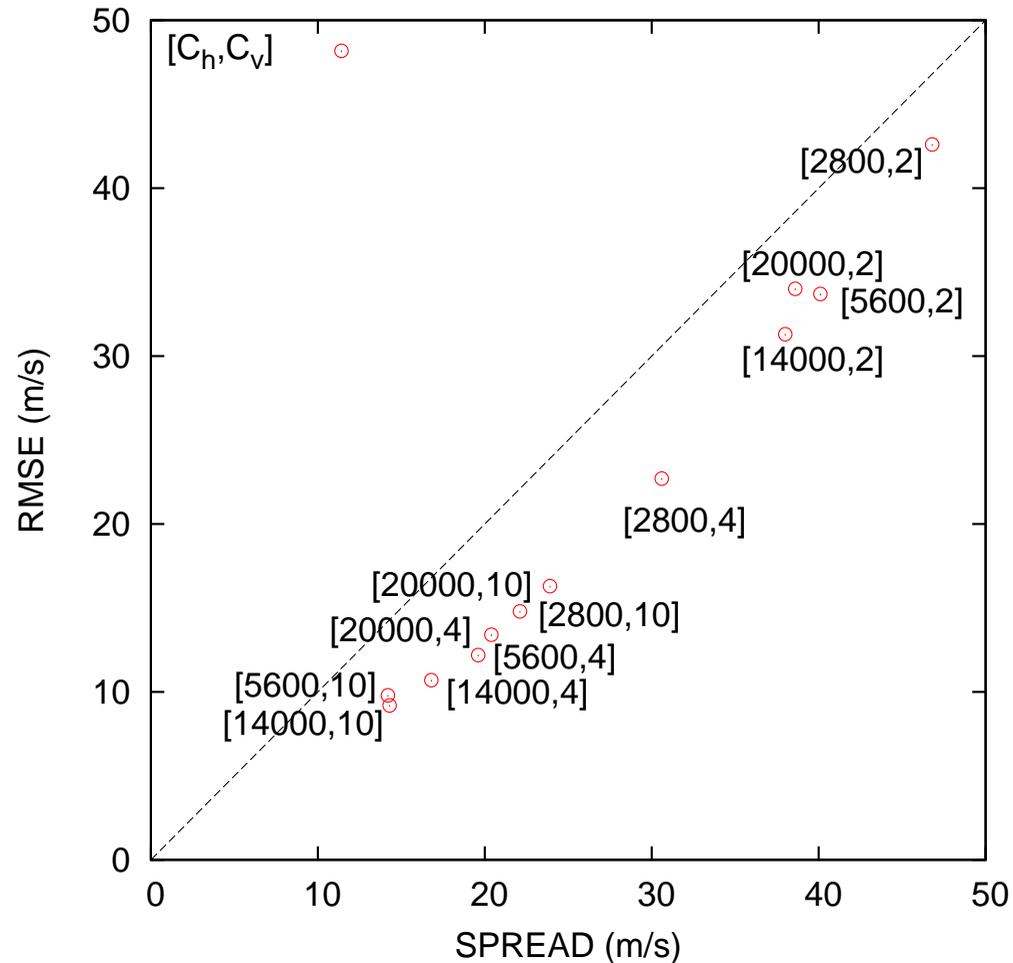
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CONCLUSIONS AND FUTURE WORK



Optimal Localization: T Assimilation

Temperature assimilation localization sensitivity study :
Time-averaged square-root of TE RMSE and SPREAD for various localization scenarios



$$TE = \frac{u'^2}{2} + \frac{v'^2}{2} + \frac{C_p}{T_{ref}} T'^2 + R_\alpha T_{ref} \left(\frac{P'_s}{P_{ref}} \right)^2$$

INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: T
- Assimilation
- Optimal Localization: O_x
- Assimilation
- Optimal Simulations

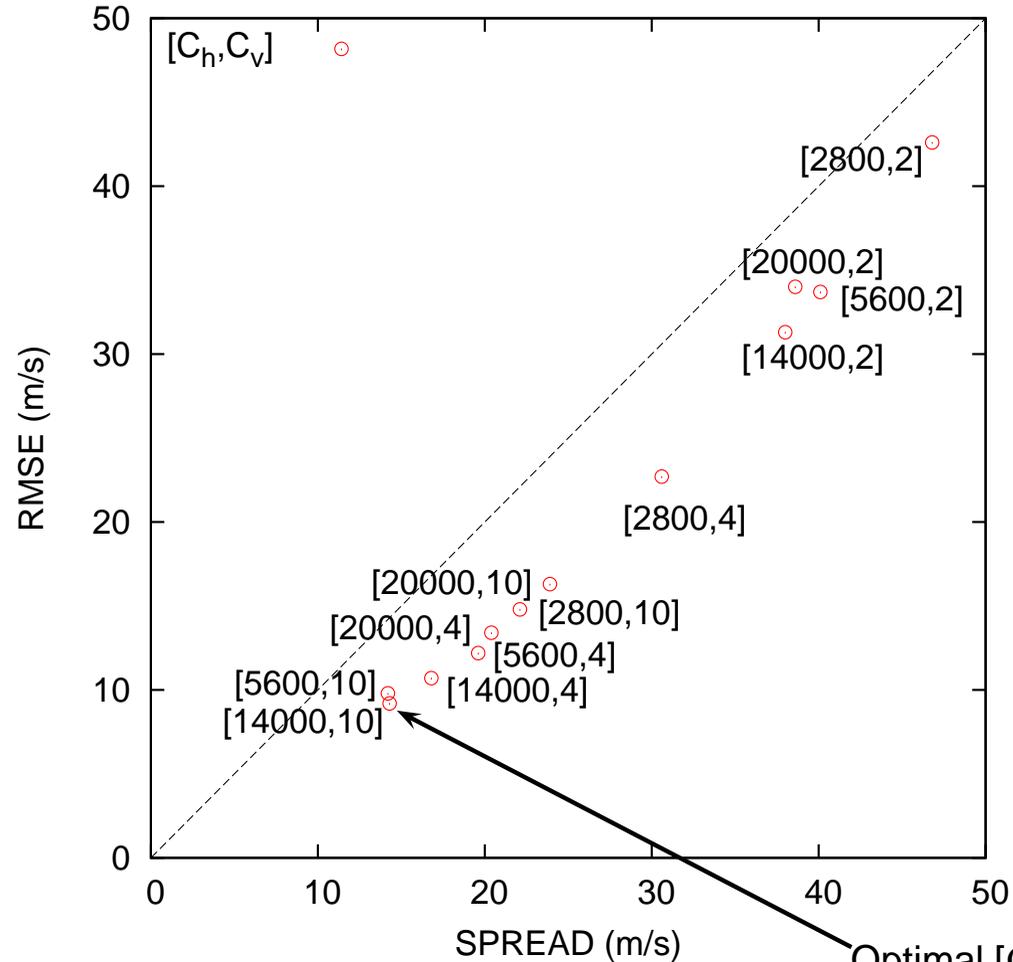
CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

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INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization

● Optimal Localization: T

Assimilation

- Optimal Localization: O_x
- Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS

INTERACTION

ENSEMBLE KALMAN

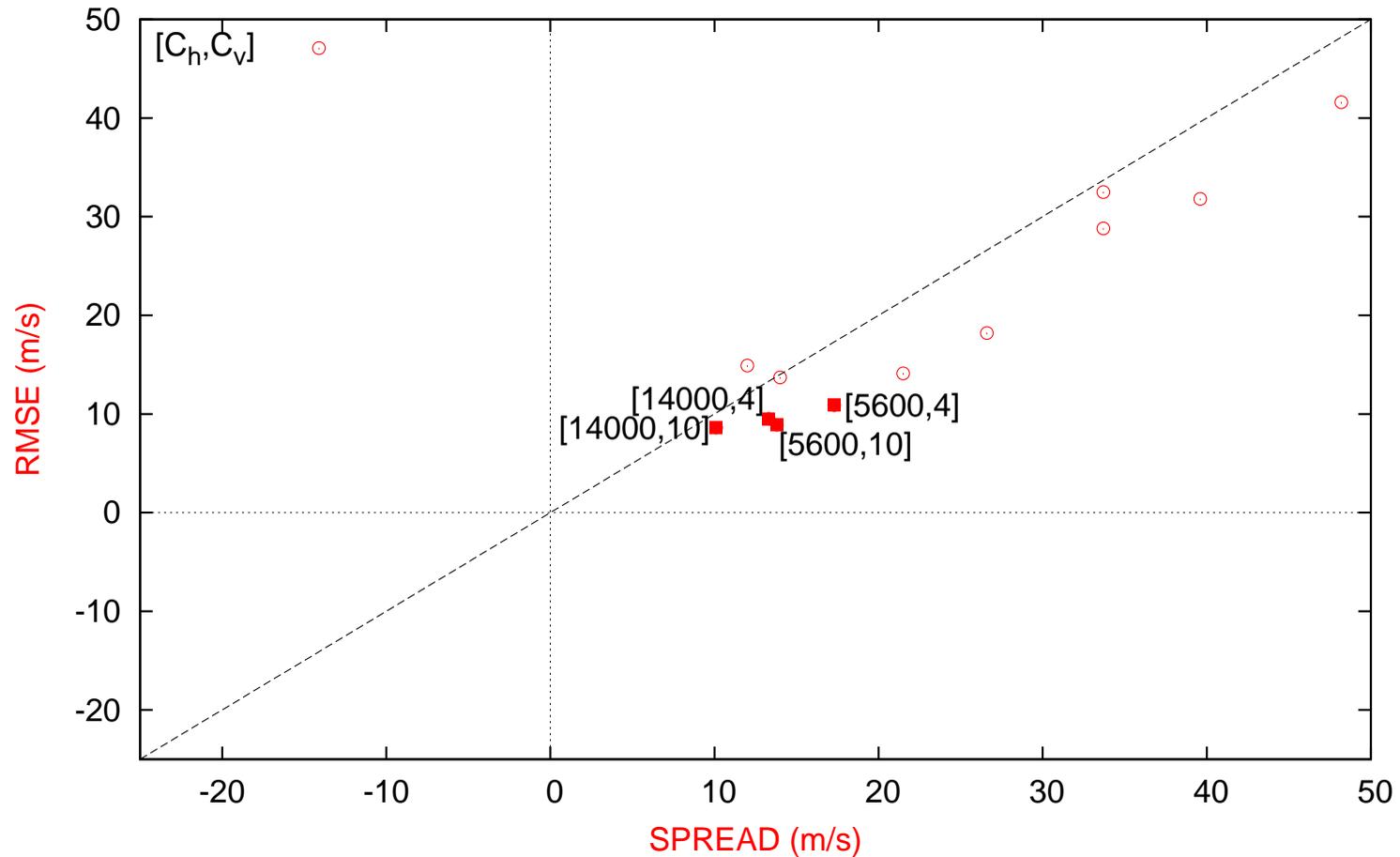
SMOOTHER

CONCLUSIONS AND FUTURE

WORK

Optimal Localization: O_x Assimilation

Ozone assimilation localization sensitivity study :
Time-averaged square-root of TE RMSE and SPREAD, alpha and beta for various localization scenarios



INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: T Assimilation
- Optimal Localization: O_x Assimilation
- Optimal Simulations

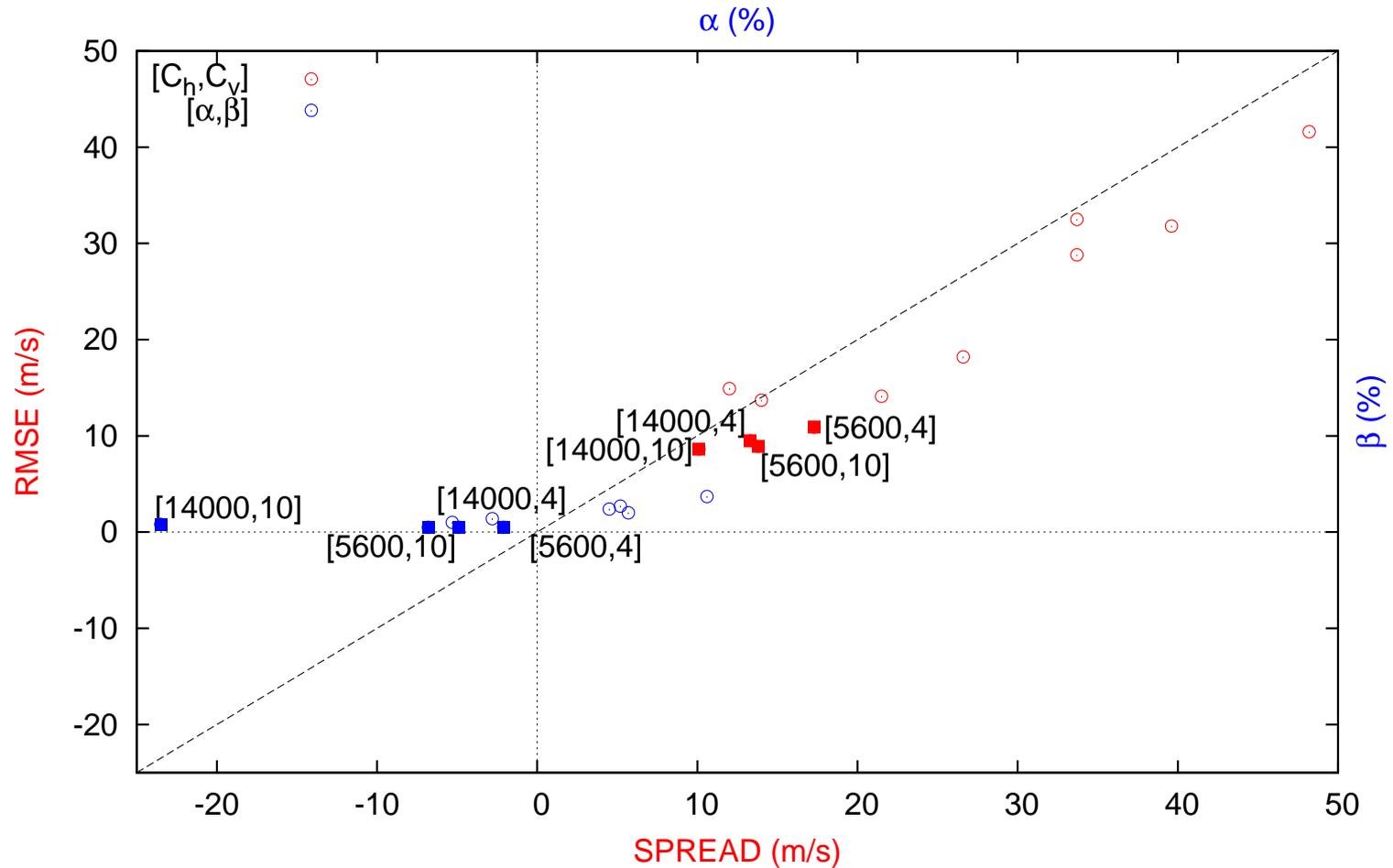
CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

Optimal Localization: O_x Assimilation

Ozone assimilation localization sensitivity study :
Time-averaged square-root of TE RMSE and SPREAD, alpha and beta for various localization scenarios



$$\alpha = \left(\frac{\text{Tr}(\langle (\mathcal{H}(\mathbf{x}^a) - \mathcal{H}(\mathbf{x}^f))(\mathbf{y} - \mathcal{H}(\mathbf{x}^f))^T \rangle)}{\text{Tr}(\mathbf{H}\mathbf{P}_e^f\mathbf{H}^T)} - 1 \right) \cdot 100\%, \quad \beta = \left(\frac{\text{Tr}(\langle (\mathbf{y} - \mathcal{H}(\mathbf{x}^a))(\mathbf{y} - \mathcal{H}(\mathbf{x}^f))^T \rangle)}{\text{Tr}(\mathbf{R})} - 1 \right) \cdot 100\%$$

INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation

● Optimal Localization: O_x Assimilation

- Optimal Simulations

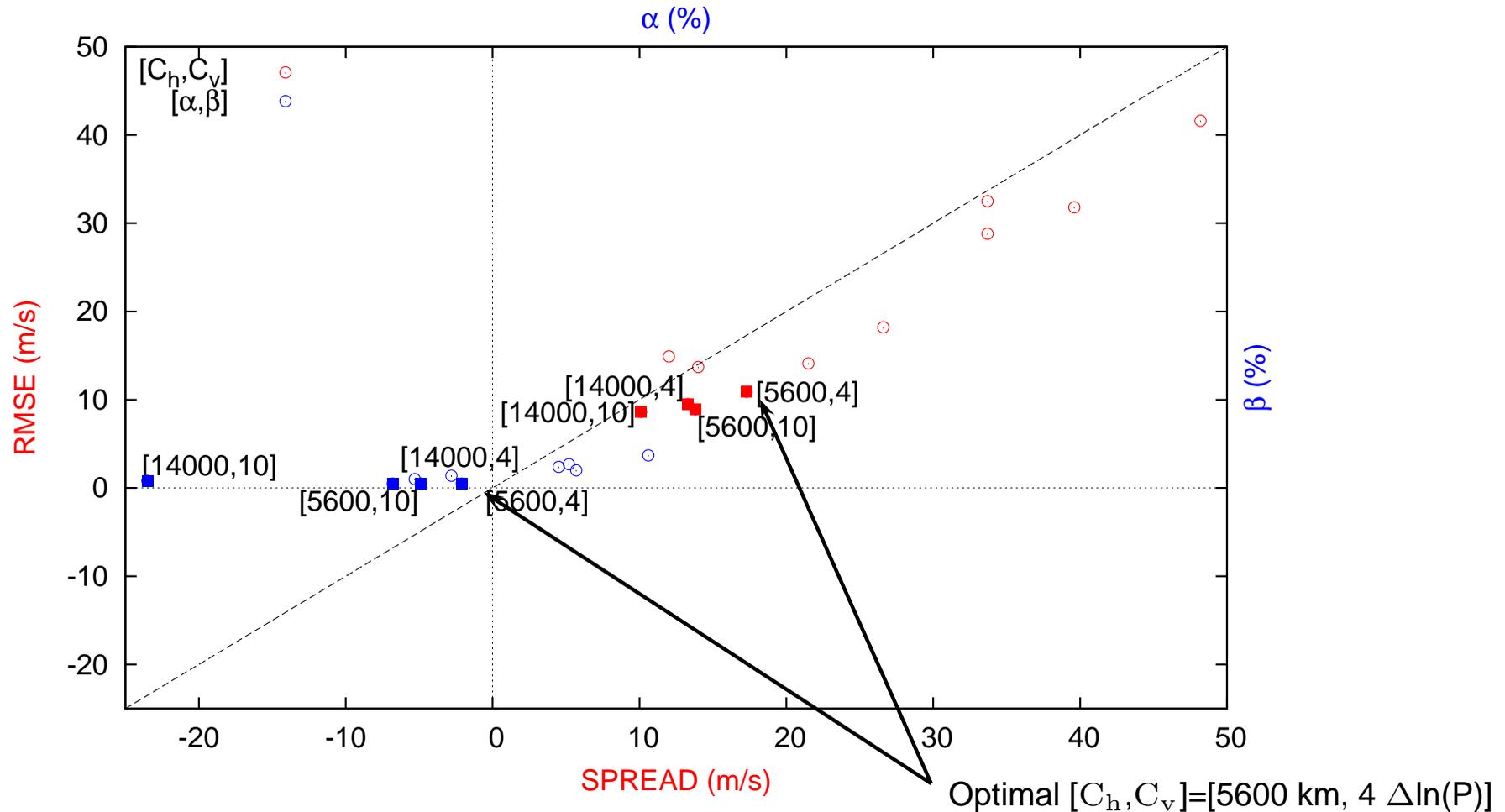
CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

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INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation

● Optimal Localization: O_x Assimilation

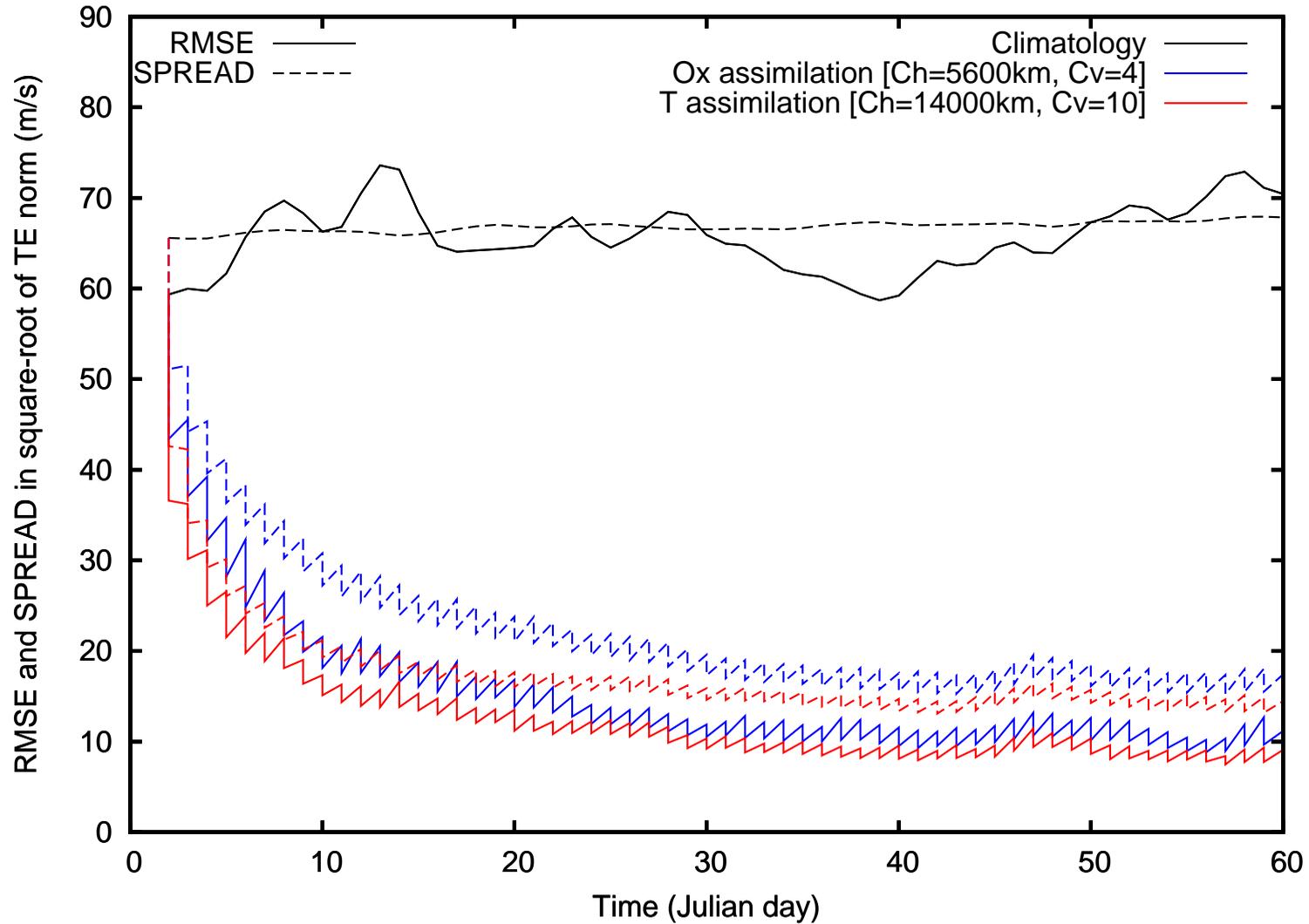
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

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CONCLUSIONS AND FUTURE WORK

Evolution of Optimal Simulation: Total Energy



$$TE = \frac{u'^2}{2} + \frac{v'^2}{2} + \frac{C_p}{T_{ref}} T'^2 + R_\alpha T_{ref} \left(\frac{P'_s}{P_{ref}} \right)^2$$

INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: O_x Assimilation

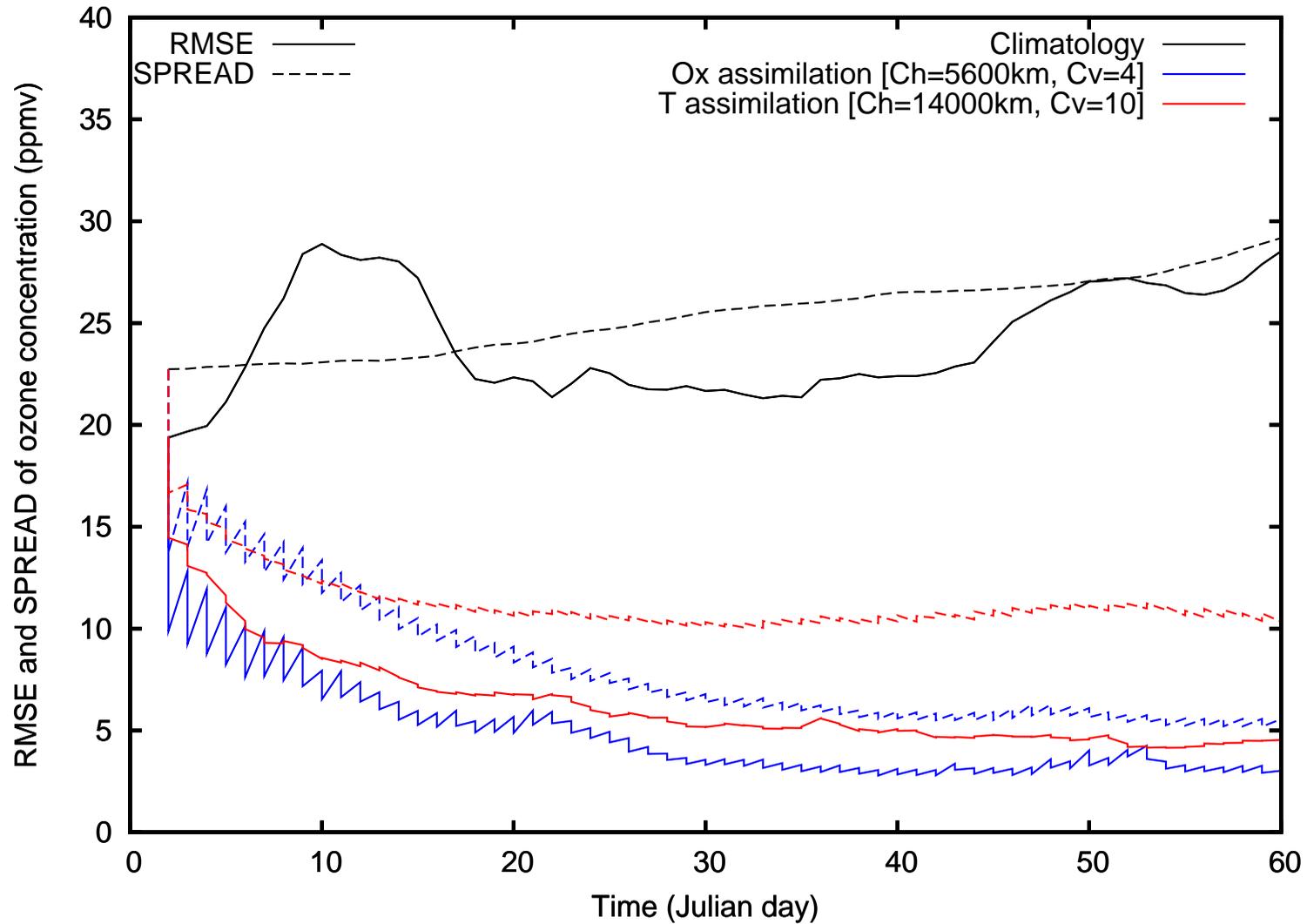
● Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

Evolution of Optimal Simulation: Ozone



INTRODUCTION

EnKF-CCM

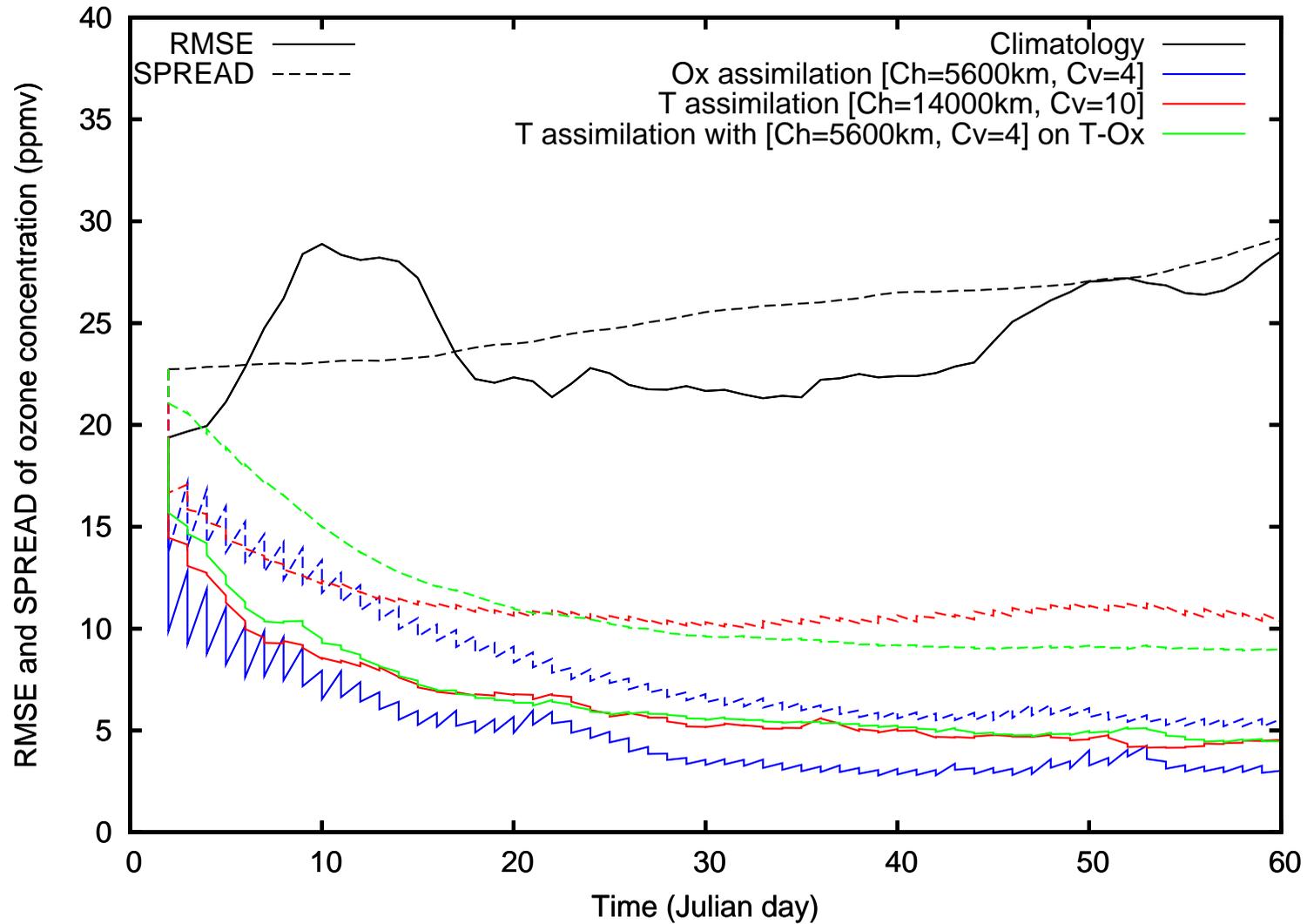
- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: \mathbf{O}_x Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

Evolution of Optimal Simulation: Ozone



INTRODUCTION

EnKF-CCM

- EnKF Theory
- Chemistry-Climate Model
- Filter Configurations
- Observations
- Localization
- Optimal Localization: \mathbf{T} Assimilation
- Optimal Localization: \mathbf{O}_x Assimilation
- Optimal Simulations

CHEMISTRY-DYNAMICS INTERACTION

ENSEMBLE KALMAN SMOOTHER

CONCLUSIONS AND FUTURE WORK

INTRODUCTION

EnKF-CCM

**CHEMISTRY-DYNAMICS
INTERACTION**

- Experiments
- **T** Assimilation:
Time-Averaged Global Analyses
- Schematics
- **O_x** Assimilation:
Time-Averaged Global Analyses
- Zonally-Averaged Analyses

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CONCLUSIONS AND FUTURE
WORK

CHEMISTRY-DYNAMICS INTERACTION

Experiments

Experiments

T assimilation	O_x assimilation
<p>Control</p> <p>T obs transmit their information to all variables</p>	<p>Control</p> <p>O_x obs transmit their information to all variables</p>
<p>NoChem</p> <p>T obs transmit their information only to u, v, T, q and P_s</p>	<p>NoDyn</p> <p>O_x obs transmit their information only to O_x, N₂O₅, NO_x, HNO₃ and P_s</p>
	<p>NoTemp</p> <p>O_x obs transmit their information to all variables except T</p>
	<p>NoWinds</p> <p>O_x obs transmit their information to all variables except u and v</p>

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

● Experiments

● T Assimilation:

Time-Averaged Global Analyses

● Schematics

● O_x Assimilation:

Time-Averaged Global Analyses

● Zonally-Averaged Analyses

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

T Assimilation: Time-Averaged Global Analyses

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS

INTERACTION

● Experiments

● **T** Assimilation:
Time-Averaged Global Analyses

● Schematics

● O_x Assimilation:

Time-Averaged Global Analyses

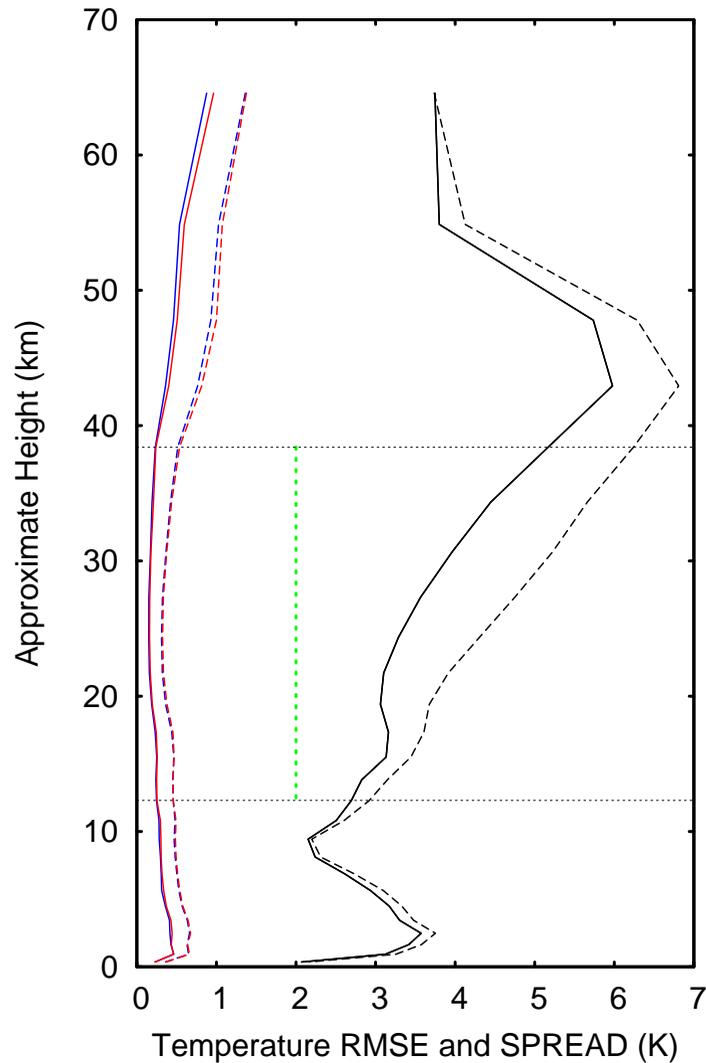
● Zonally-Averaged Analyses

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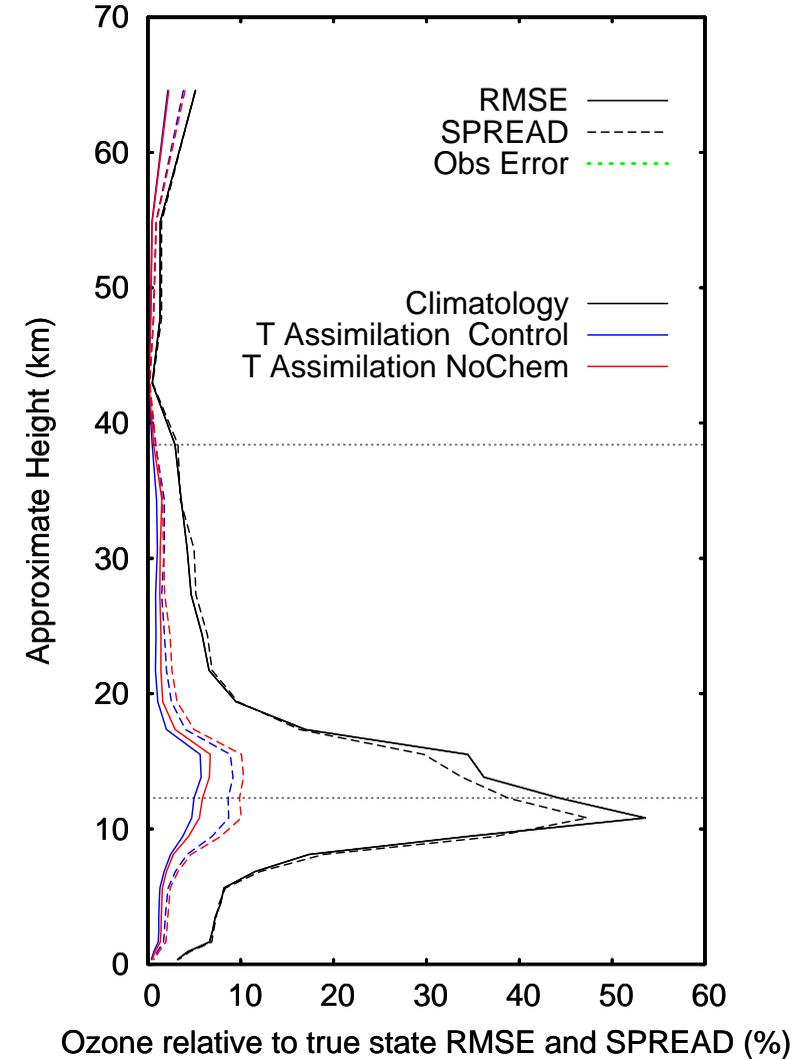
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CONCLUSIONS AND FUTURE

WORK



→ useless T-Chem error covariances
or ineffective radiation ?



→ T-Chem error covariances
further reduce O_x error.

Schematics

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

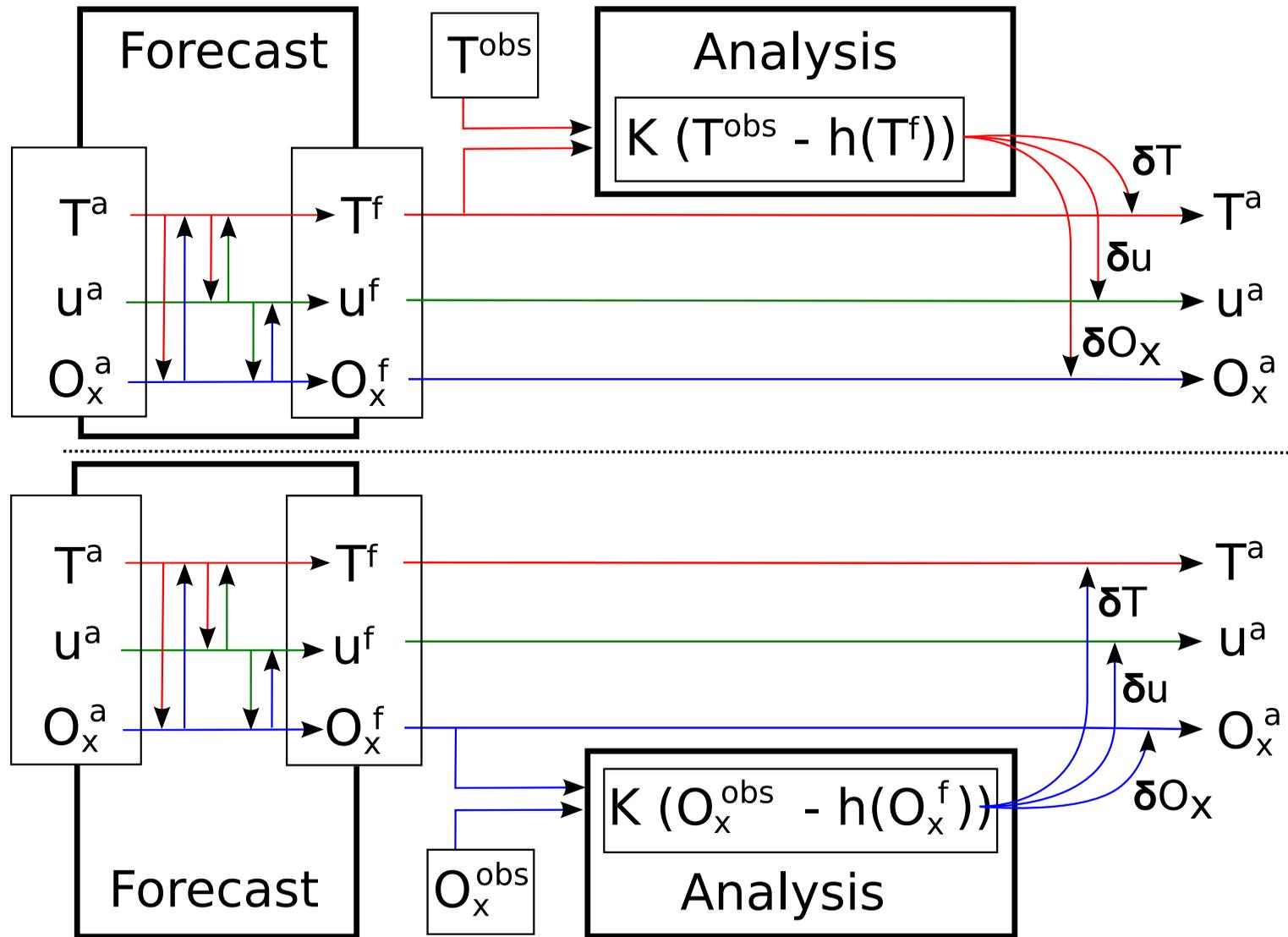
● Experiments
● \mathbf{T} Assimilation:
Time-Averaged Global Analyses

● Schematics

● \mathbf{O}_x Assimilation:
Time-Averaged Global Analyses
● Zonally-Averaged Analyses

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SMOOTHER

CONCLUSIONS AND FUTURE
WORK



O_x Assimilation: Time-Averaged Global Analyses

- INTRODUCTION

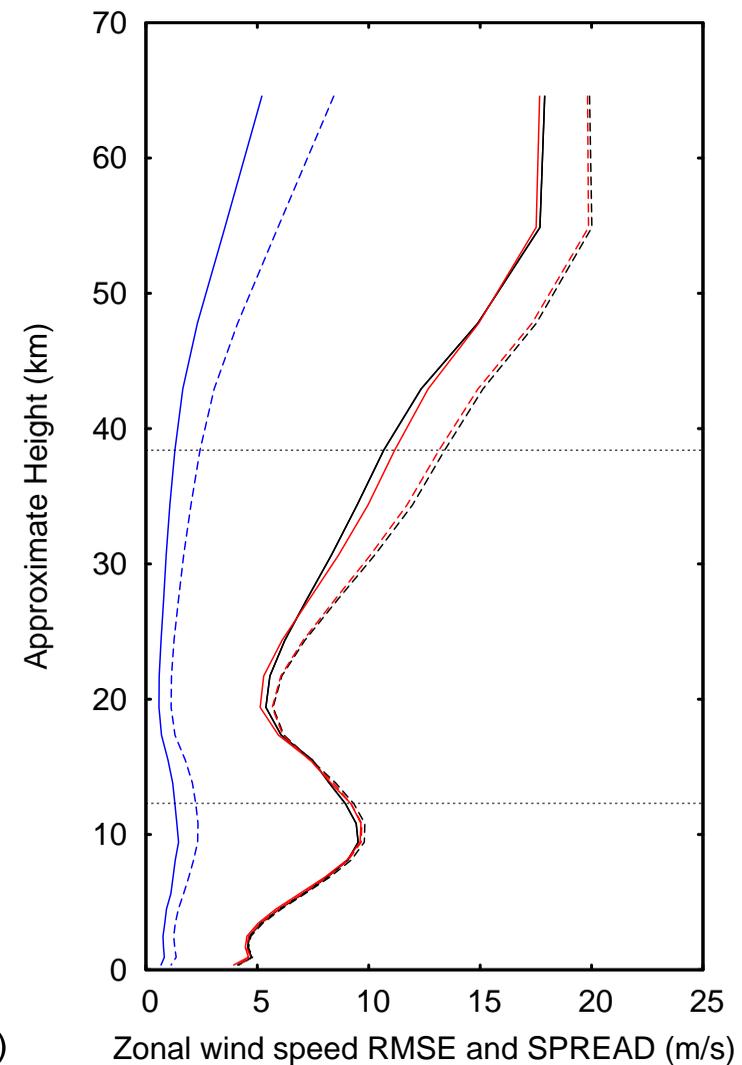
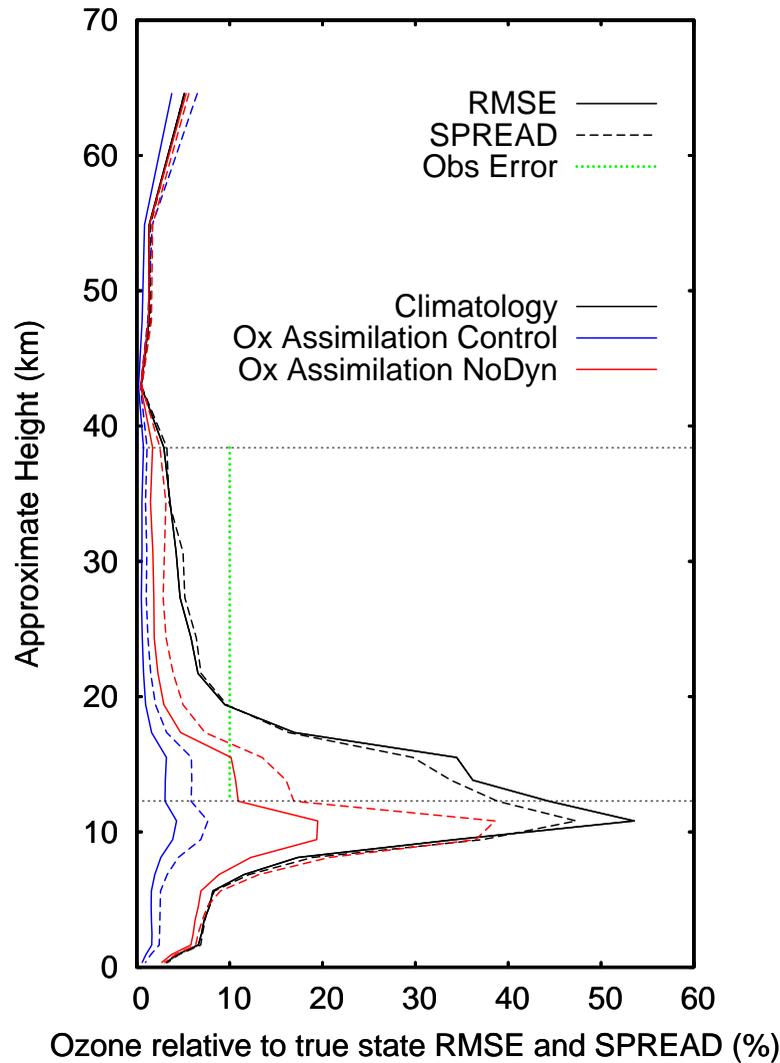
- EnKF-CCM

- CHEMISTRY-DYNAMICS INTERACTION

- Experiments
- **T** Assimilation: Time-Averaged Global Analyses
- Schematics
- **O_x** Assimilation: Time-Averaged Global Analyses
- Zonally-Averaged Analyses

- ENSEMBLE KALMAN SMOOTHER

- CONCLUSIONS AND FUTURE WORK



→ O_x-Dyn error covariances account for about half the O_x error reduction.

→ O_x-Dyn covariances are necessary to constrain well the u error.

O_x Assimilation: Time-Averaged Global Analyses

- INTRODUCTION

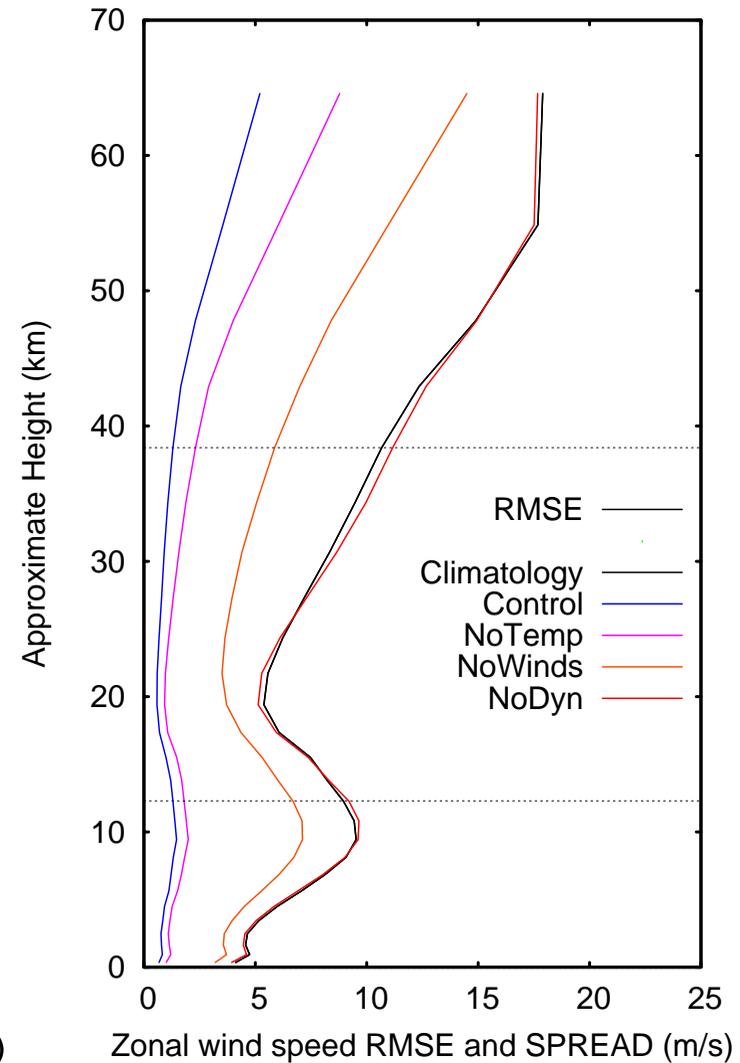
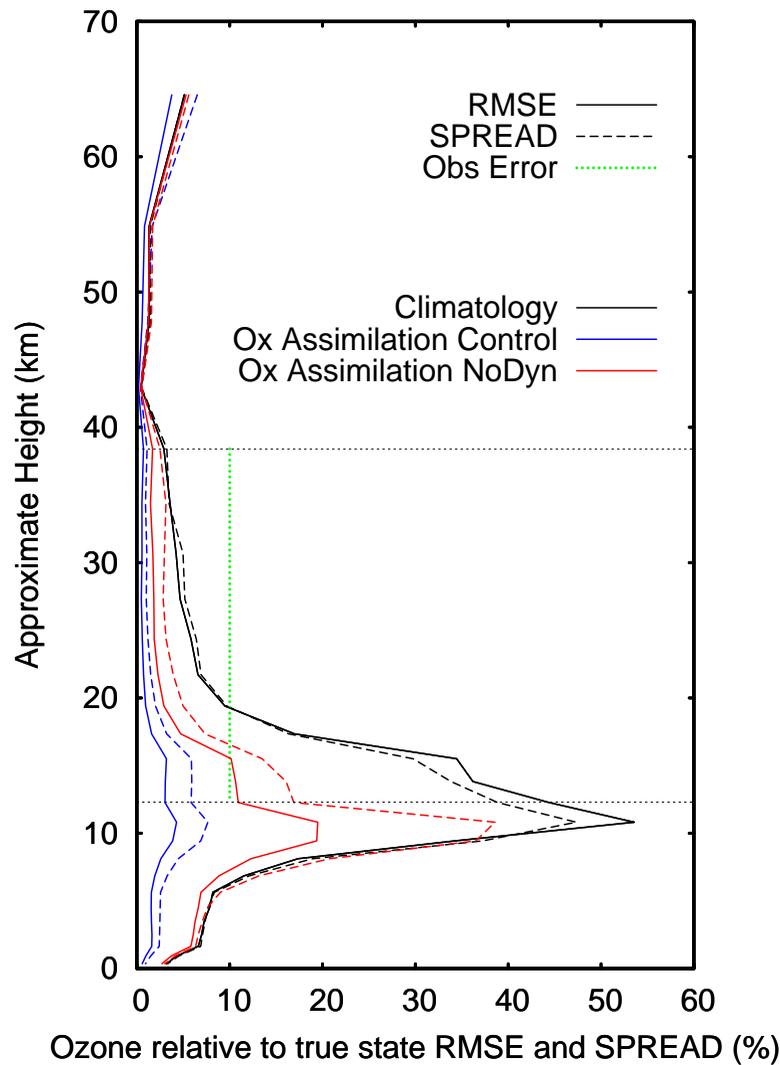
- EnKF-CCM

- CHEMISTRY-DYNAMICS INTERACTION

- Experiments
- **T** Assimilation: Time-Averaged Global Analyses
- Schematics
- **O_x** Assimilation: Time-Averaged Global Analyses
- Zonally-Averaged Analyses

- ENSEMBLE KALMAN SMOOTHER

- CONCLUSIONS AND FUTURE WORK

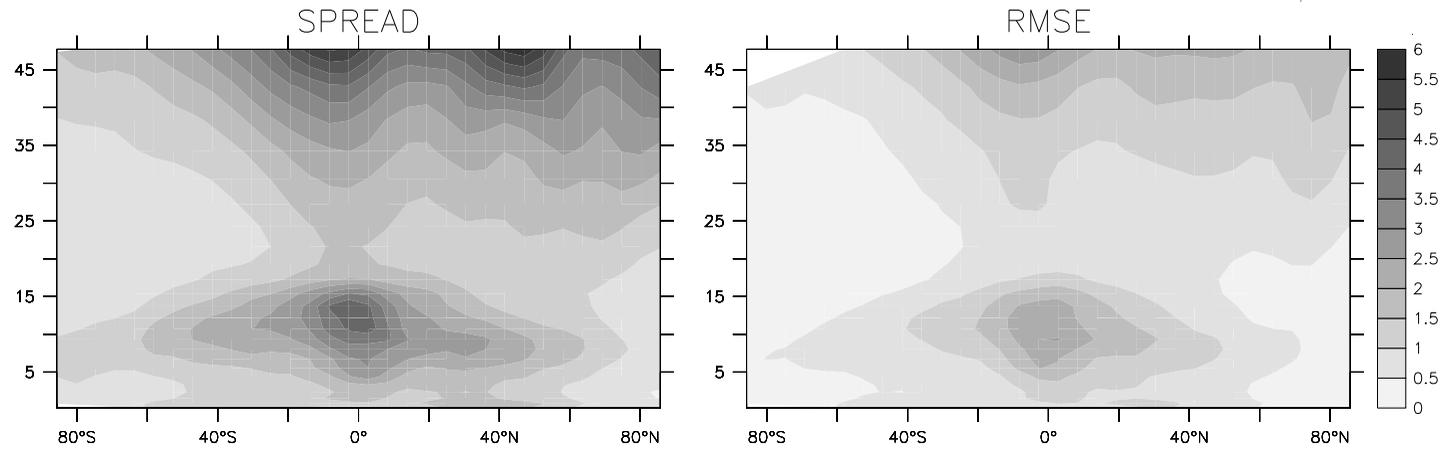


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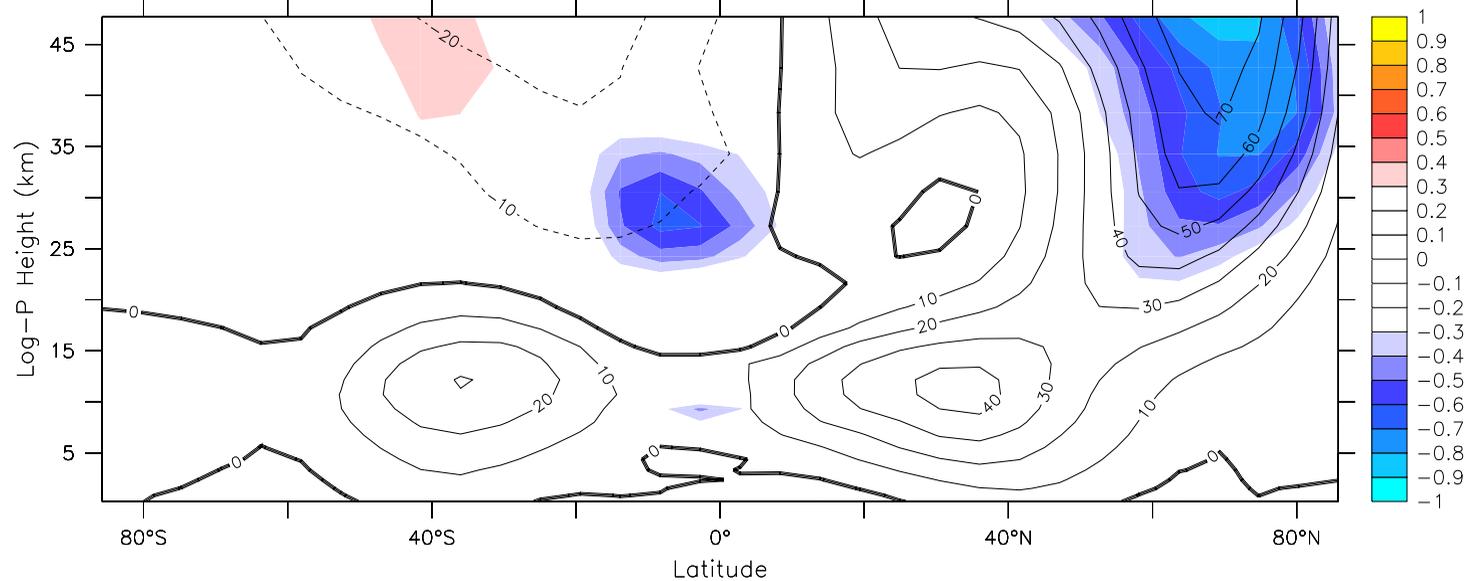
→ O_x-Wind error covariances are the most critical to constrain the dynamics.

Zonally-Averaged Zonal Wind Analyses

Control O_x assimilation ZONAL WIND (m/s)



ENSEMBLE-AVERAGE and BIAS



INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

● Experiments
● **T** Assimilation:
Time-Averaged Global Analyses

● Schematics
● **O_x** Assimilation:
Time-Averaged Global Analyses

● **Zonally-Averaged Analyses**

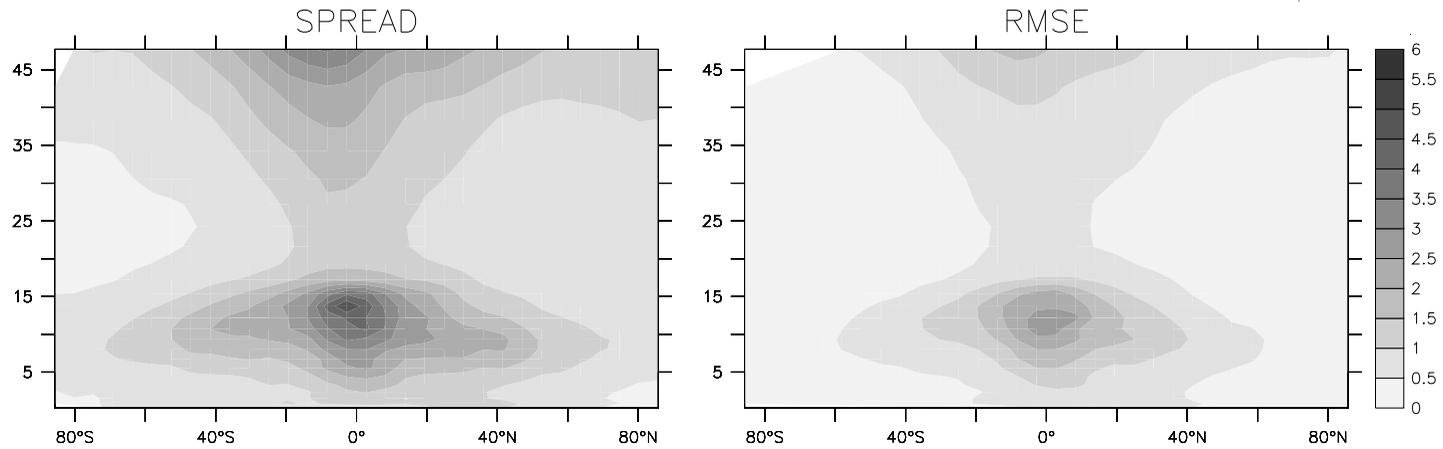
ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

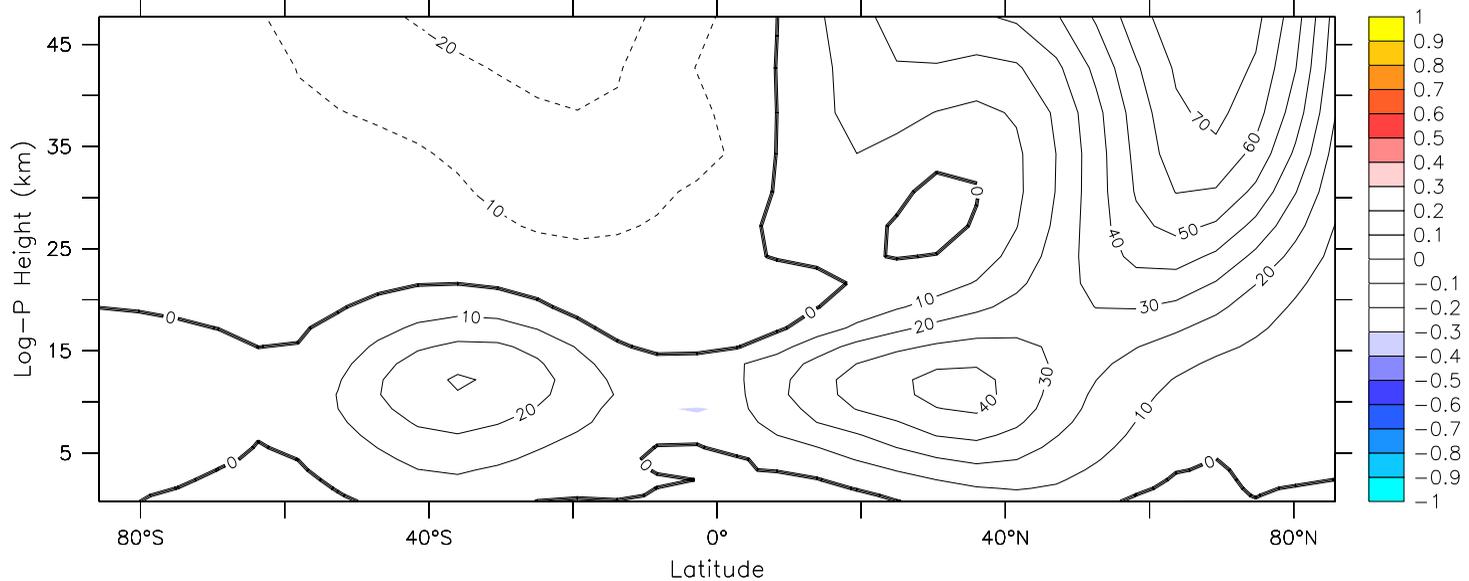
Zonally-Averaged Zonal Wind Analyses

Control T assimilation

ZONAL WIND (m/s)



ENSEMBLE-AVERAGE and BIAS



INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

● Experiments
● T Assimilation:
Time-Averaged Global Analyses

● Schematics
● O_x Assimilation:
Time-Averaged Global Analyses

● Zonally-Averaged Analyses

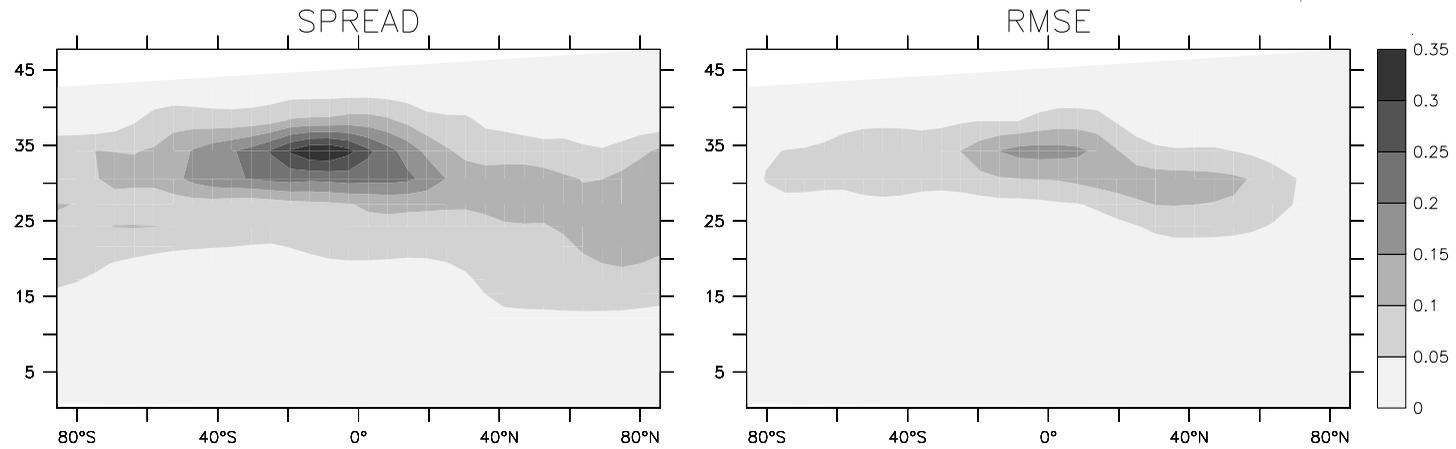
ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

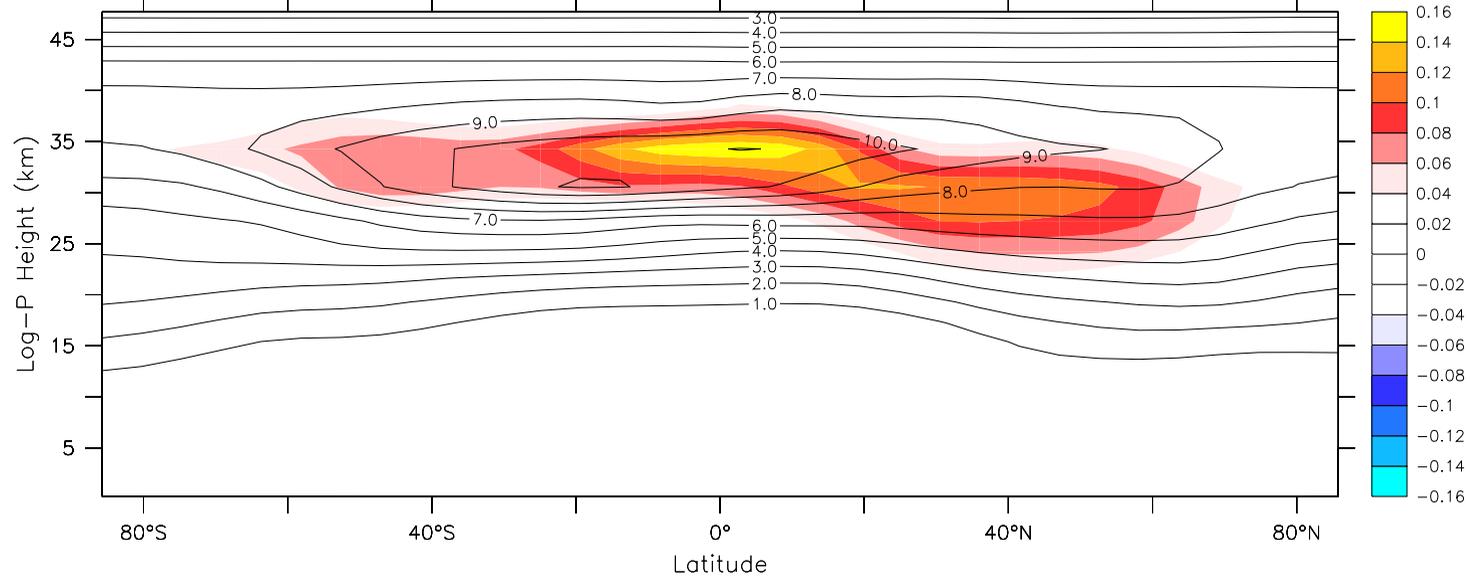
Zonally-Averaged Ozone Analyses

Control T assimilation

O_x (ppmv)



ENSEMBLE-AVERAGE and BIAS



INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

● Experiments
● T Assimilation:
Time-Averaged Global Analyses

● Schematics
● O_x Assimilation:
Time-Averaged Global Analyses

● Zonally-Averaged Analyses

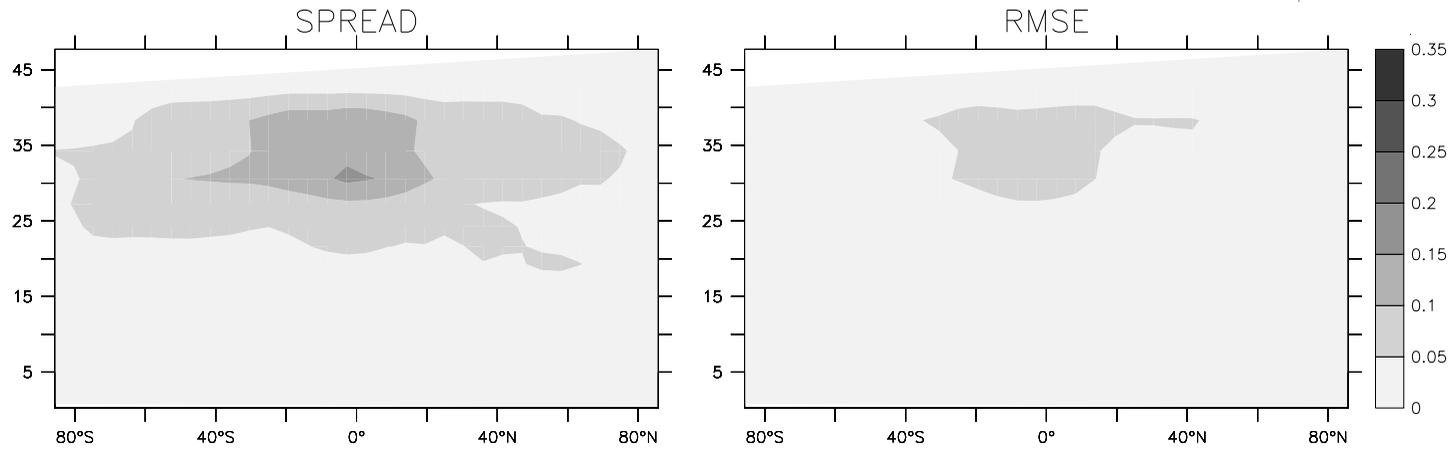
ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

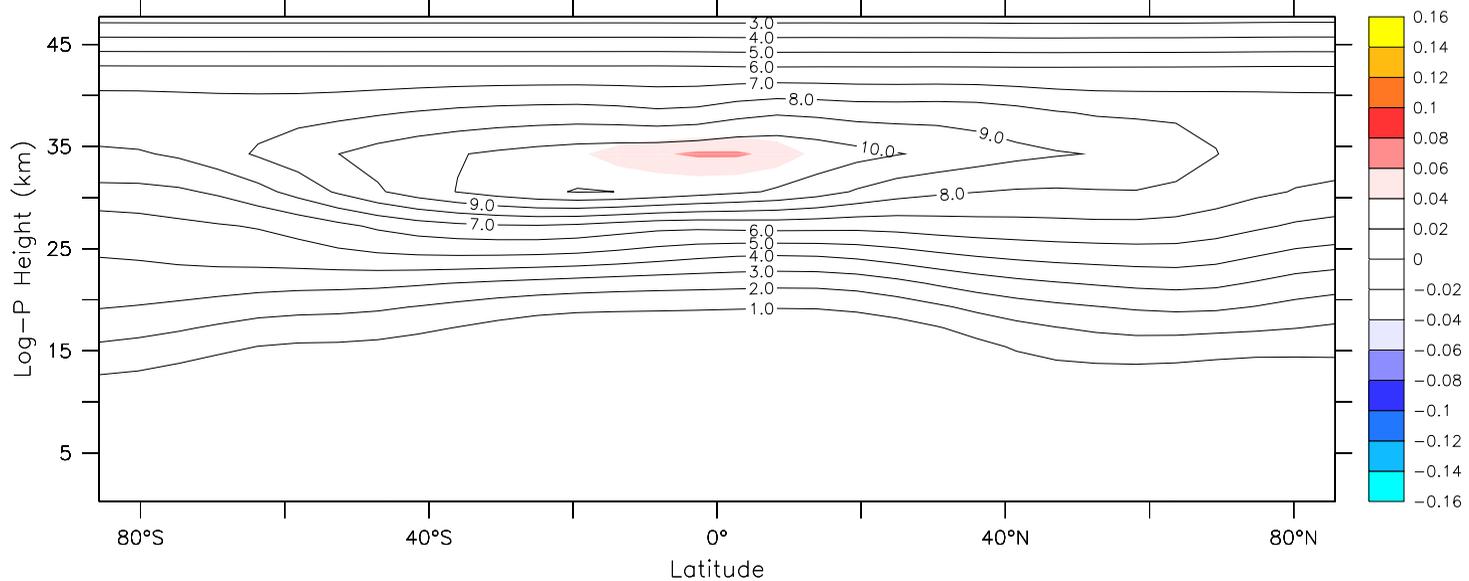
Zonally-Averaged Ozone Analyses

Control O_x assimilation

O_x (ppmv)



ENSEMBLE-AVERAGE and BIAS



INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

● Experiments
● **T** Assimilation:
Time-Averaged Global Analyses

● Schematics
● **O_x** Assimilation:
Time-Averaged Global Analyses

● **Zonally-Averaged Analyses**

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

**ENSEMBLE KALMAN
SMOOTHER**

- EnKS Configurations
- EnKS Analyses
- Experiment
- EnKS vs. EnKF

CONCLUSIONS AND FUTURE
WORK

ENSEMBLE KALMAN SMOOTHER

EnKS Configurations

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

● EnKS Configurations

● EnKS Analyses

● Experiment

● EnKS vs. EnKF

CONCLUSIONS AND FUTURE
WORK

model state
vector

$$\begin{pmatrix} u \\ v \\ T \\ P_s \\ Q \\ O_x \\ N_2O_5 \\ NO_x \\ HNO_3 \end{pmatrix}$$

- Perfect-model twin experiment
- Initial ensemble is **ensemble of analyses from Control Mipas O_x and Mipas T assimilation experiments, every three days in February**
- Double-**EnKS** assimilation of a **single-batch** of daily observations
- **Compressed Row Storage (CRS)** for **sparse background error-covariance matrix**
- No covariance inflation

O_x Assimilation: Time-Lagged EnKS Analyses

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

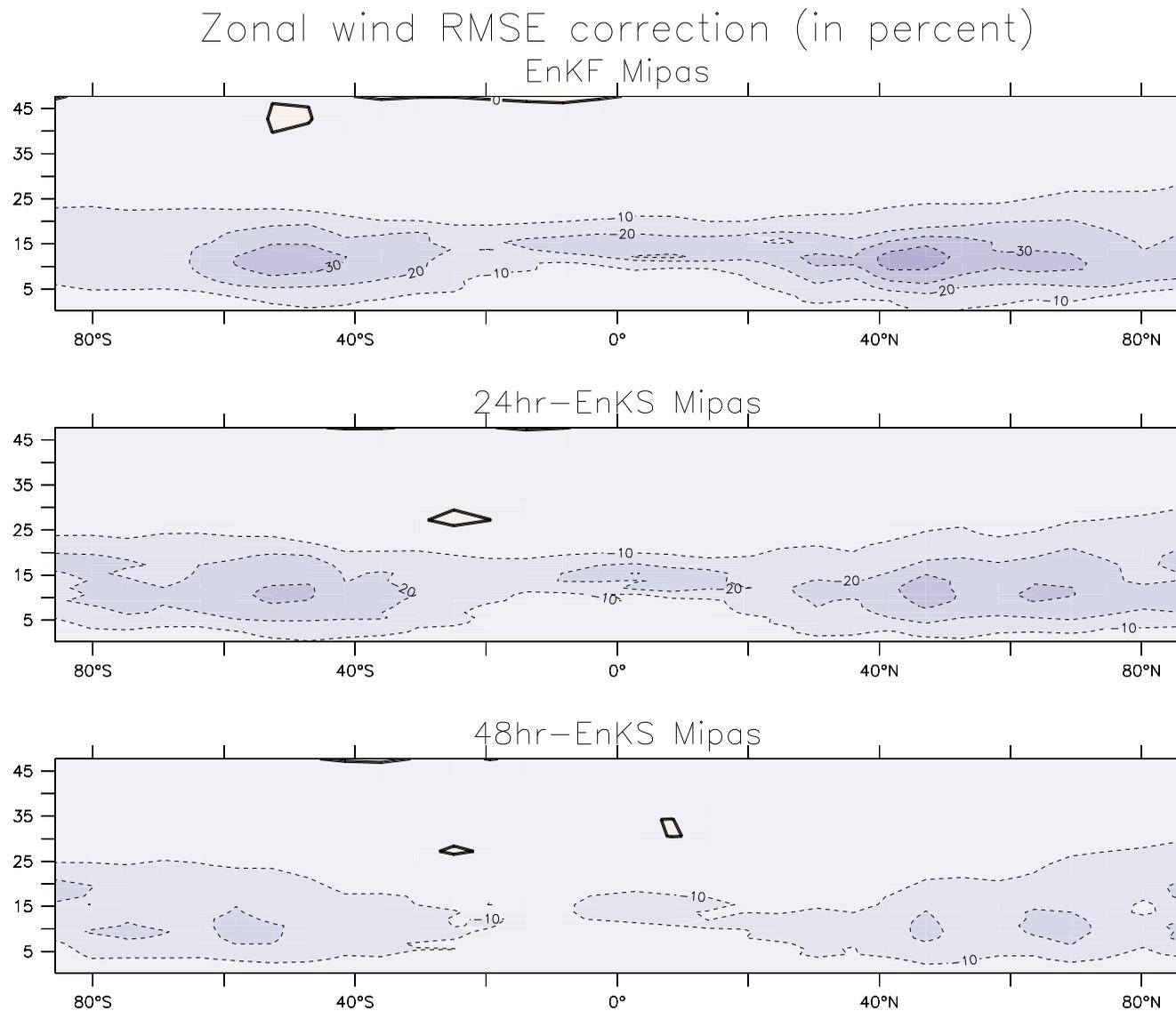
● EnKS Configurations

● EnKS Analyses

● Experiment

● EnKS vs. EnKF

CONCLUSIONS AND FUTURE
WORK



O_x Assimilation: Time-Lagged EnKS Analyses

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

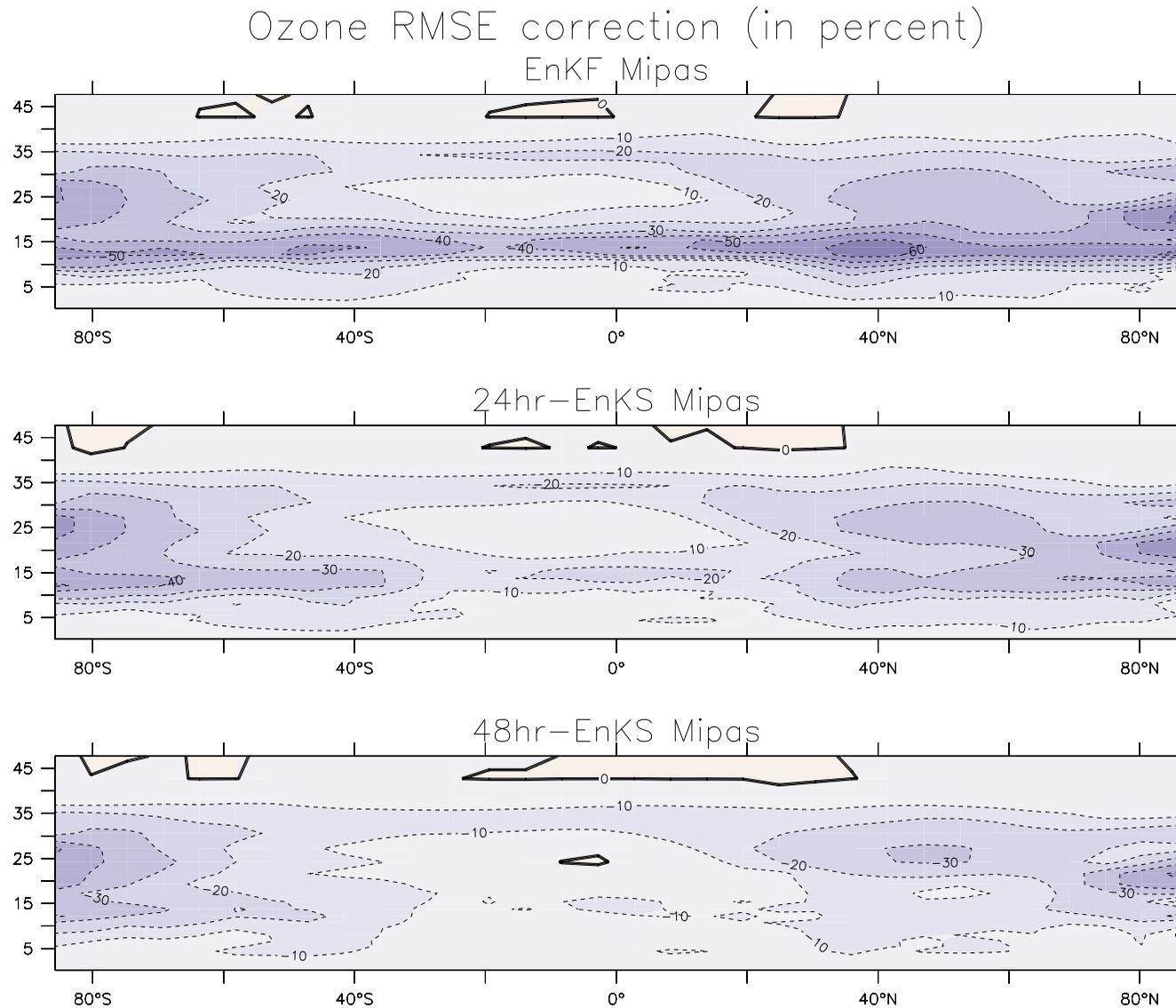
● EnKS Configurations

● EnKS Analyses

● Experiment

● EnKS vs. EnKF

CONCLUSIONS AND FUTURE
WORK



Ensemble Kalman Smoother

Experiments

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

- EnKS Configurations
- EnKS Analyses
- Experiment
- EnKS vs. EnKF

CONCLUSIONS AND FUTURE
WORK

EnKF Mipas	daily Mipas data assimilation
EnKS Mipas	daily EnKF Mipas data assimilation + two days of posterior Mipas data assimilation
EnKF 3×Mipas	daily assimilation of three times the amount of Mipas data, same amount of observations as EnKS Mipas

EnKF vs. EnKS: U RMSE Corrections

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

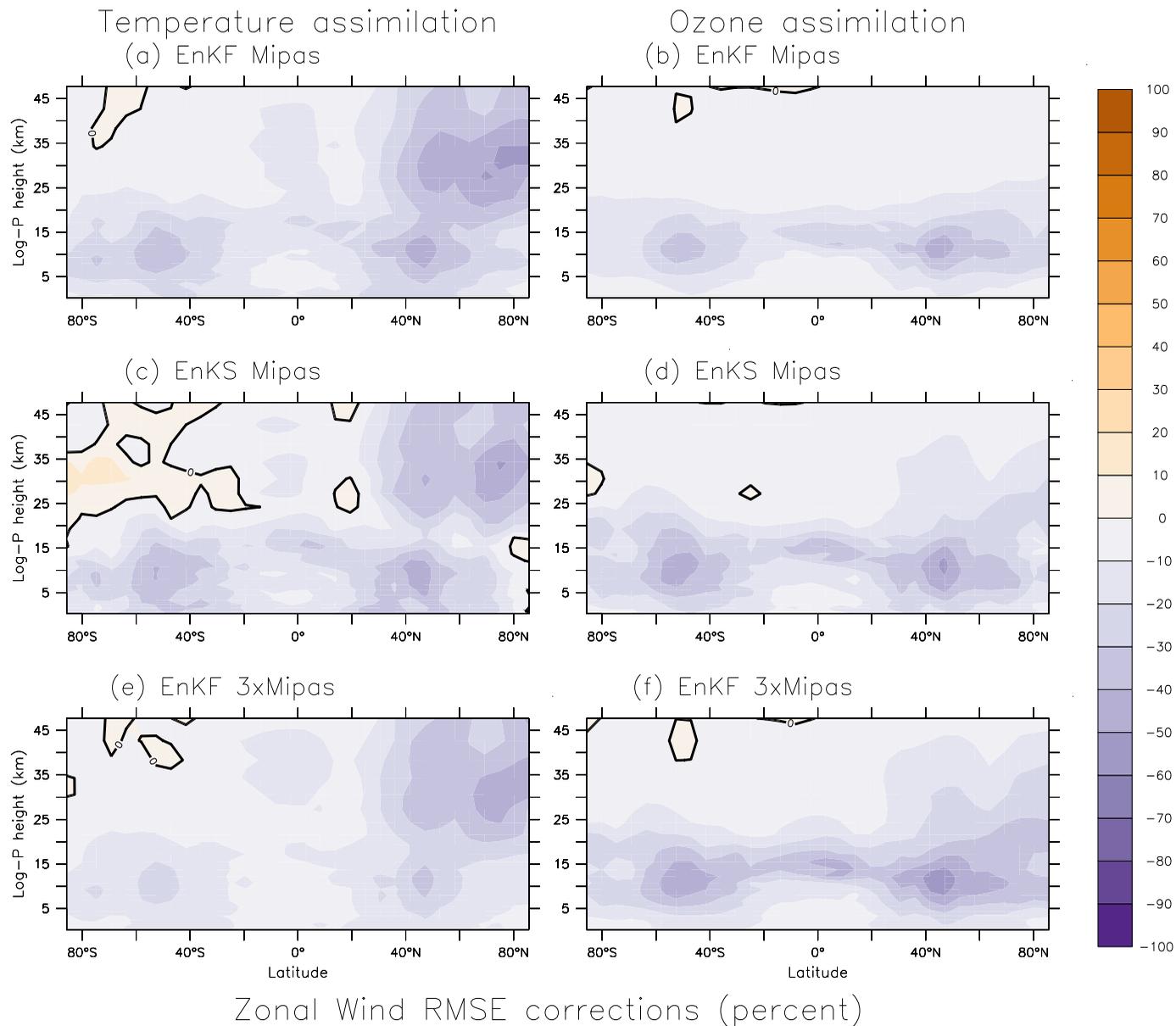
● EnKS Configurations

● EnKS Analyses

● Experiment

● EnKS vs. EnKF

CONCLUSIONS AND FUTURE
WORK



EnKF vs. EnKS: O₃ RMSE Corrections

INTRODUCTION

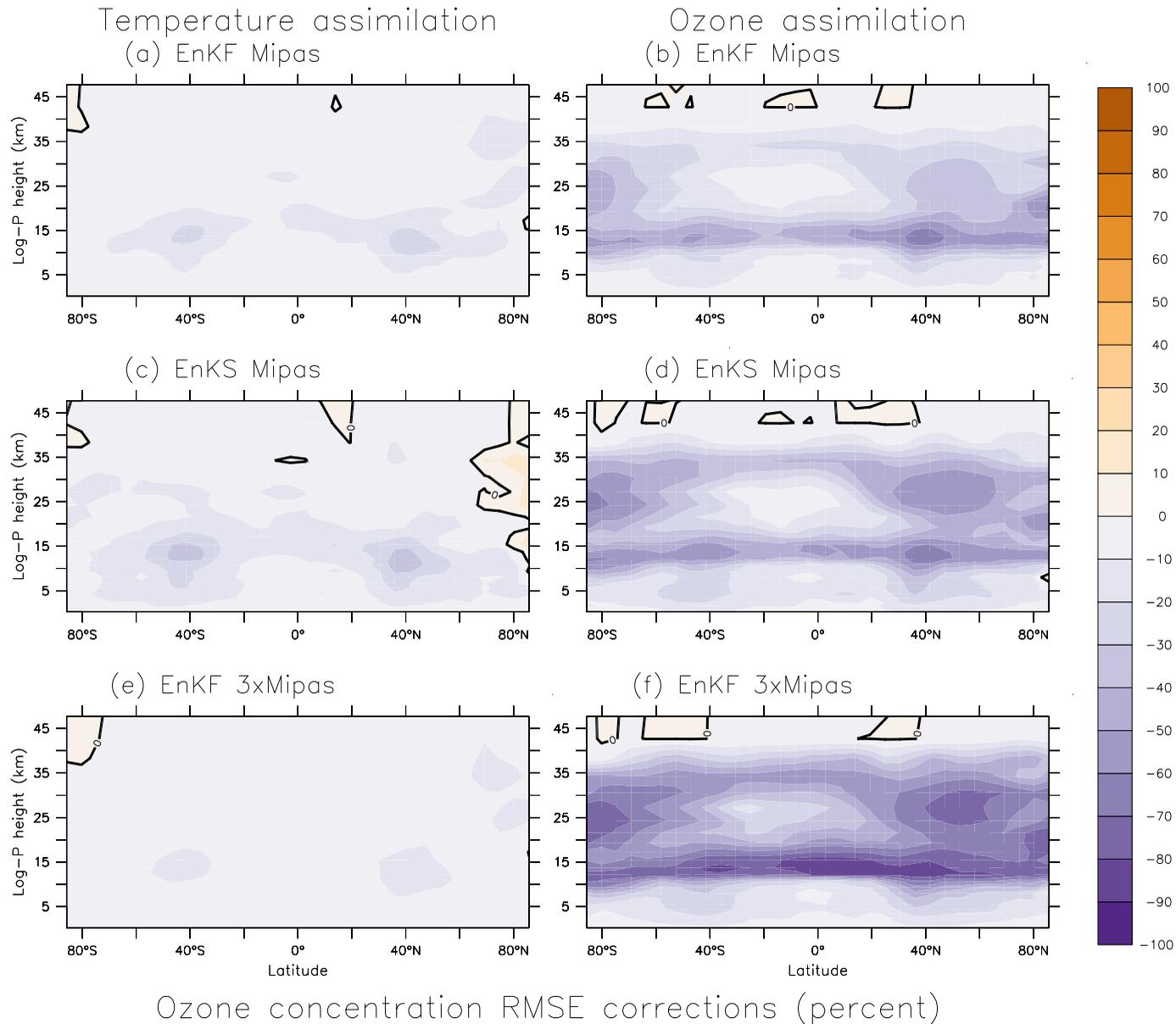
EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

- EnKS Configurations
- EnKS Analyses
- Experiment
- EnKS vs. EnKF

CONCLUSIONS AND FUTURE
WORK



INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

**CONCLUSIONS AND FUTURE
WORK**

- Summary
- Future Work

CONCLUSIONS AND FUTURE WORK

Summary

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

● Summary

● Future Work

EnKF-CCM :

- EnKF assimilation of MIPAS-like stratospheric observations in the IGCM-FASTOC can efficiently constrain the whole model state.
- Two-month ozone (10% error) or temperature (2K error) assimilation experiments yield approximately the same constraint on the dynamical state of the system.
- Temperature assimilation has however more problems constraining the chemical state.

CHEMISTRY-DYNAMICS INTERACTION:

- $\mathbf{T} \rightarrow \mathbf{O}_x$ covariances permit to slightly improve the ozone analysis, compared to using only $\mathbf{T} \rightarrow$ dynamics covariances.
- $\mathbf{O}_x \rightarrow \mathbf{u}$ and $\mathbf{O}_x \rightarrow \mathbf{T}$ covariances permit to constrain wind motion during ozone assimilation, but particularly $\mathbf{O}_x \rightarrow \mathbf{u}$ covariances.

Milewski, T. and M.S. Bourqui, 2011a: Assimilation of stratospheric temperature and ozone with an Ensemble Kalman Filter in a Chemistry-Climate Model. Monthly Weather Review 139, pp.3389-3404

ENSEMBLE KALMAN SMOOTHING :

- Steady decrease of analysis corrections with time-lag, but still beneficial corrections for 48 hours.
- For \mathbf{O}_x assimilation, analysis improvements from adding posterior data almost as good as from adding horizontal data.

Milewski, T. and M.S. Bourqui, 2011b: Impact of synchronous and asynchronous ensemble assimilation of stratospheric observations. In preparation.

Future Work

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

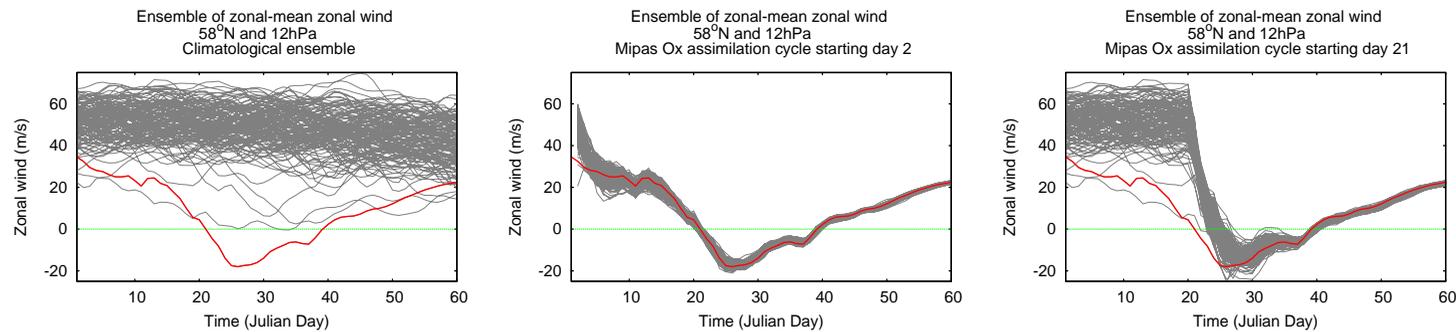
ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

● Summary

● Future Work

- Extreme events : stratospheric sudden warmings



- Imperfect-twin experiment : model errors
 - Additive or multiplicative inflation do not work in sparsely-observed systems.
 - Bias correction.

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

- Summary
- Future Work

THANK YOU !
ANY FEEDBACK IS HIGHLY APPRECIATED

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

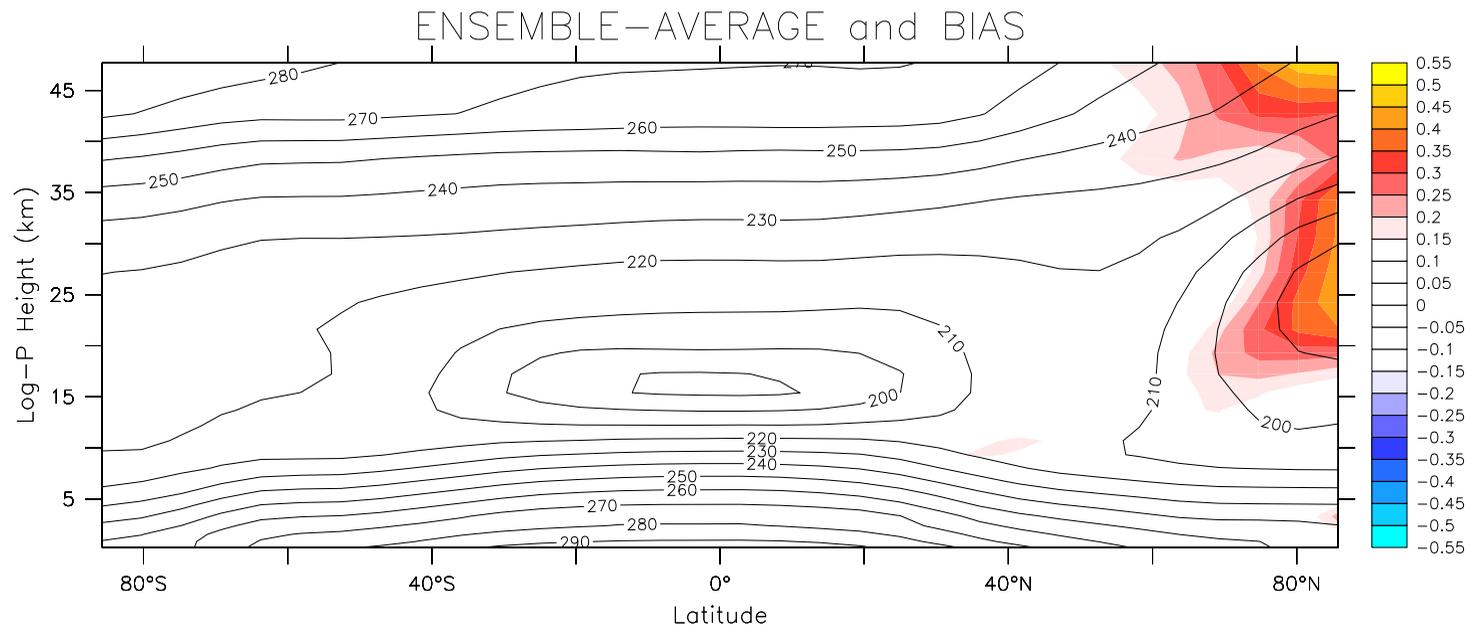
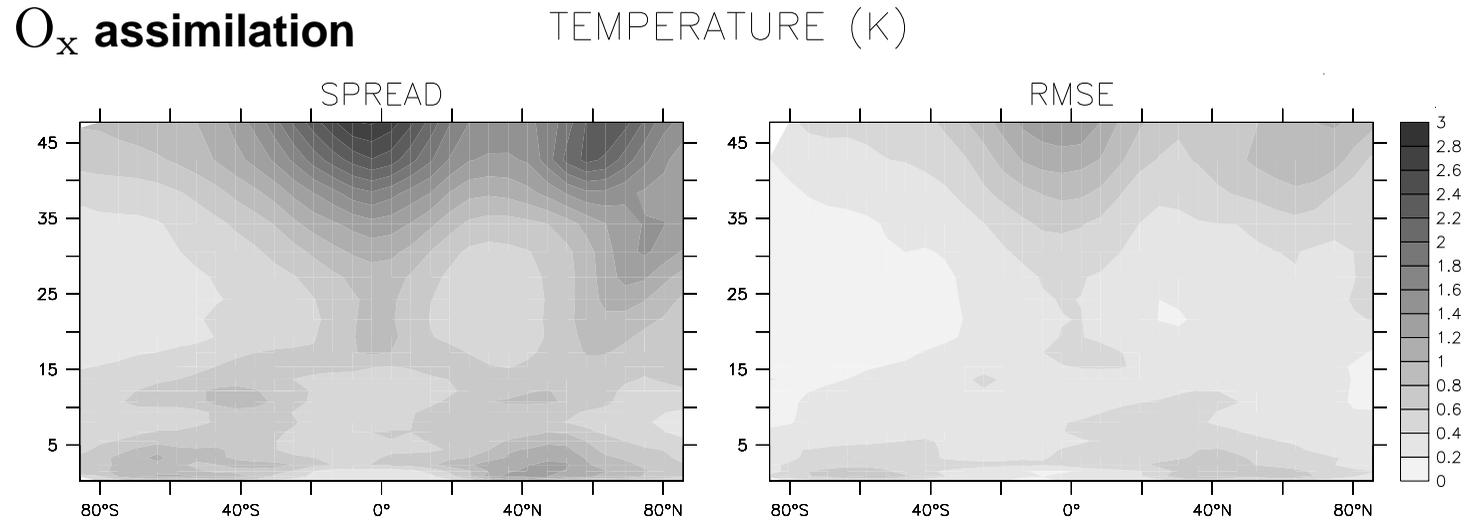
EXTRA FIGURES

EXTRA FIGURES

Zonally-Averaged Temperature Analyses

- INTRODUCTION
- EnKF-CCM
- CHEMISTRY-DYNAMICS INTERACTION
- ENSEMBLE KALMAN SMOOTHER
- CONCLUSIONS AND FUTURE WORK

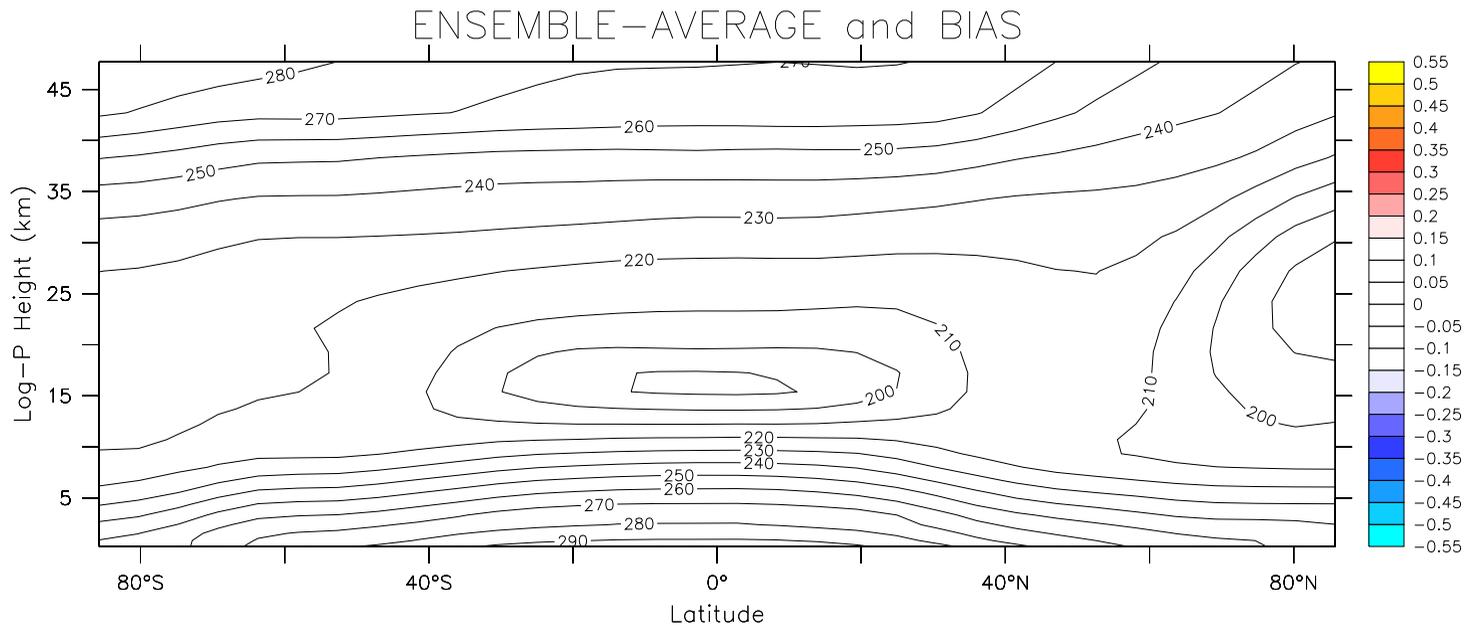
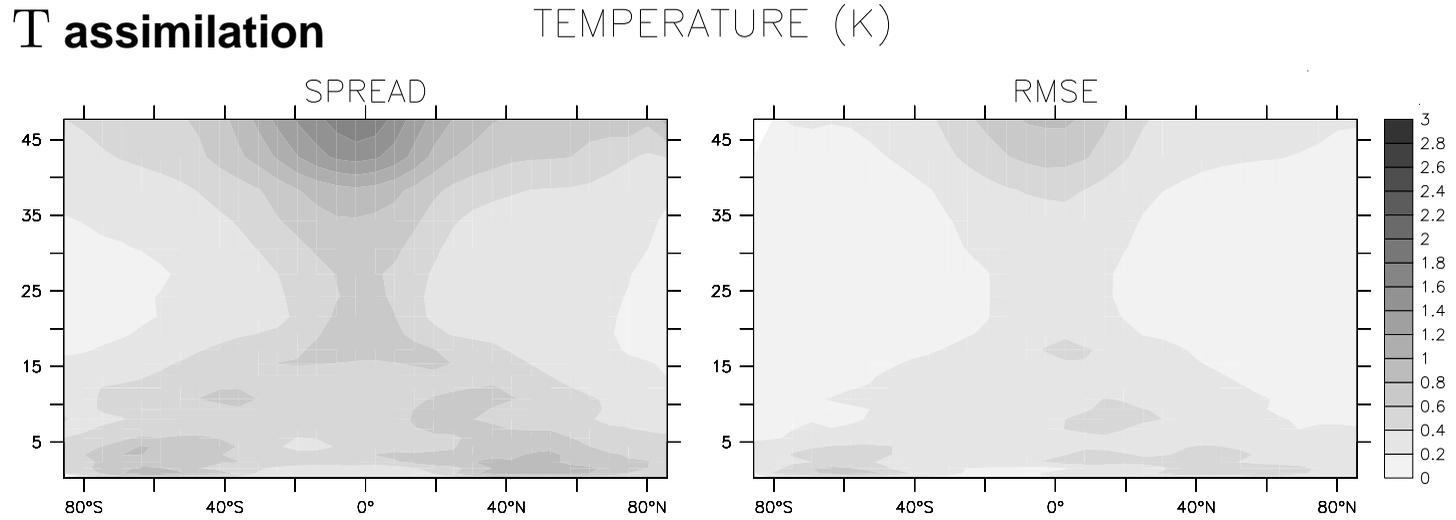
EXTRA FIGURES



Zonally-Averaged Temperature Analyses

- INTRODUCTION
- EnKF-CCM
- CHEMISTRY-DYNAMICS INTERACTION
- ENSEMBLE KALMAN SMOOTHER
- CONCLUSIONS AND FUTURE WORK

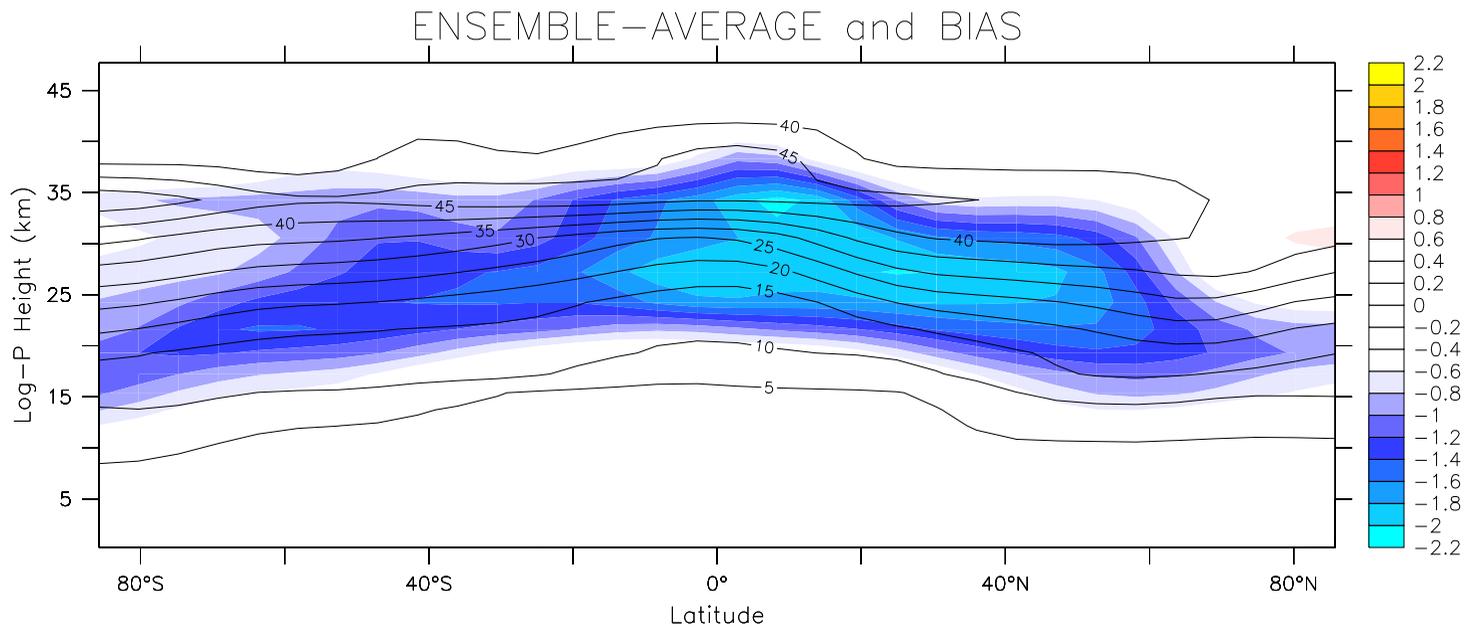
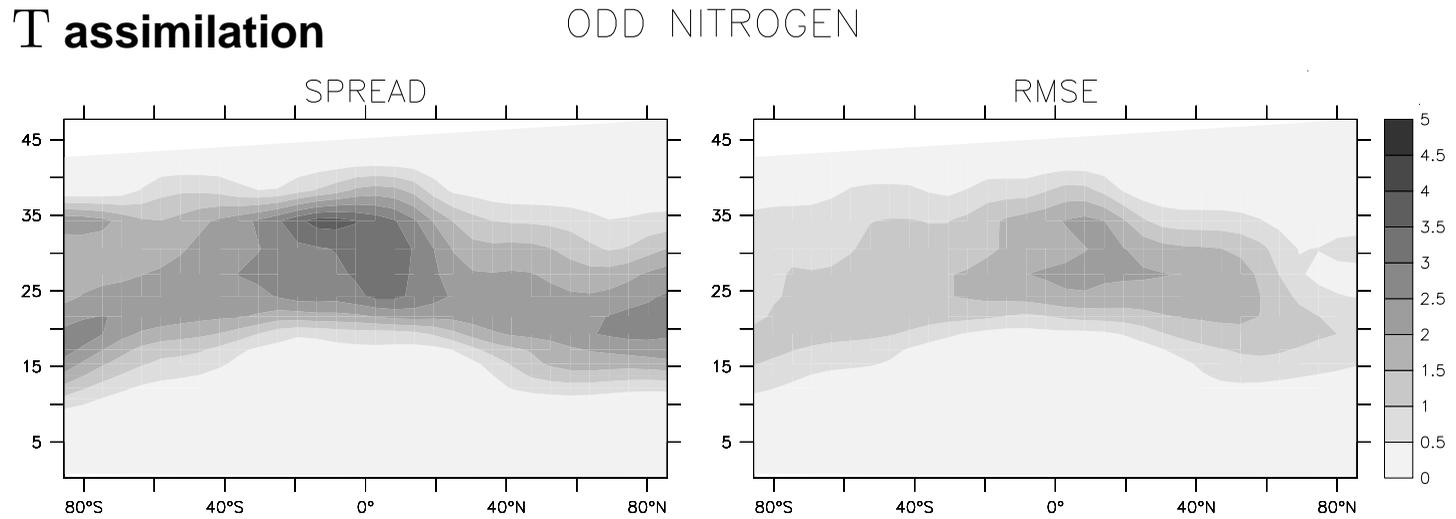
EXTRA FIGURES



Zonally-Averaged Odd Nitrogen Analyses

- INTRODUCTION
- EnKF-CCM
- CHEMISTRY-DYNAMICS INTERACTION
- ENSEMBLE KALMAN SMOOTHER
- CONCLUSIONS AND FUTURE WORK

EXTRA FIGURES



Zonally-Averaged Odd Nitrogen Analyses

INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

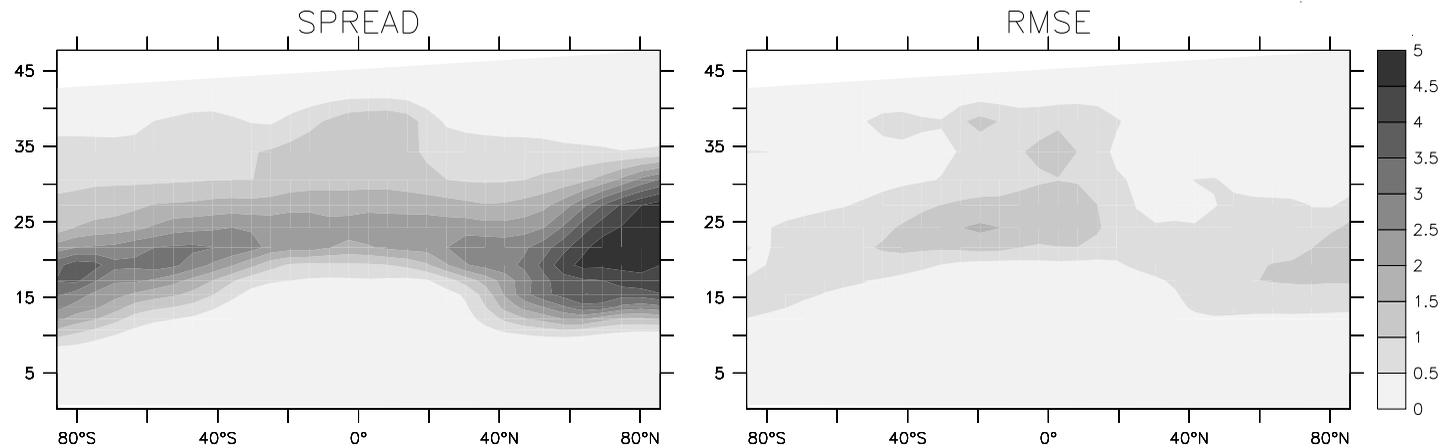
ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

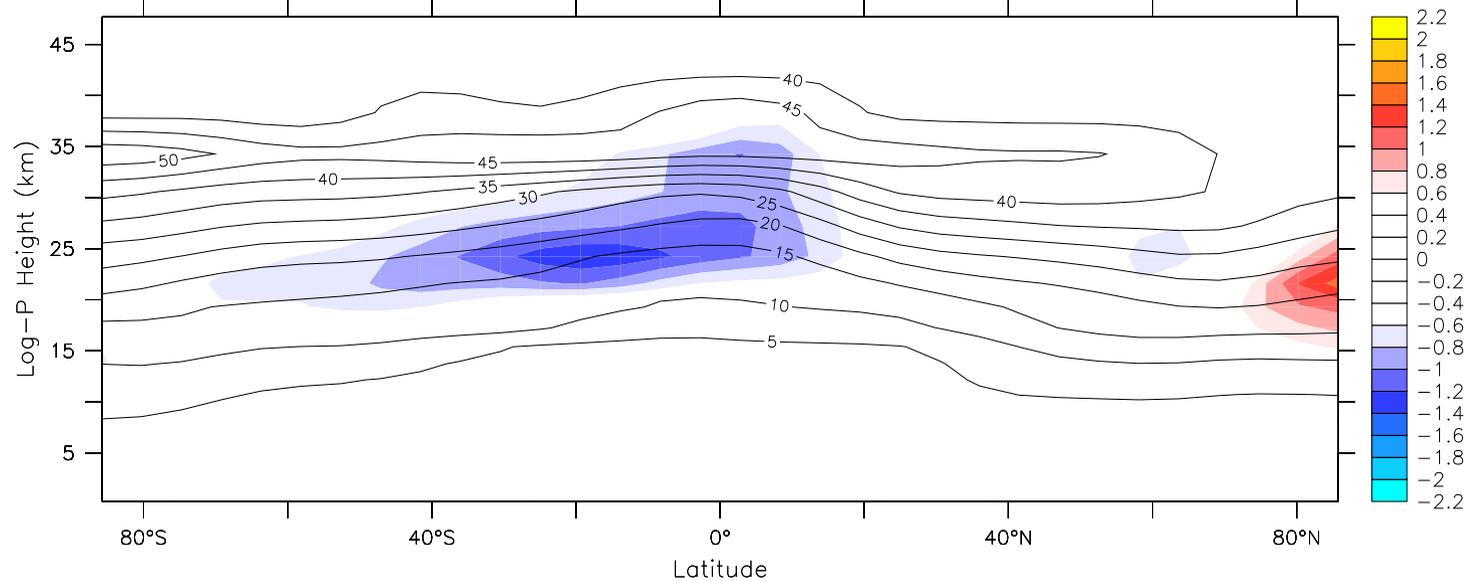
EXTRA FIGURES

O_x assimilation

ODD NITROGEN



ENSEMBLE-AVERAGE and BIAS



INTRODUCTION

EnKF-CCM

CHEMISTRY-DYNAMICS
INTERACTION

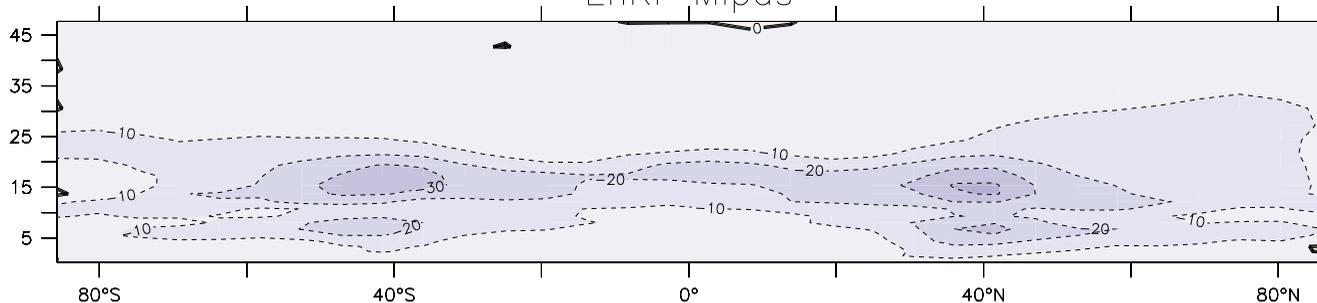
ENSEMBLE KALMAN
SMOOTHER

CONCLUSIONS AND FUTURE
WORK

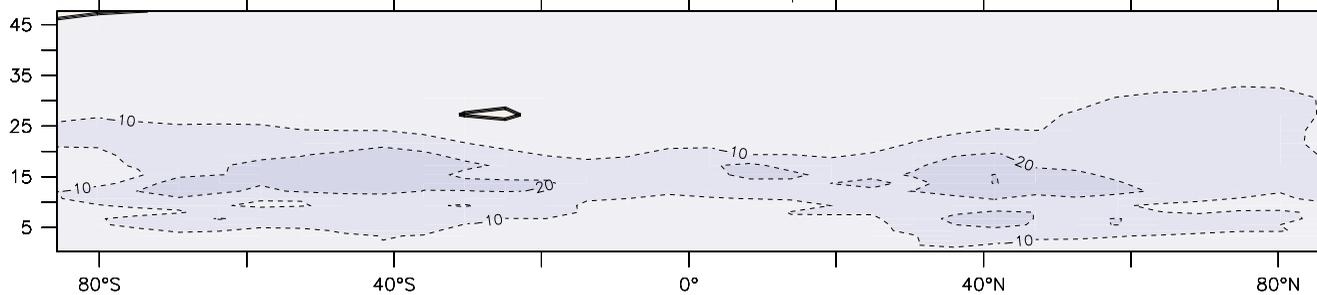
EXTRA FIGURES

Temperature RMSE correction (in percent)

EnKF Mipias



24hr-EnKS Mipias



48hr-EnKS Mipias

