



Tour d'horizon et assimilation de données à l'échelle convective

Jean-François Caron

Met Office @ Reading.

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2. a) Nowcasting & covariances d'erreurs de prévision à l'échelle kilométrique
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1. Prévision d'ensemble à aire limitée : comment produire des perturbations aux conditions initiales cohérentes avec le SPE porteur ?



Où est le Met Office ?





Met Office

Deterministic NWP Models

Global

- 25-km 70 L (80-km top)
- 4DVAR (60-km inner-loop)
- 144 hour forecast - 00/12Z
- 60 hour forecast - 06/18Z

NAE (North Atlantic and Europe)

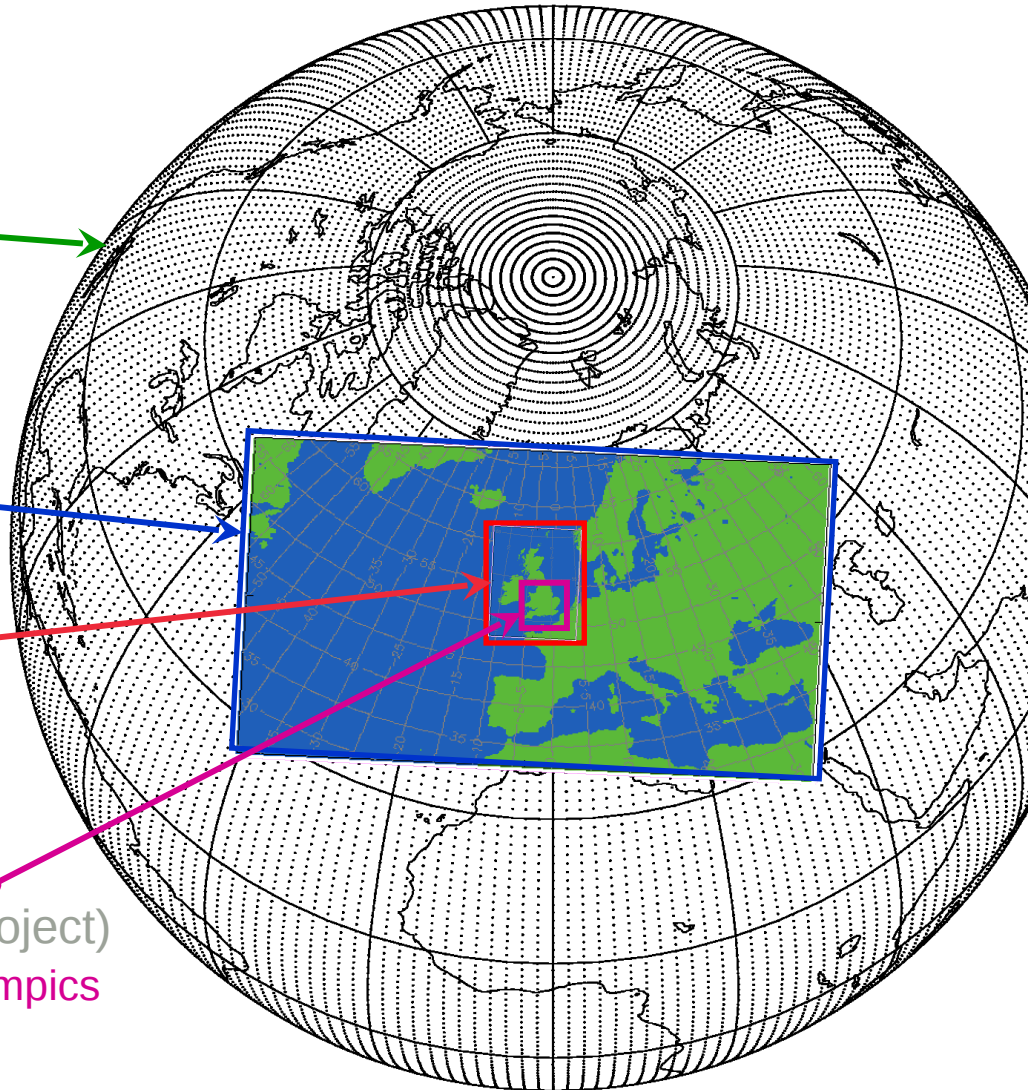
- 12-km 70 L (80-km top)
- 4DVAR (36-km inner-loop)
- 60 hour forecast - 4 times per day

UK4 & UKV

- 4/1.5-km 70 L (40-km top)
- 3DVAR
- 36 hour forecast - 4 times per day

NDP (Nowcasting Demonstration Project)

- In development for London's 2012 Olympics
- 1.5-km 70 L (40-km top)
- 4DVAR (3-km inner loop)
- 10 hour forecast - 24 times per day





Met Office

Met Office Global and Regional EPS

MOGREPS-G

- 24 member ETKF – 60 km
- 72 hour forecasts - 2 times per day (00/12Z)
- 15 day forecasts run @ ECMWF

MOGREPS-R (EU)

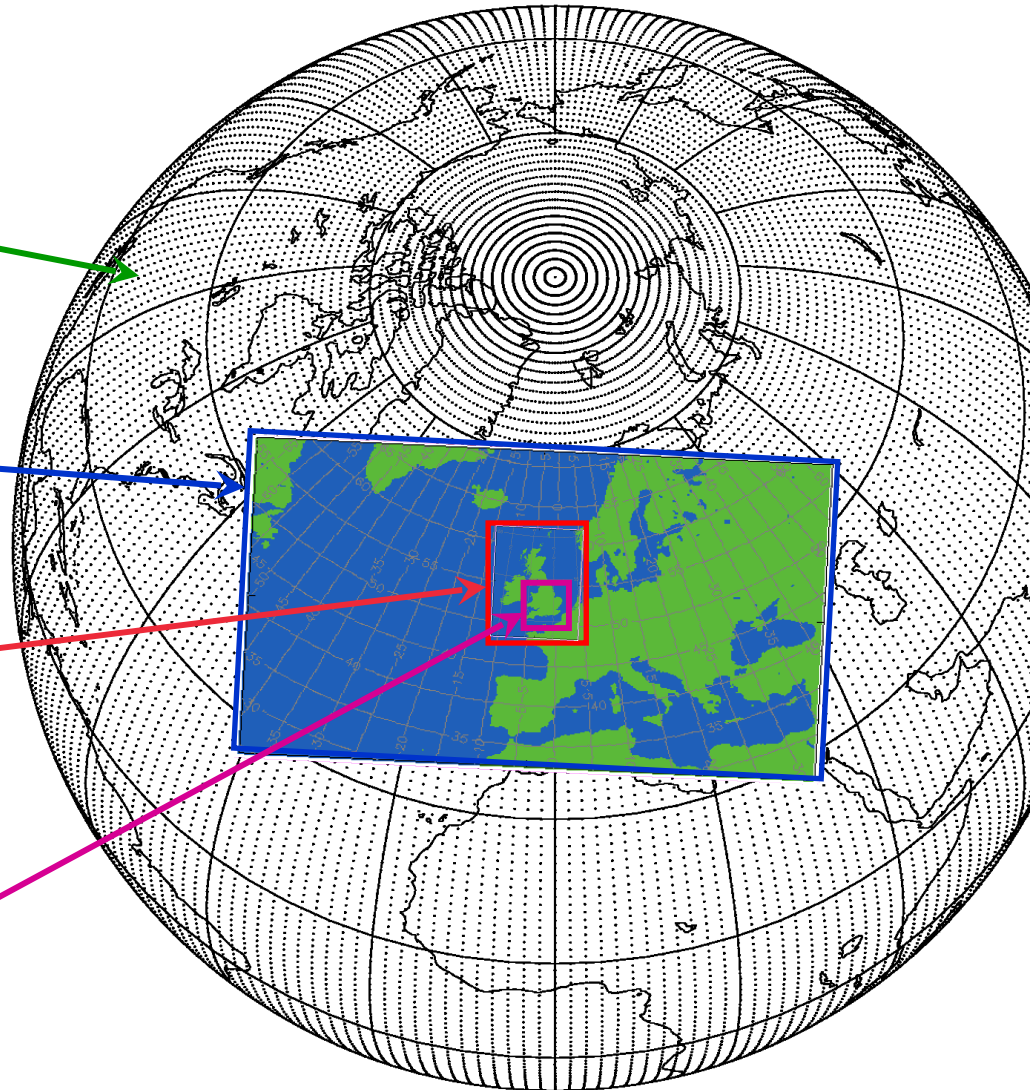
- Downscaled forecast pert. from global
- 24 member - 18-km
- 60 hour forecast - 2 times per day

MOGREPS-UK

- In development – For 2012
- Downscaling of MOGREPS-R
- 12 member - 2.2-km
- 36 hour forecast - 4 times per day

1.5-km ETKF

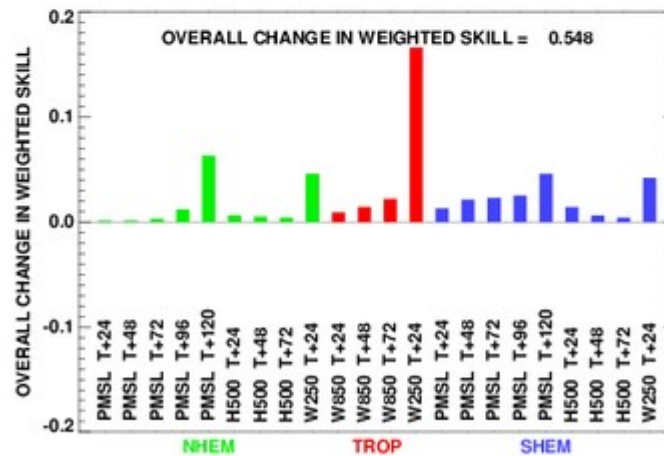
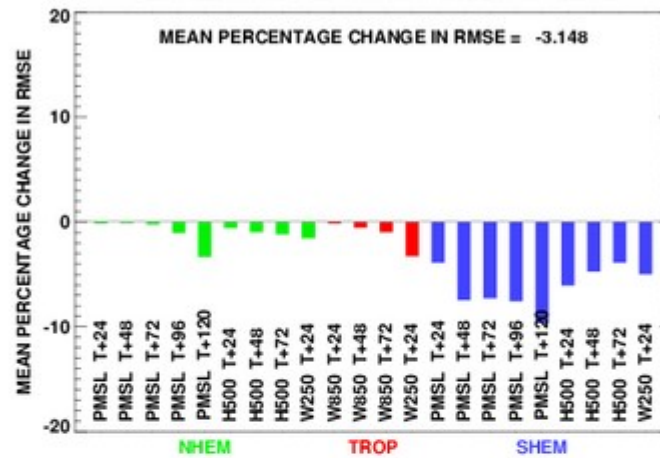
- Research only
- 24 member ETKF - 1.5-km
- 1 hour forecast - 24 times per day



4DVAR hybride Global

PARALLEL SUITE 27A: HYBRID ASSIMILATION (ON) VERIFICATION VS OBSERVATIONS

OVERALL CHANGE IN NWP INDEX = 2.373

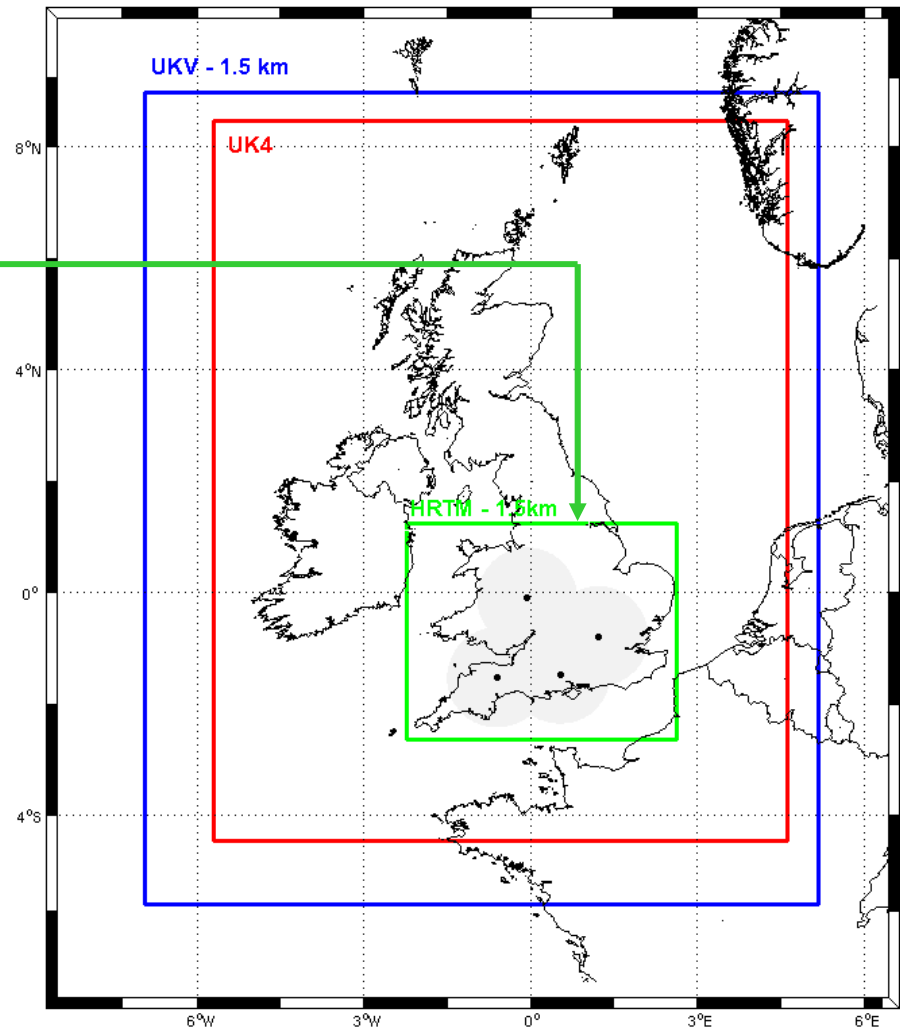




Background error covariances for the nowcasting demonstration project

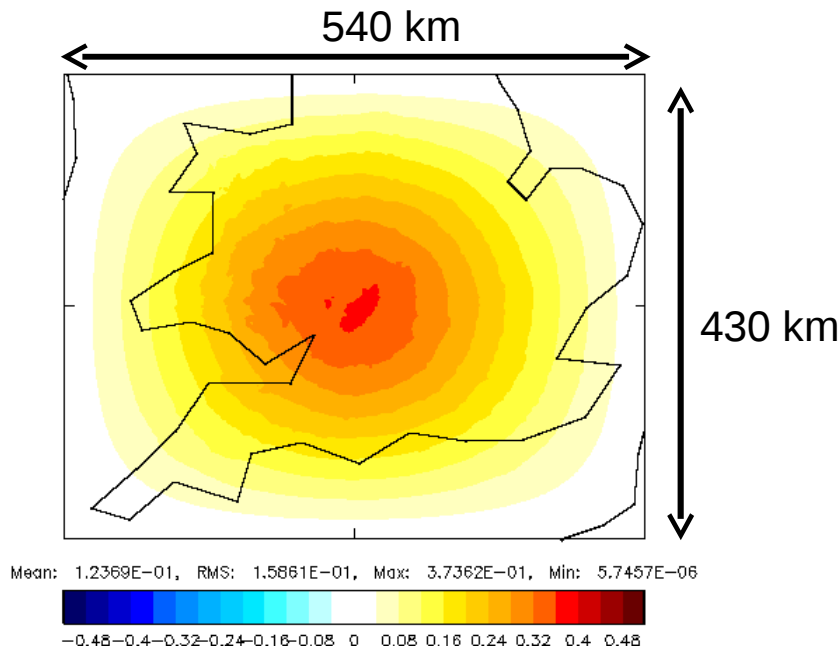
Nowcasting Demonstration Project

- 1.5 km NWP-based nowcasting system
- Southern UK only
- 4DVAR with Doppler winds
- Hourly cycling
- ~12 hour forecasts
- To be run in experimental mode during London's 2012 summer Olympics
- So far, all trials were made using UKV background error covariances



The elephant in the room

Response to the assimilation of a pseudo temperature at level 20 (~850 hPa)



theta at level 20 (~850 hPa)

Training data

- A set of 74 NMC 24h-12h (using same LBC) from UK4 model

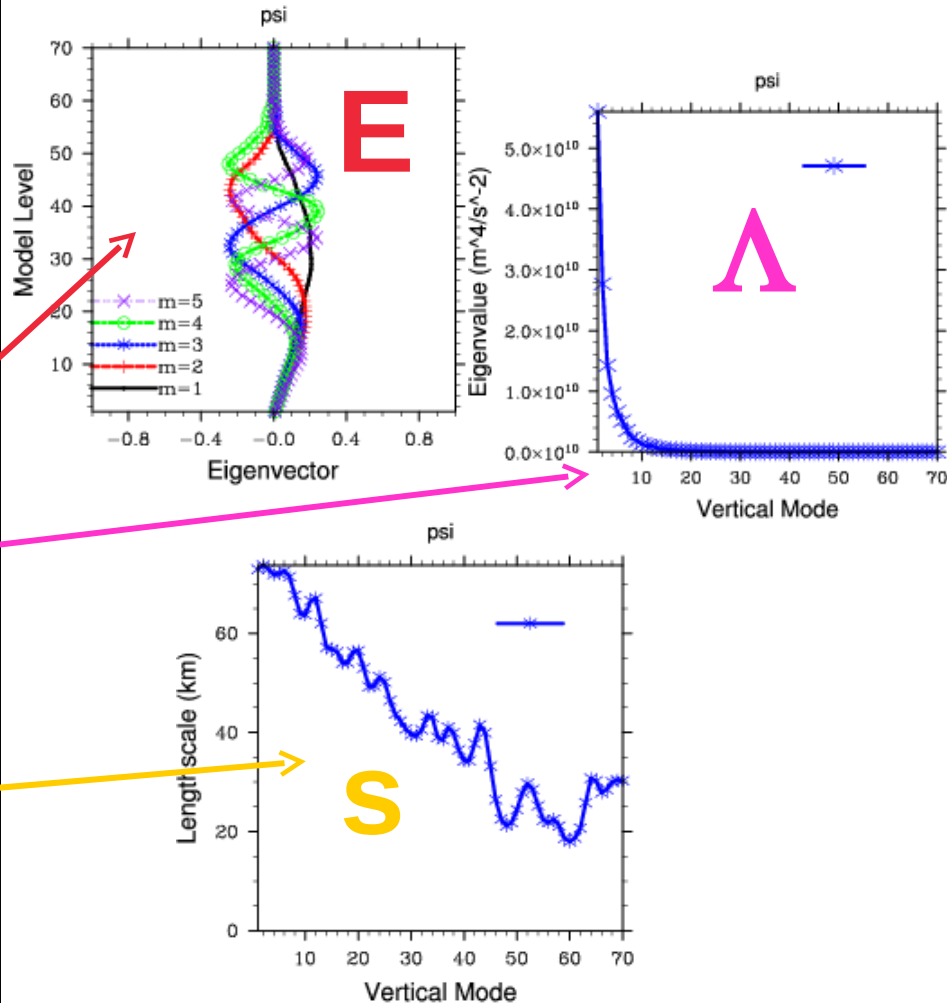
LAM covariance model

- Horizontal correlation modelled using **ad hoc lengthscales**.
- Control variables : **Same as in global and NAE**
- Temperature increments obtained from P increments assuming **hydrostatic balance**
- **No Ekman pumping, no coupling with humidity, no coupling between latent heating and w ...**

MetO VAR-LAM covariance model

Steps from model variable space to covariance model space

1	\mathbf{x}'	$u', v', \theta', \rho', p', q'$
2	Parameter transform $\mathbf{v}_p = \mathbf{T}_p \mathbf{x}'$	ψ', χ', p'_u, r' $p'_u = p'_h \mathbf{N} f(\psi')$ $\mathbf{N} = \langle f(\psi'), p'_h \rangle \langle p'_h, p'_h \rangle^{-1}$
3	Vertical transform \mathbf{T}_v	<p style="color: cyan;">auto covariances only</p> $\mathbf{B}_v(i, j) = \mathbf{E} \mathbf{\Lambda} \mathbf{E}^T$ $\mathbf{v}_p(i, j, k) \rightarrow \mathbf{v}_v(i, j, m)$
4	Horizontal transform \mathbf{T}_h	$\mathbf{B}_h(r) = \mathbf{B}_h(0) \left(1 + \frac{r}{s} \right) e^{r/s}$



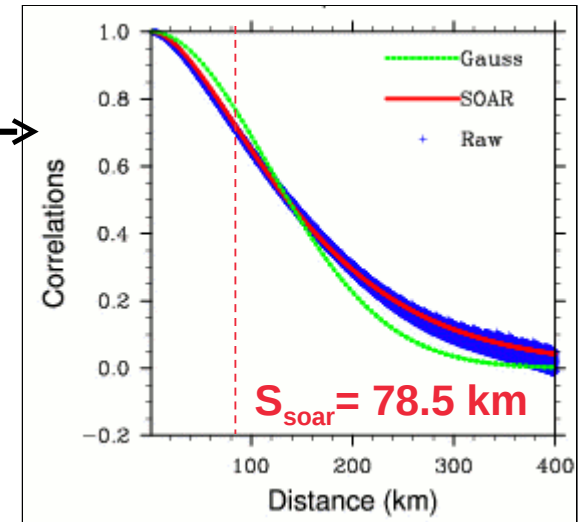


Gen_BE at the Met Office

(J.-F. Caron and Dale Barker)

Gen_BE is WRF's utility to compute background error covariances

- Clean and flexible conception: easy to test ideas
- Horizontal length scale in vertical mode space can be evaluated objectively. →
- WRF covariance model similar to the UK MetO VAR (e.g. vertical transform first)
- Already a multi-model tool: used by NCAR (WRF) and KMA.



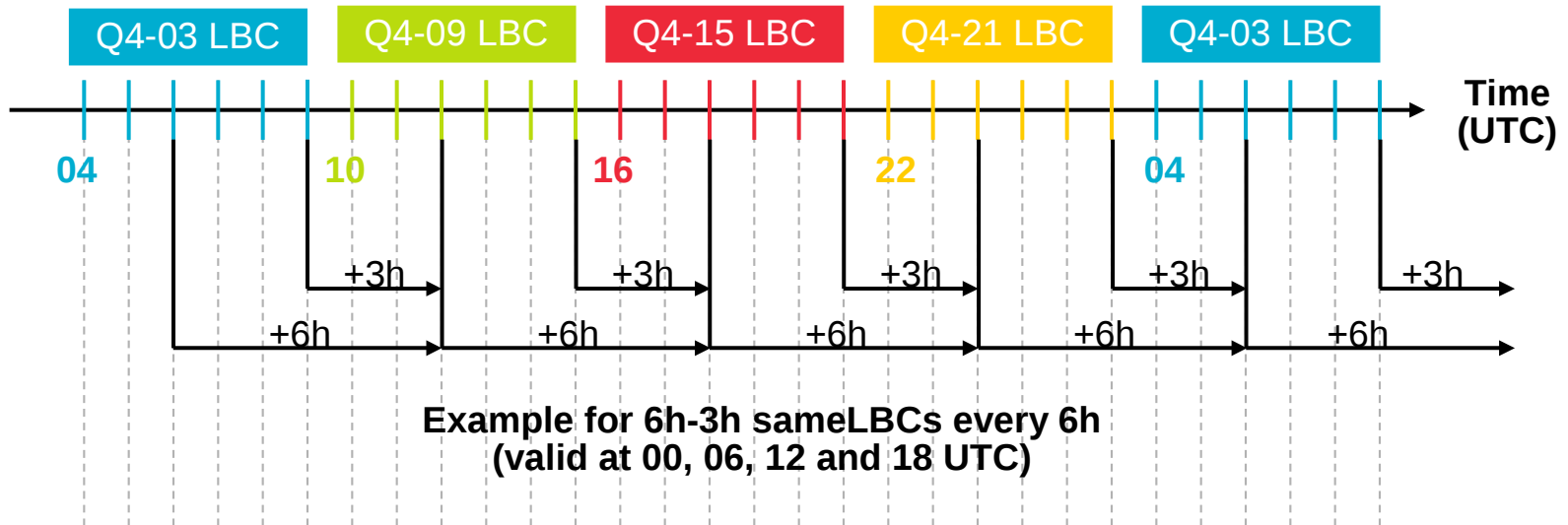
Objective #1: Adapt Gen_BE to provide the appropriate input to the actual VAR's covariance model in limited area mode (completed)

Objective #2: Use Gen_BE to conduct examinations of B at the convective scale (in progress)

Training data for the NDP

NMC method

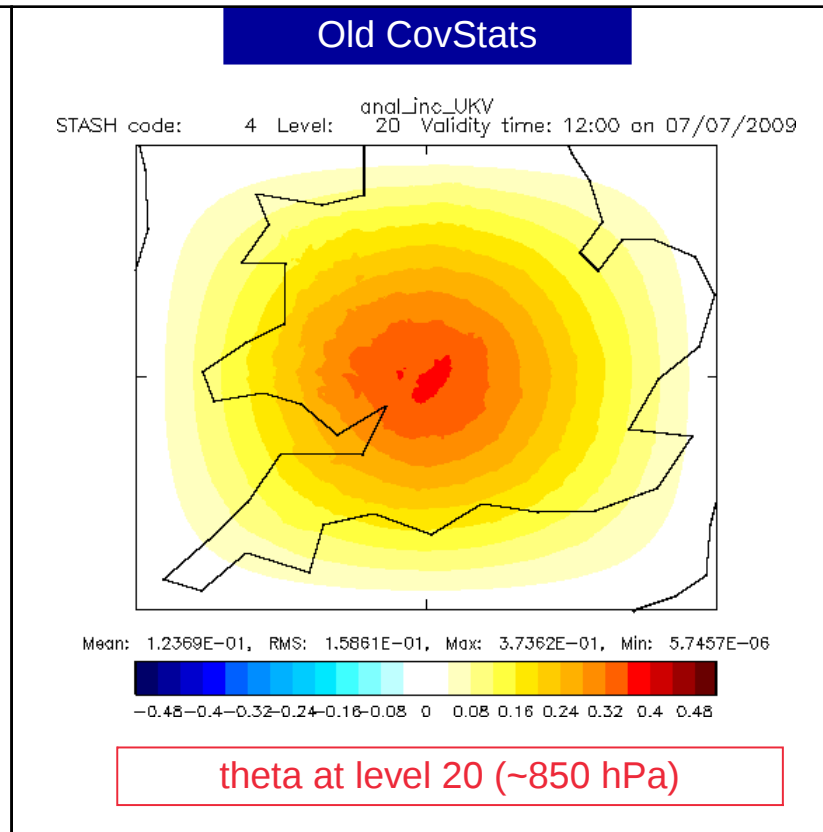
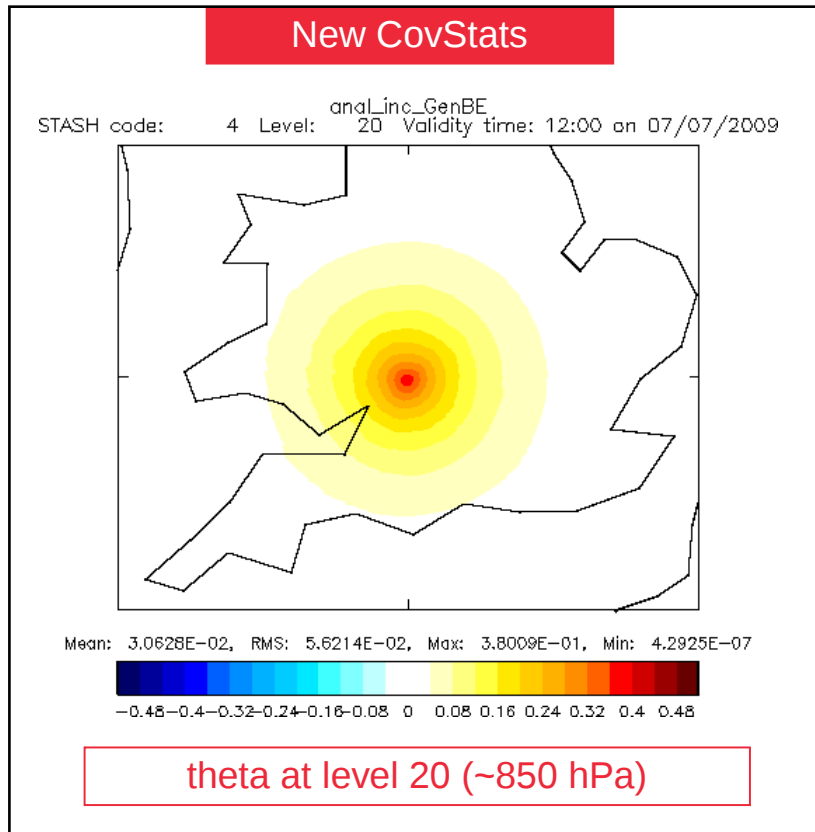
- From a three week cycle of the hourly SUK 1.5km (16/03/10 to 06/04/10) in 3DVAR mode using the UKV Covstats
- Outputs from +1h to +6h, every 1h
- Possibility to compute lagged forecast differences with forecasts having the same LBCs or different LBCs



Single pseudo-obs experiments

http://www-nwp/~frfc/work/hrtm/1obs/gen_be1.html

Response to the assimilation of a pseudo temperature at level 20 (~850 hPa)
 innovation = 2 K and observation error = 1 K



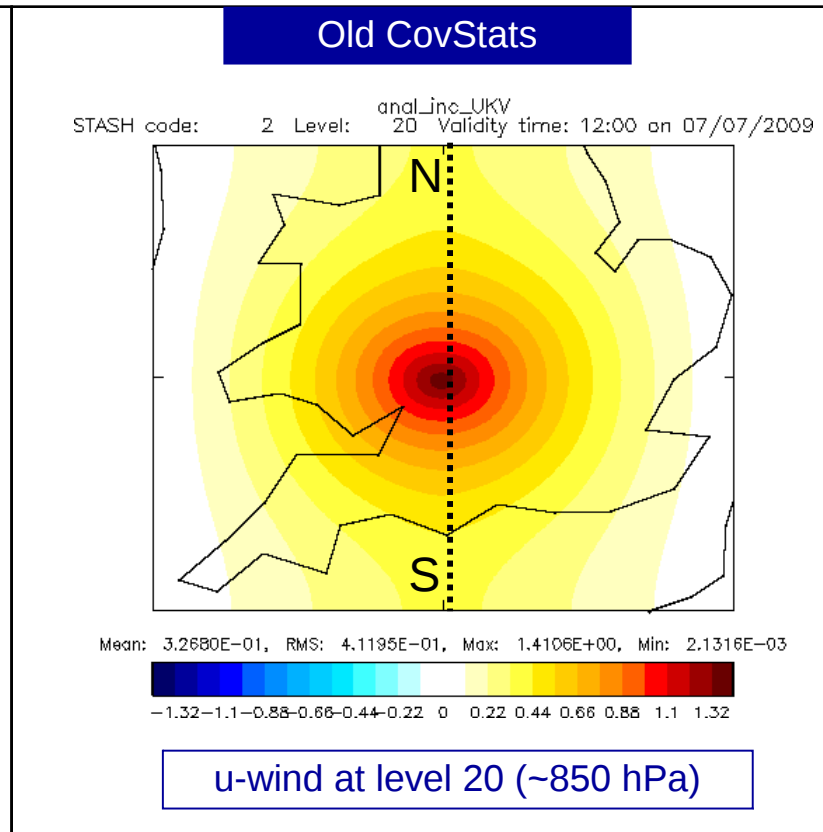
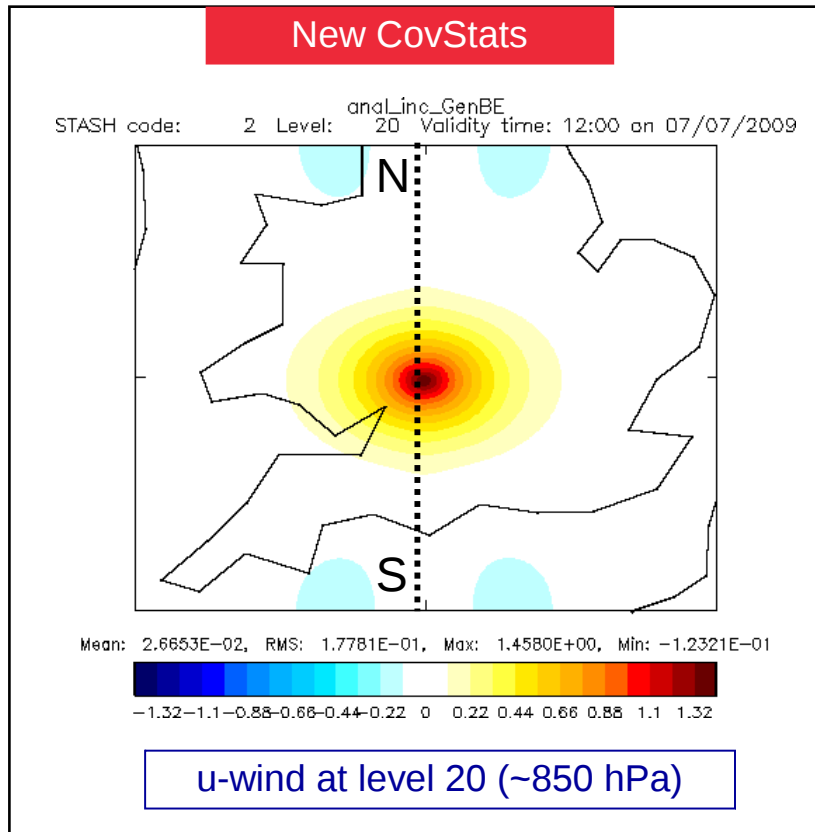
Single pseudo-obs experiments

http://www-nwp/~frfc/work/hrtm/1obs/gen_be1.html

Response to the assimilation of a pseudo

u-wind at level 20 (~850 hPa)

innovation = 2 m/s and observation error = 1m/s

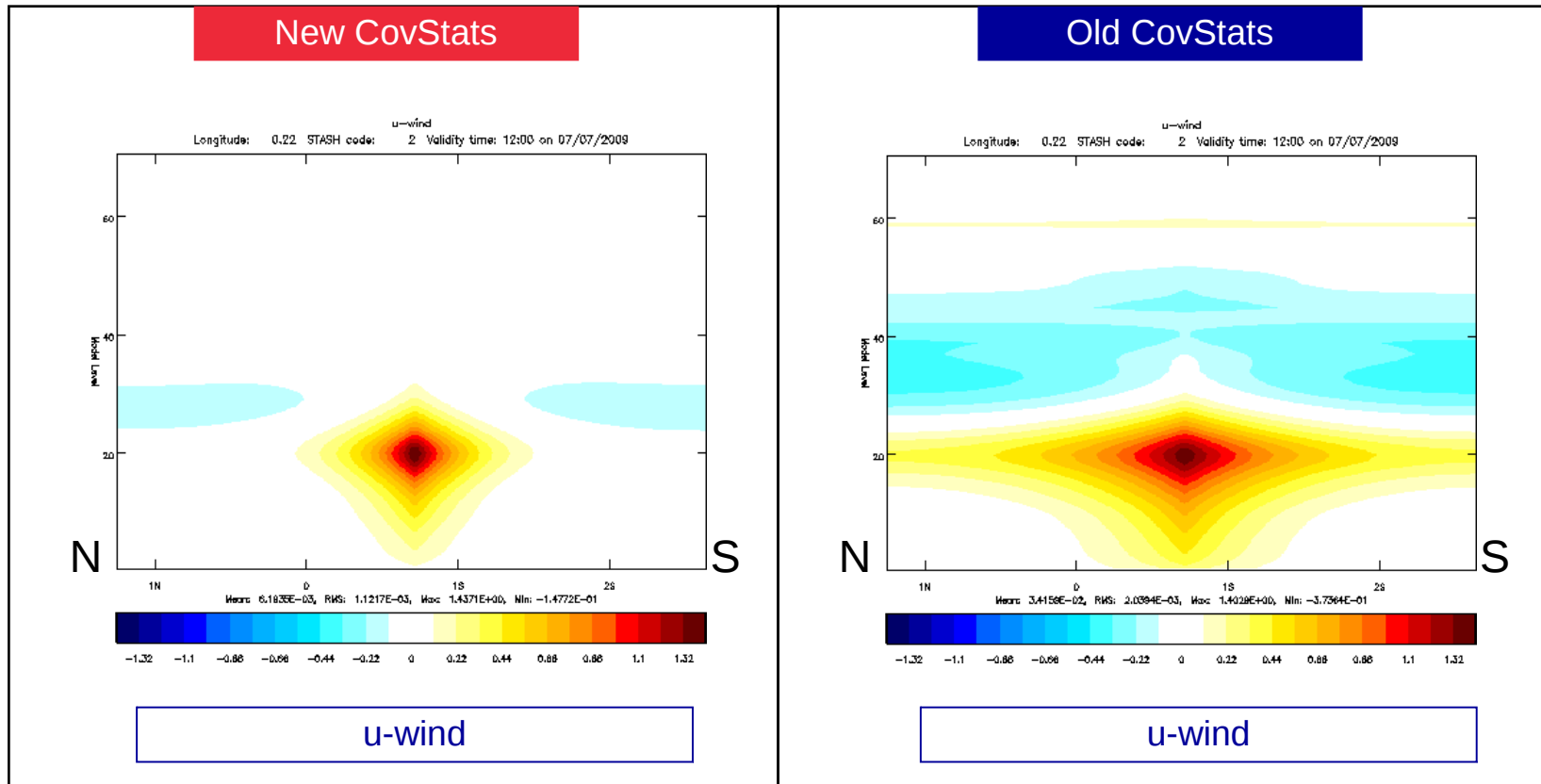


Single pseudo-obs experiments

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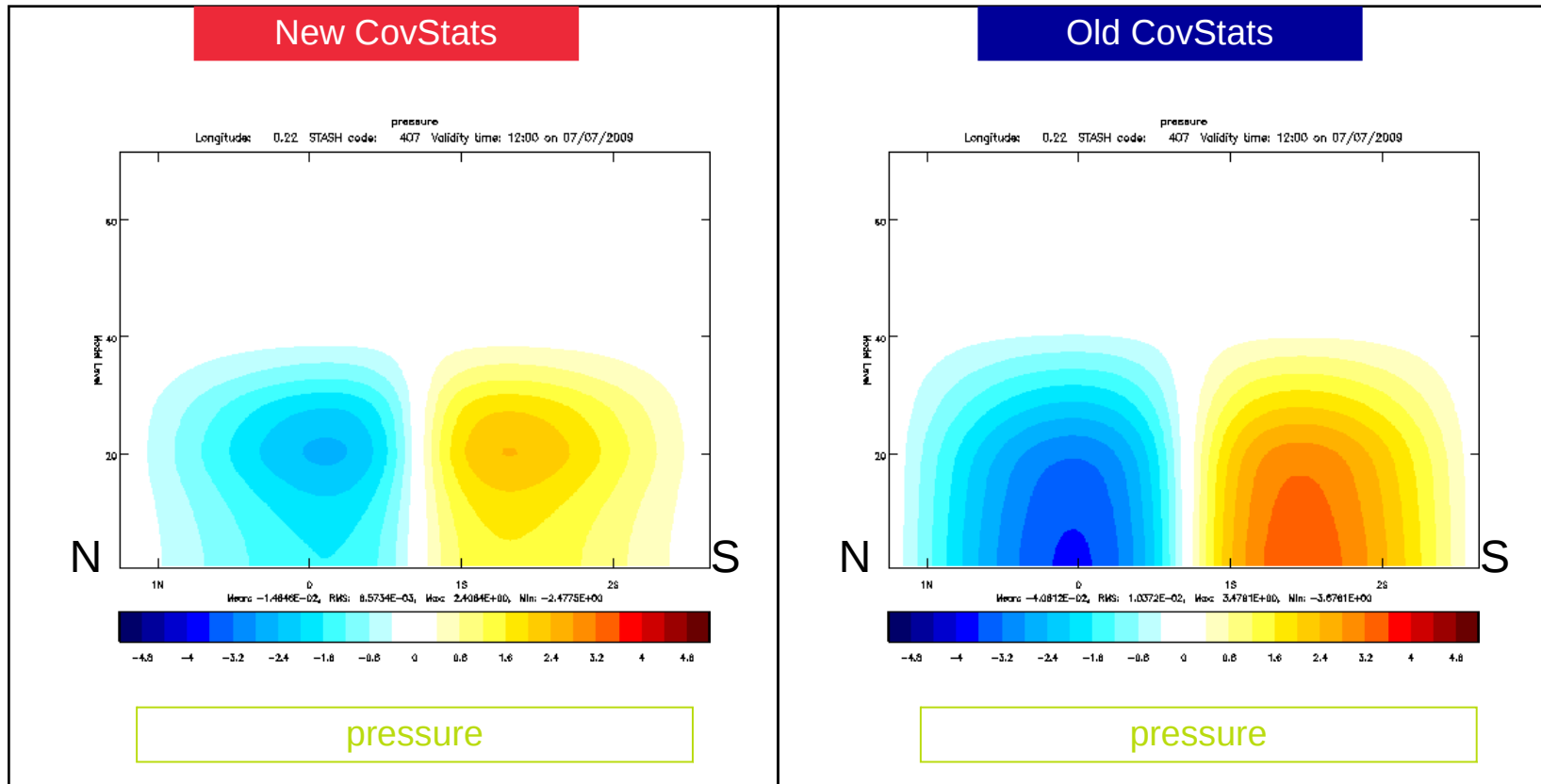


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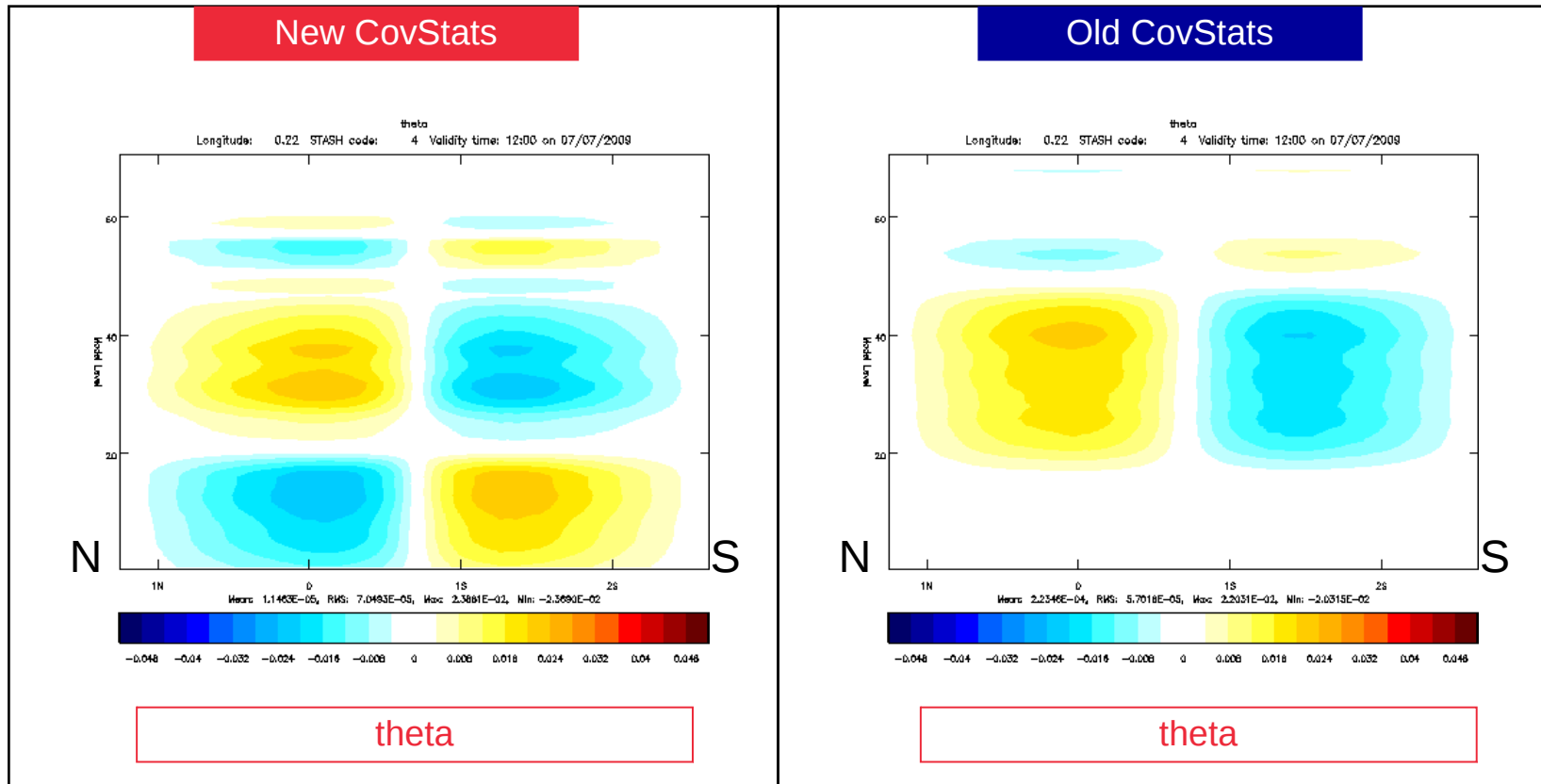


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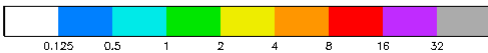
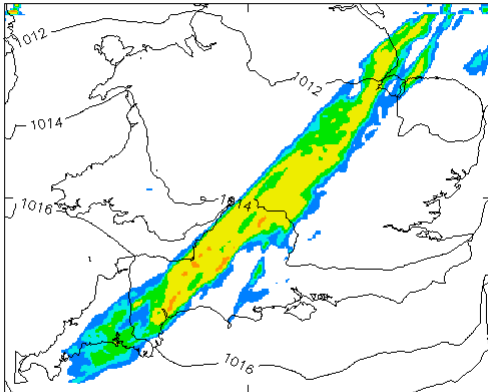


Forecast impact (from Zihong Li)

So far so good...

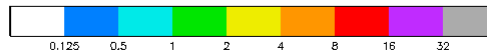
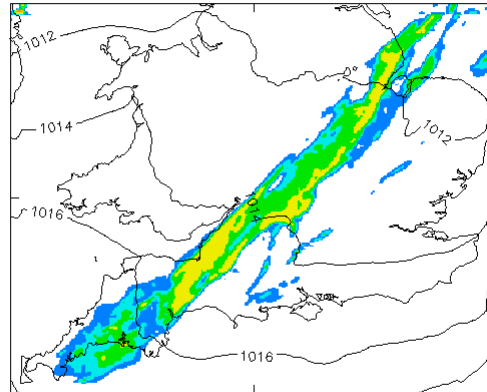
30-min forecasts from 4DVAR analyses

(conv off) large scale rain and snow rate only (4203 and 4204)
At 15Z on 18/ 5/2011, from 15Z on 18/ 5/2011



Old CovStats

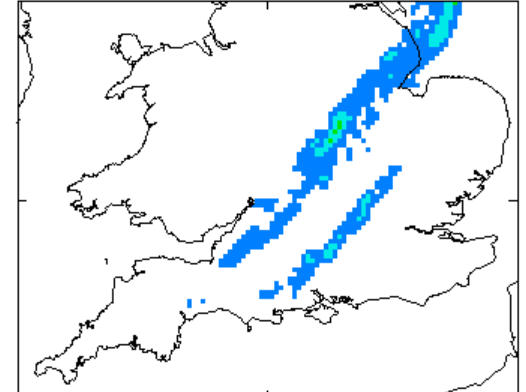
(conv off) large scale rain and snow rate only (4203 and 4204)
At 15Z on 18/ 5/2011, from 15Z on 18/ 5/2011



New CovStats

Radar

15:00 RADAR RAINFALL RATE
At 15Z on 18/ 5/2011, from 15Z on 18/ 5/2011





Summary and future work

- Before: An elephant was in the room

We were using **inappropriate estimates of forecast errors** in a **covariance model developed for large scale DA** using **ad hoc horizontal length scales**

- Now: An hippopotamus is in the room

We are using **more appropriate estimates of forecast errors** in a **covariance model developed for large scale DA** using **objective estimates of horizontal length scales**

- Next

Develop an appropriate covariance model for convective scale DA



Using Gen_BE to conduct examinations of B at the convective scale

B_{precip} vs. $B_{\text{non precip}}$



Intro

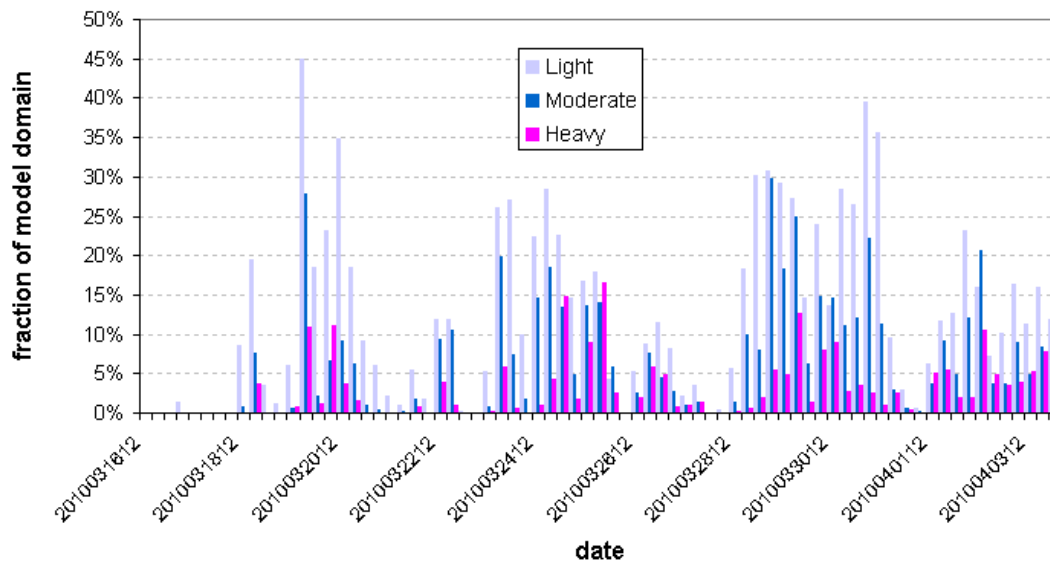
- Short-term precipitation forecast is the key objective in meso- γ scale forecasting & radar observations are the main component of the meso- γ scale observing system.
- Is **B** very different in precipitation areas ?
- How to take these differences into account in VAR without '**B** of the day' ?

Caron and Fillion, 2010	<ul style="list-style-type: none">• (Linear) Balance decreases with precipitation intensity at synoptic and meso-β scales• Introduced the idea of deriving rain-dependent background error statistics.• Showed that statistical balance operators can be improved over precipitation areas by using only profiles from precipitation areas in training data.
Montmerle and Berre, 2010	<ul style="list-style-type: none">• Compared $\mathbf{B}_{\text{precip}}$ vs $\mathbf{B}_{\text{non precip}}$ in summer convective regime at meso-γ scale and showed significant differences• Implemented rain-dependent B formulation in AROME 3DVAR
Michel et al, 2011 (in press)	<ul style="list-style-type: none">• $\mathbf{B}_{\text{precip} + \text{hydrometeors}}$
Ménétrier and Montmerle, 2011 (submitted)	<ul style="list-style-type: none">• \mathbf{B}_{fog}

Definition of precipitations areas

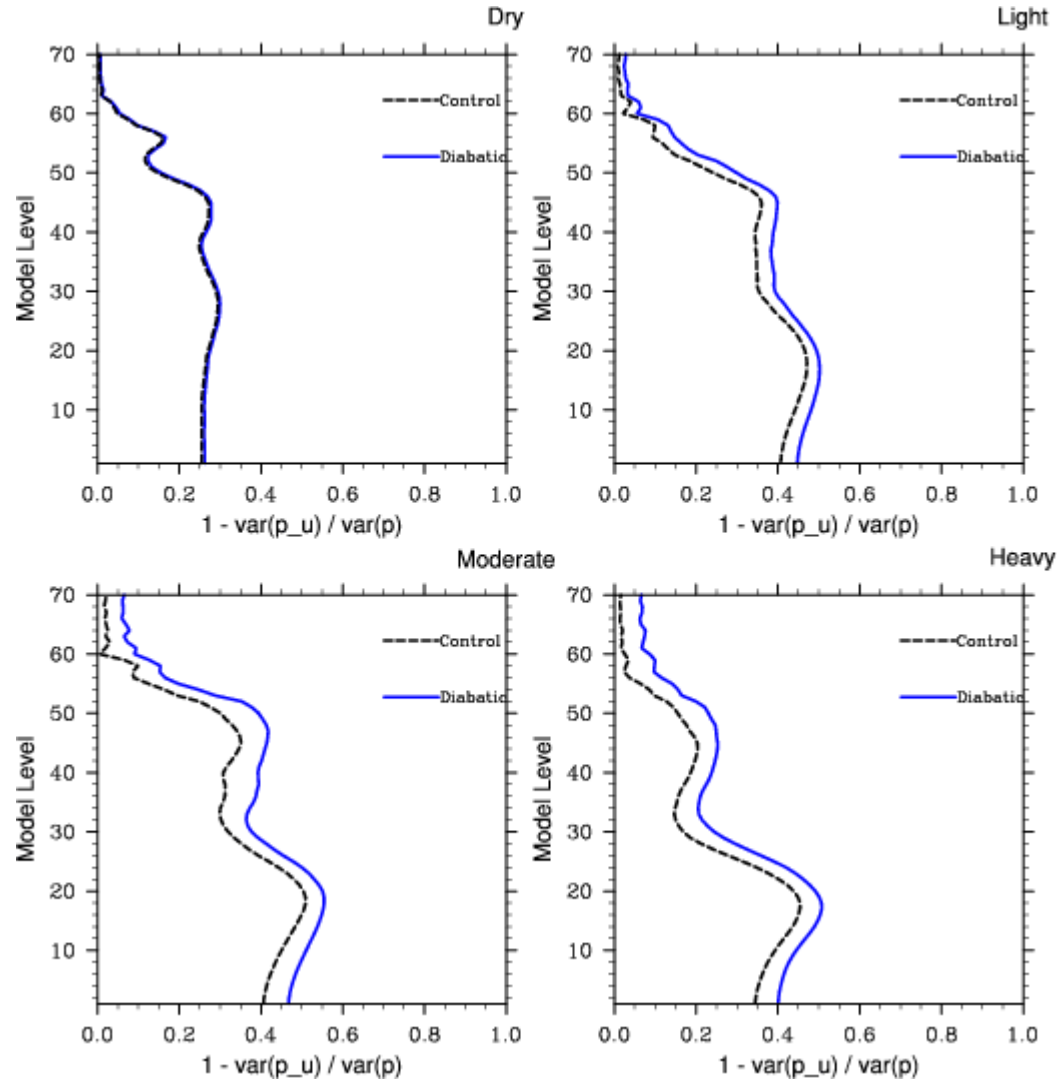
Category	$P_M = \text{Max}(P_{6h}, P_{3h})$ (mm/h)	Fraction of total points (%)	Fraction of precip points (%)
Dry	$P_M < 0.1$	77.9	—
Light	$0.1 \leq P_M < 1.0$	12.6	56.9
Moderate	$1.0 \leq P_M < 2.75$	6.6	29.8
Heavy	$2.75 \leq P_M$	2.9	13.3

Time-series of precip categories in (NMC) training data



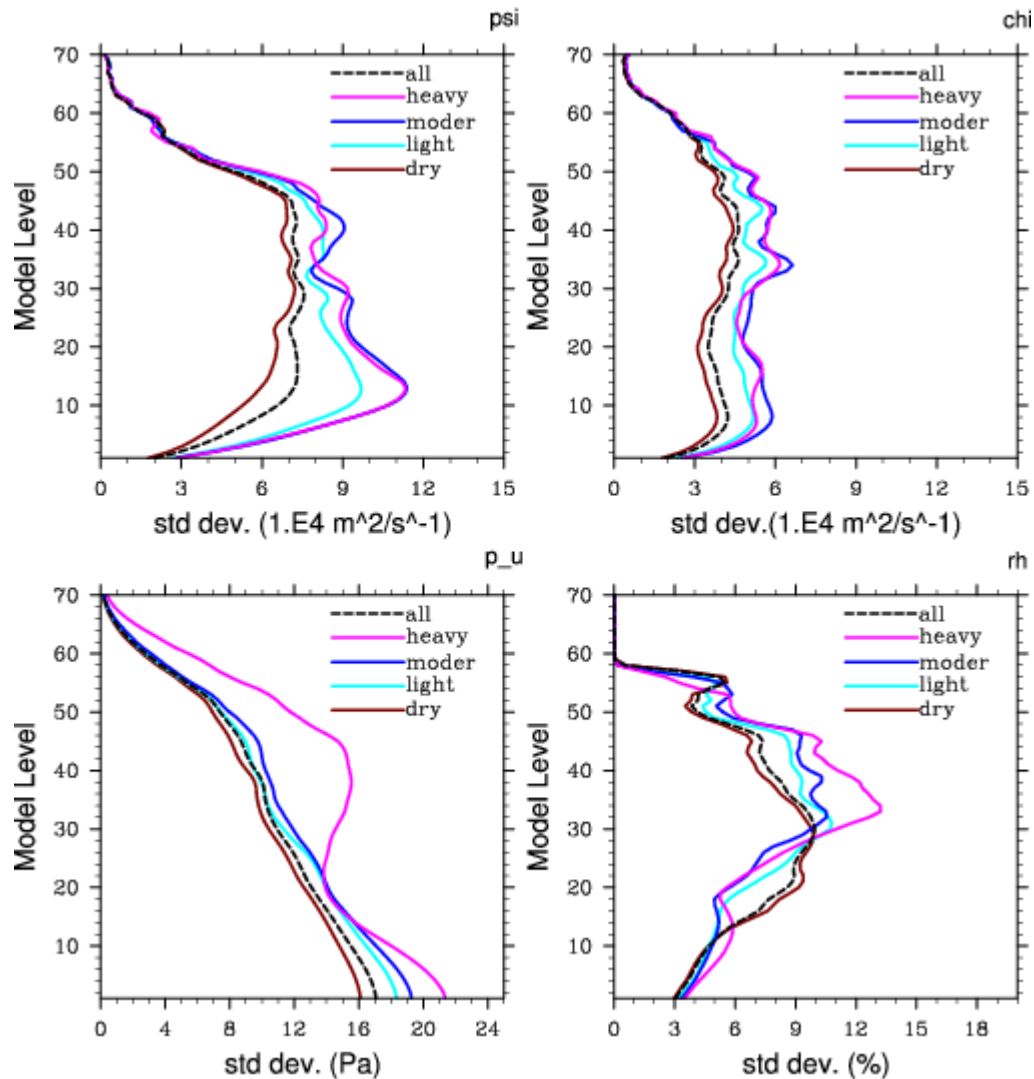
B_{precip} vs $B_{\text{non-precip}}$

Explained pressure variance by the statistical balance operators



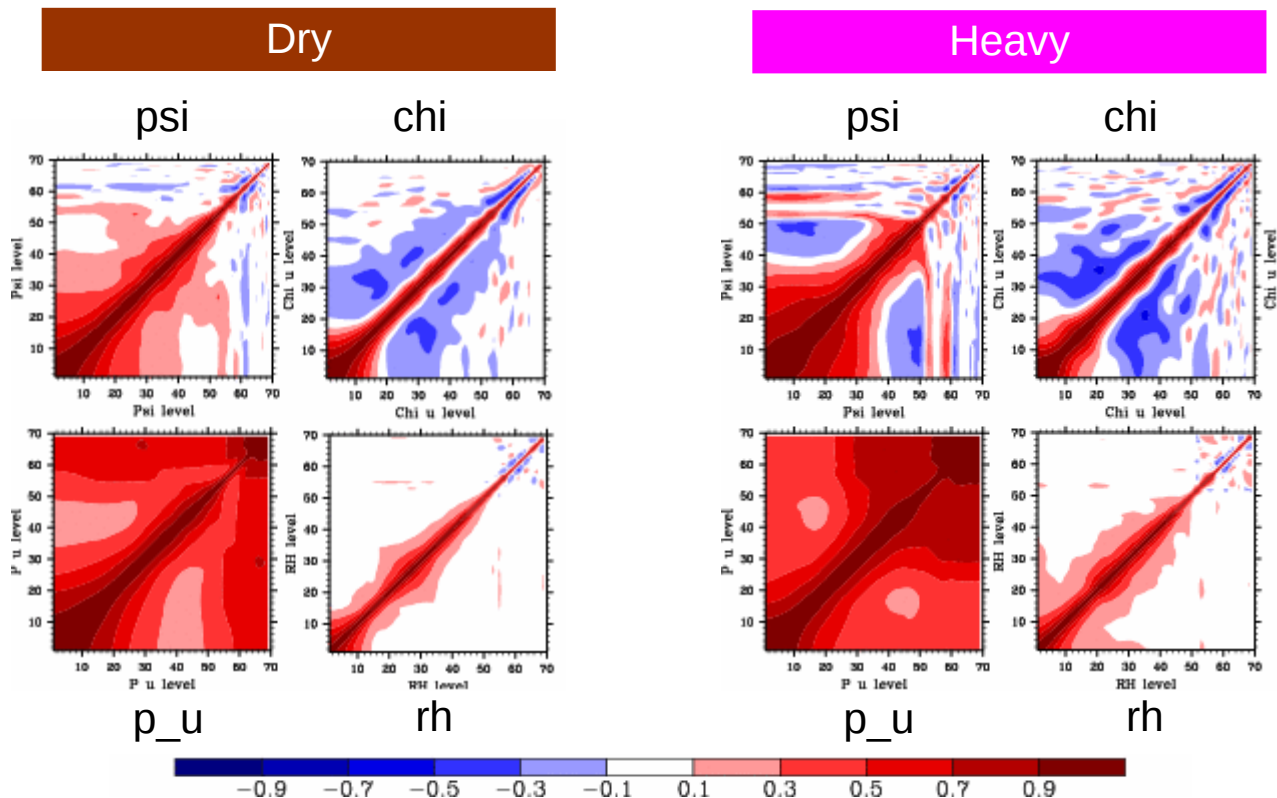
B_{precip} vs $B_{\text{non-precip}}$

Standard deviation profiles



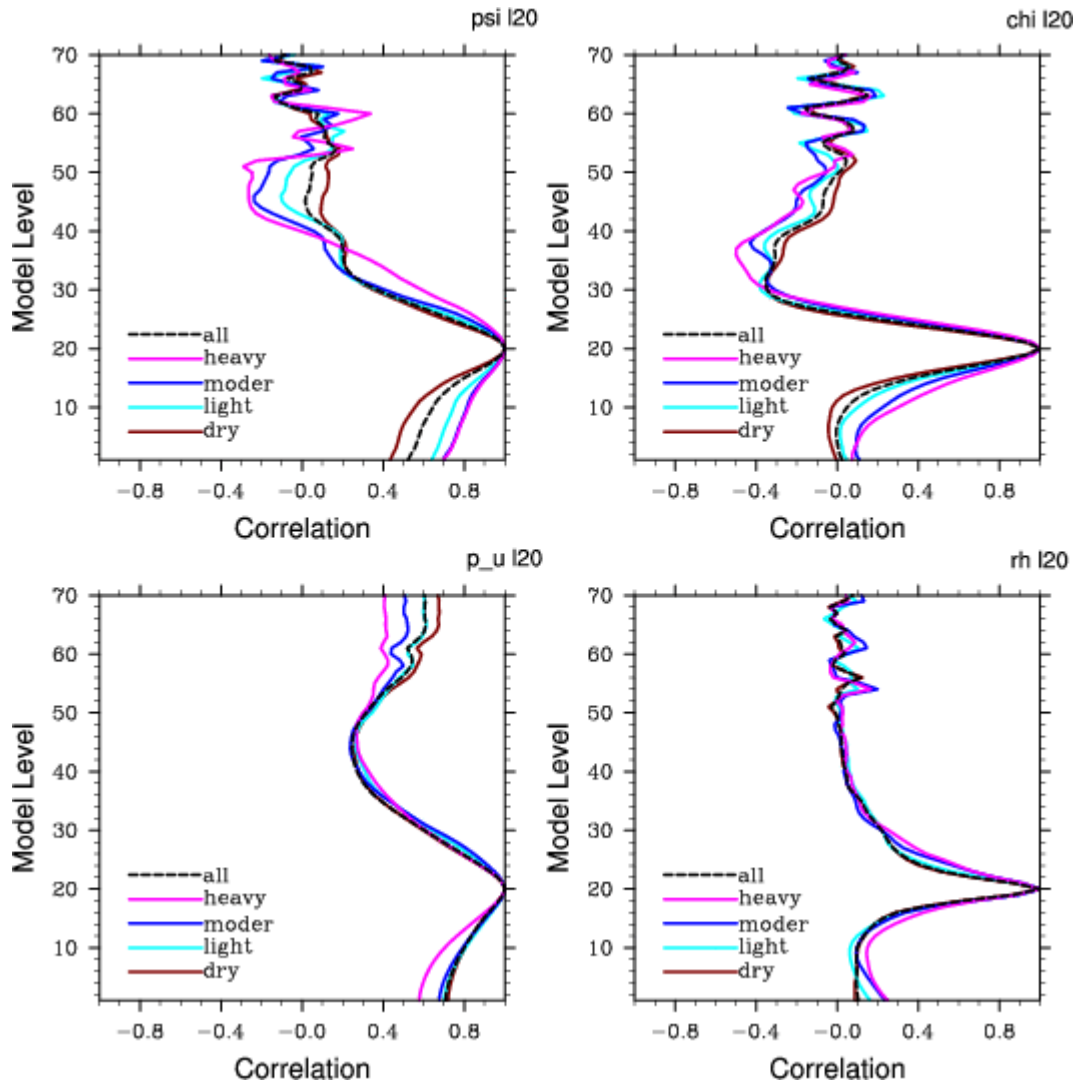
B_{precip} vs $B_{\text{non-precip}}$

Vertical correlation matrices



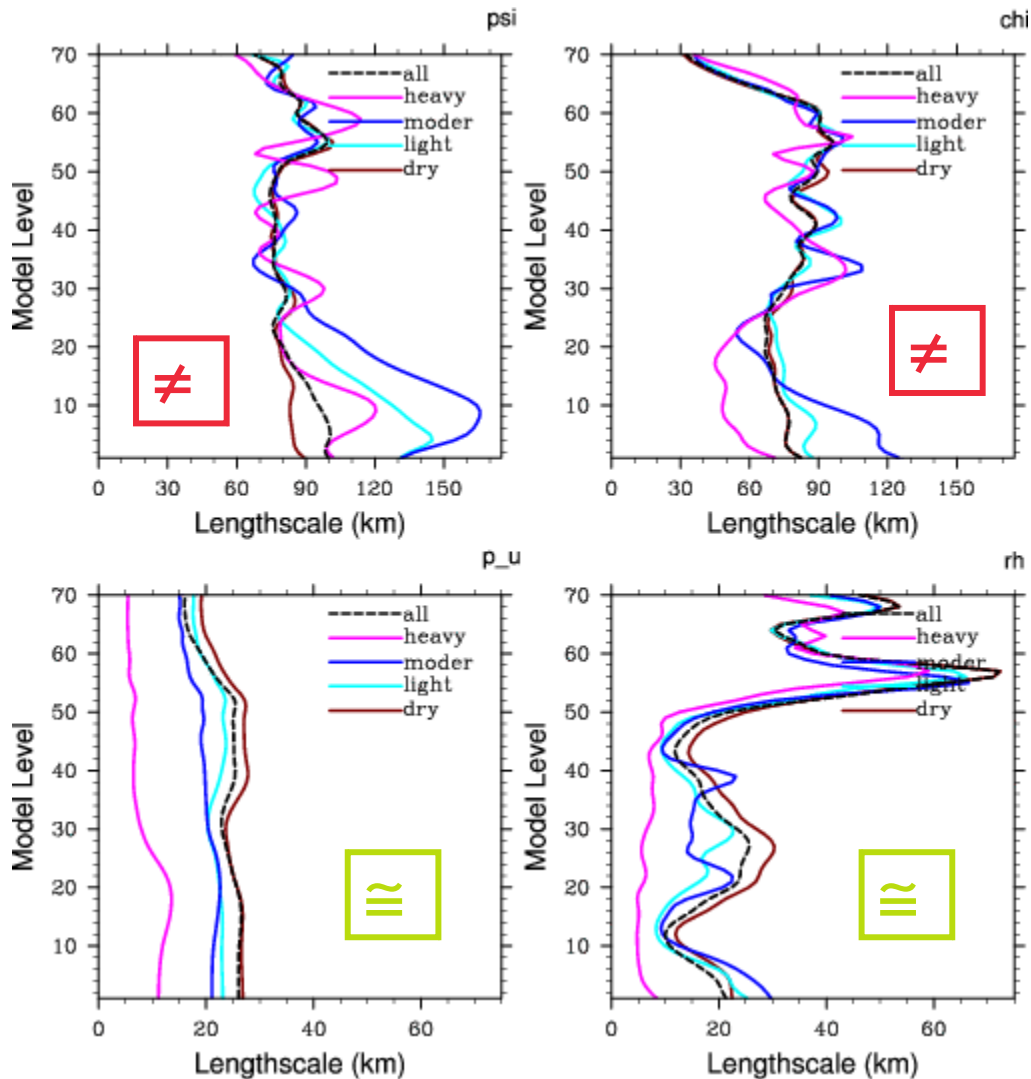
B_{precip} vs $B_{\text{non-precip}}$

Vertical correlations with level 20 (~1.5 km)



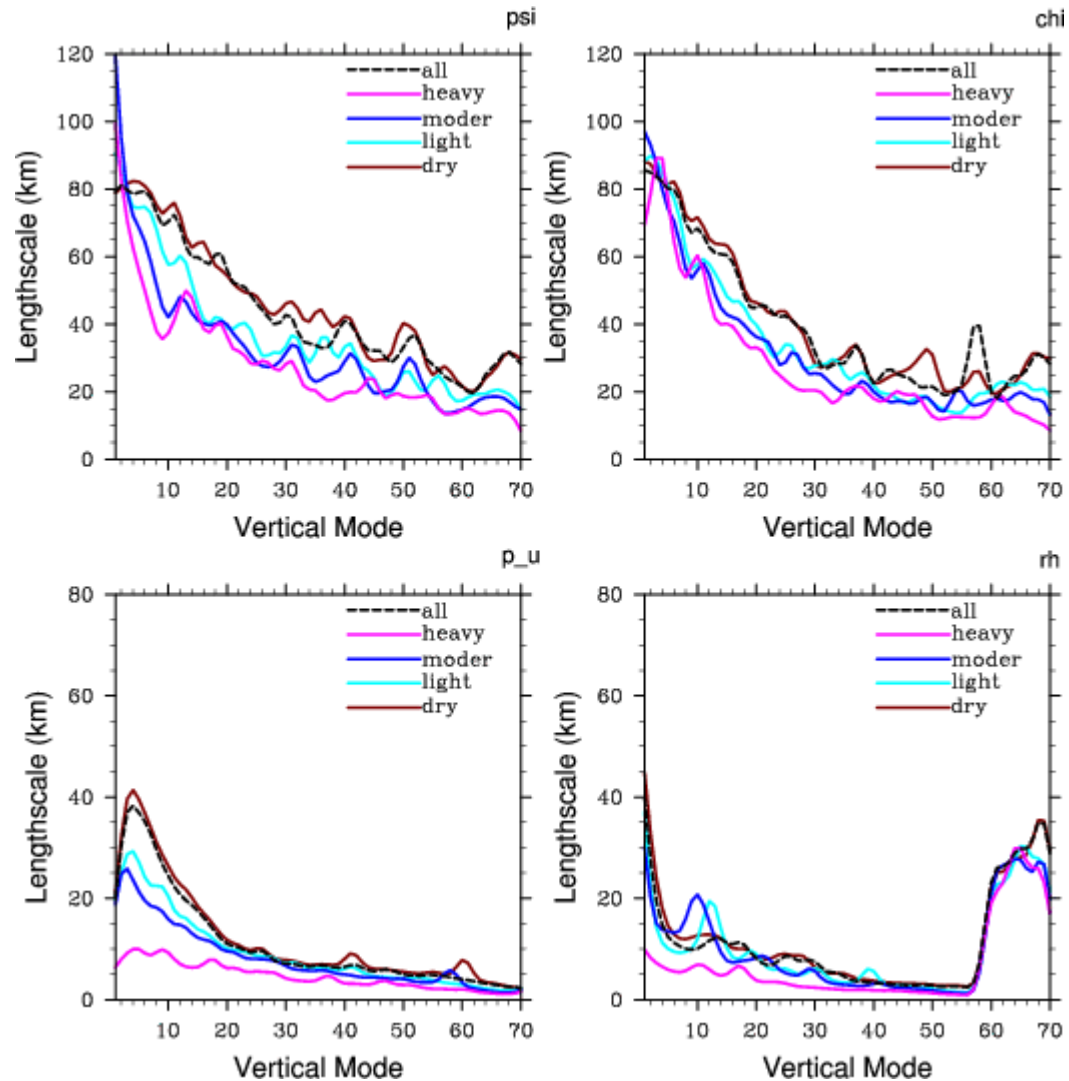
B_{precip} vs $B_{\text{non-precip}}$

SOAR horizontal length scales on **model levels**



B_{precip} vs $B_{\text{non-precip}}$

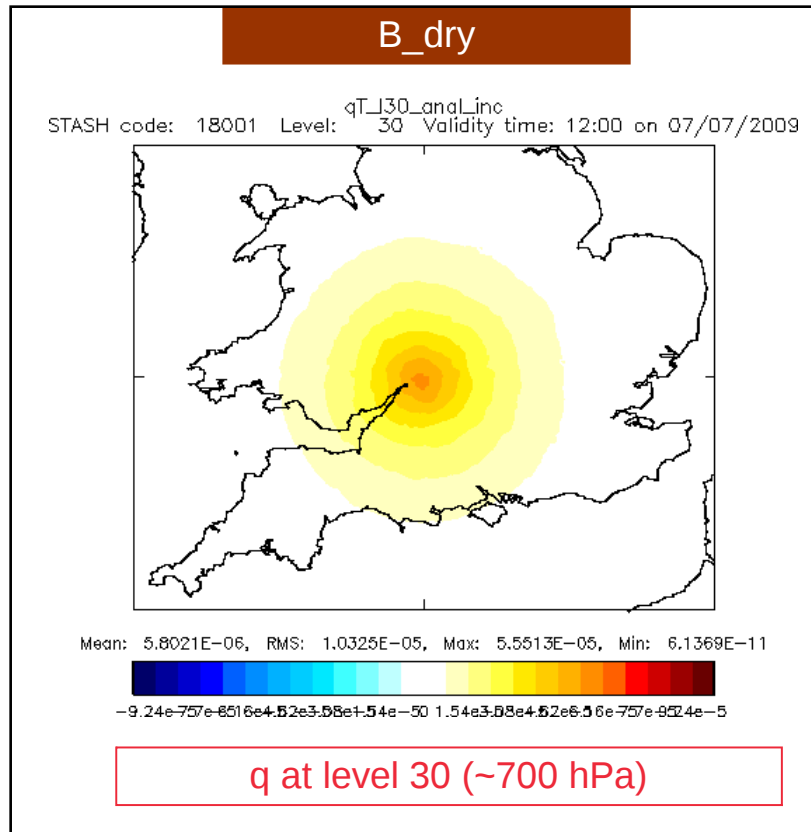
SOAR horizontal length scales in vertical mode space



Single pseudo-obs experiments

Response to the assimilation of a pseudo

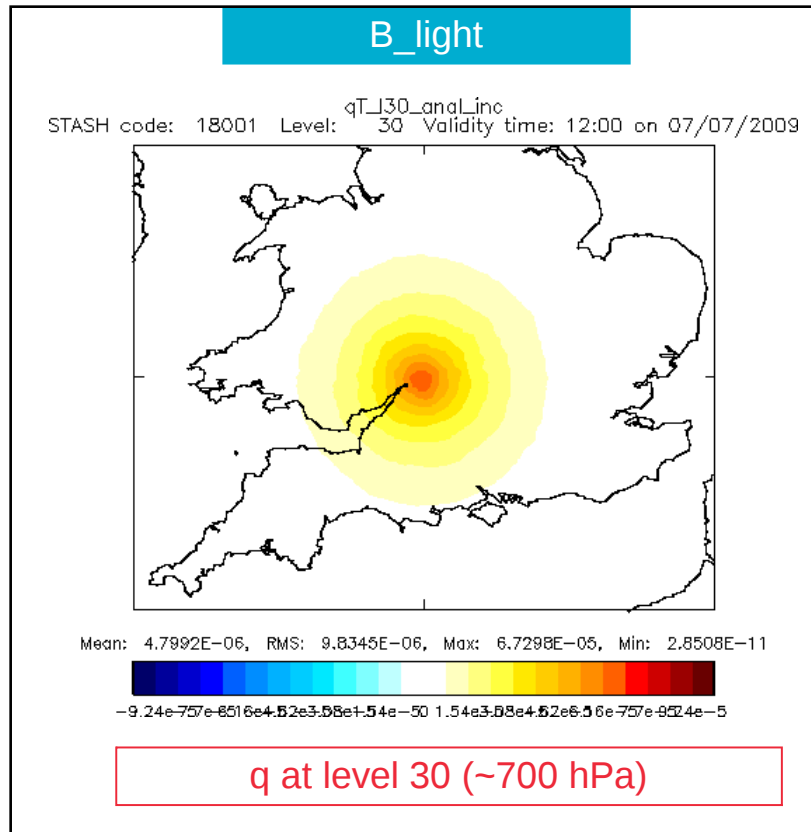
Humidity (q) at level 30 (~700 hPa)



Single pseudo-obs experiments

Response to the assimilation of a pseudo

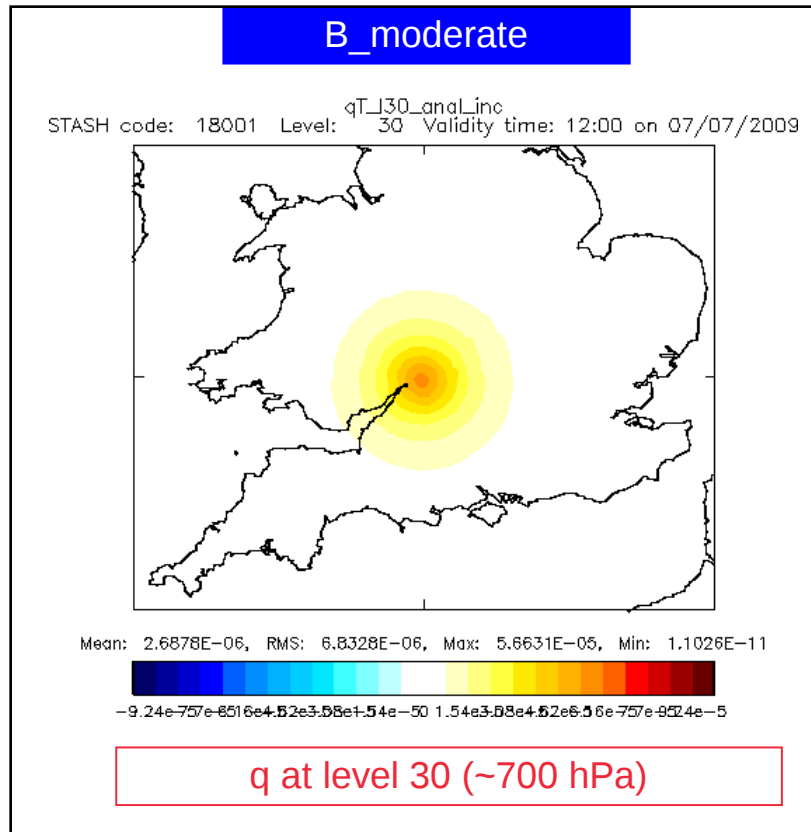
Humidity (q) at level 30 (~700 hPa)



Single pseudo-obs experiments

Response to the assimilation of a pseudo

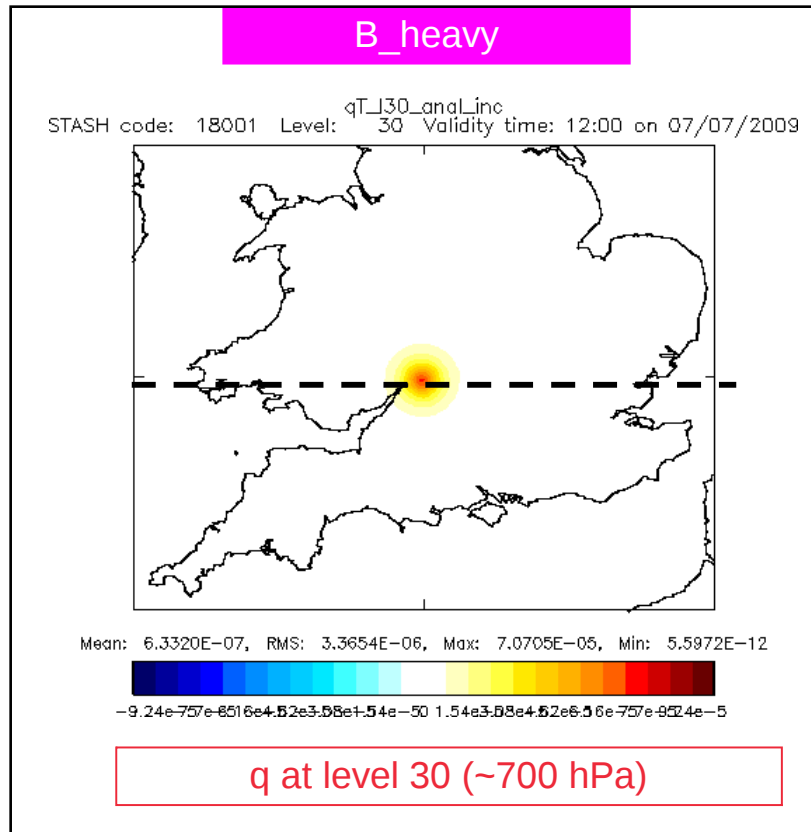
Humidity (q) at level 30 (~700 hPa)



Single pseudo-obs experiments

Response to the assimilation of a pseudo

Humidity (q) at level 30 (~700 hPa)

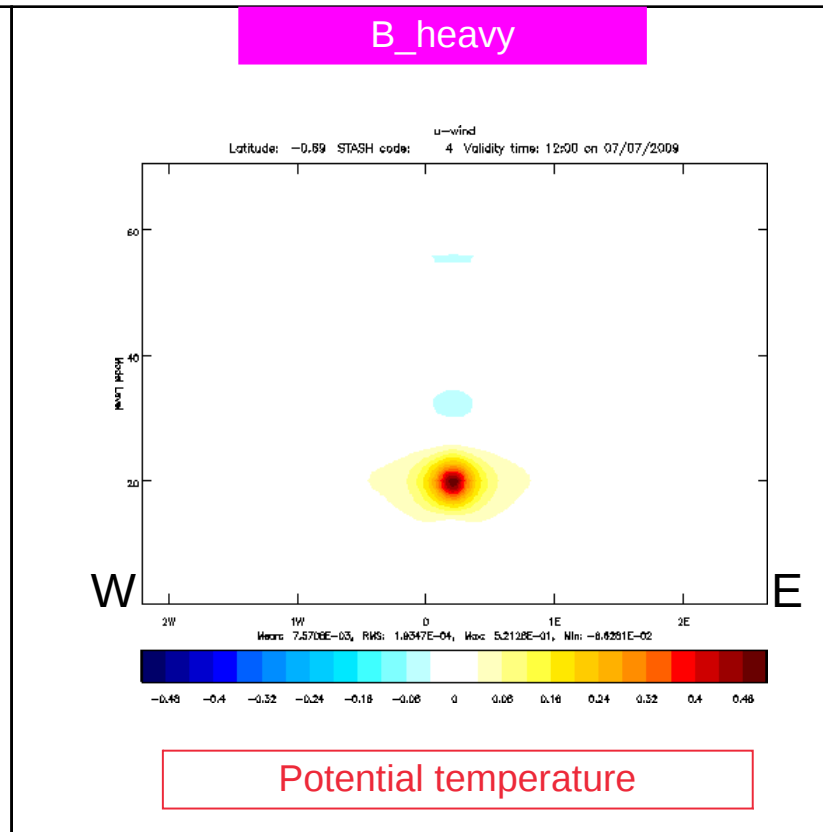
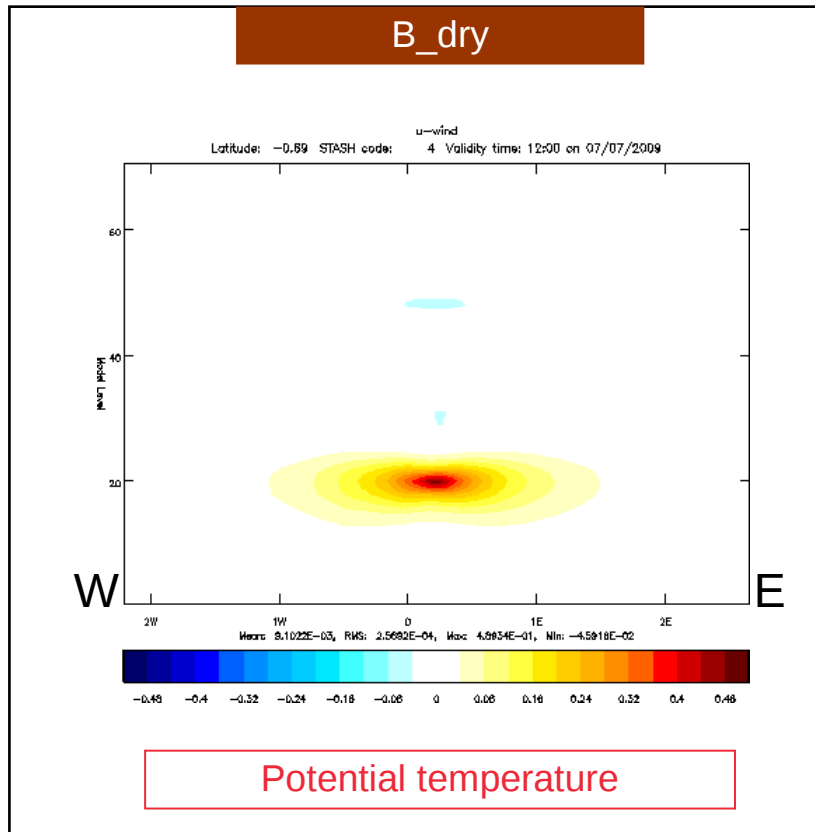


Single pseudo-obs experiments

Response to the assimilation of a pseudo

Temperature at level 20 (~850 hPa)

innovation = 2 K and observation error = 1 K

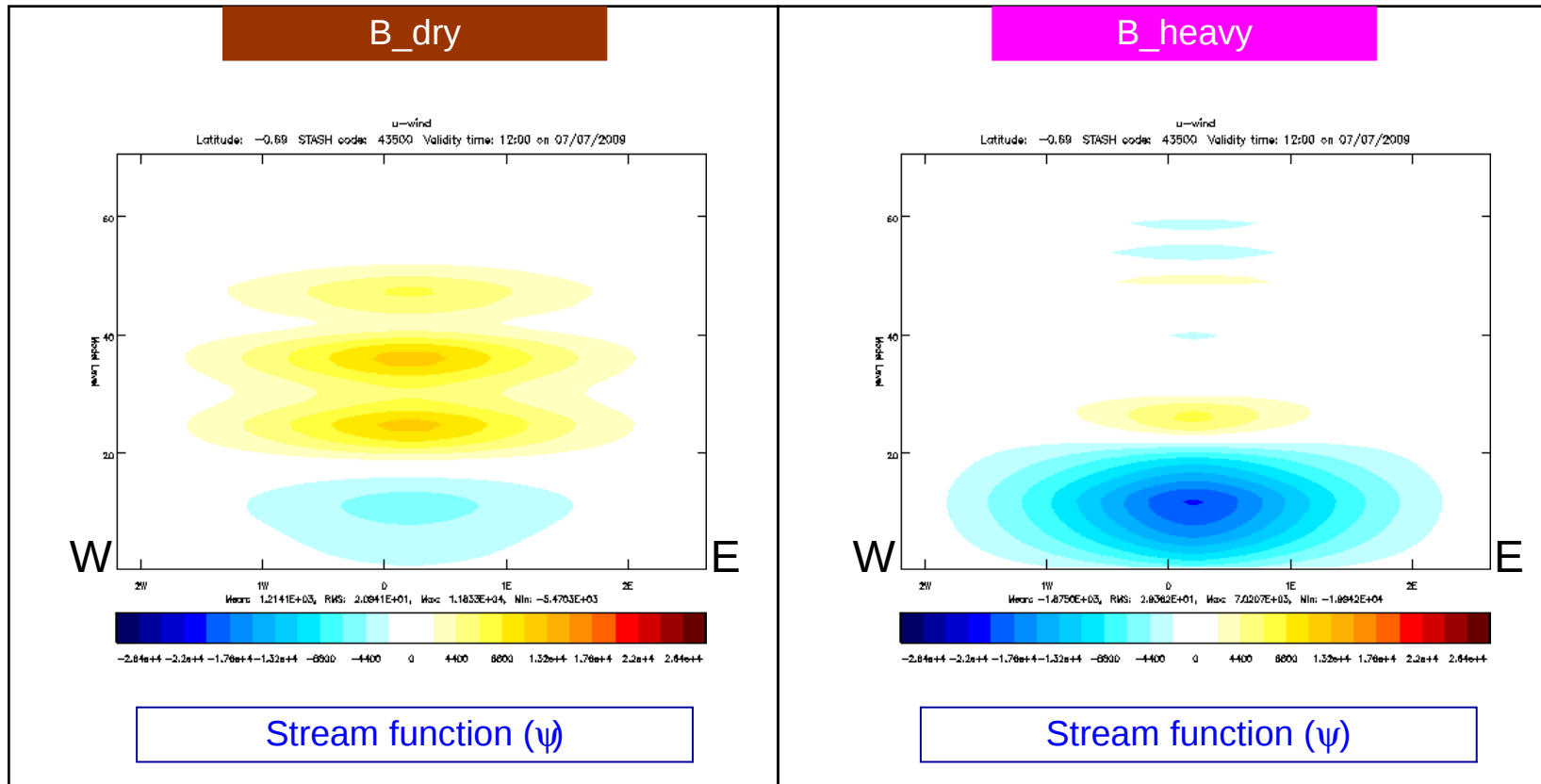


Single pseudo-obs experiments

Response to the assimilation of a pseudo

Temperature at level 20 (~850 hPa)

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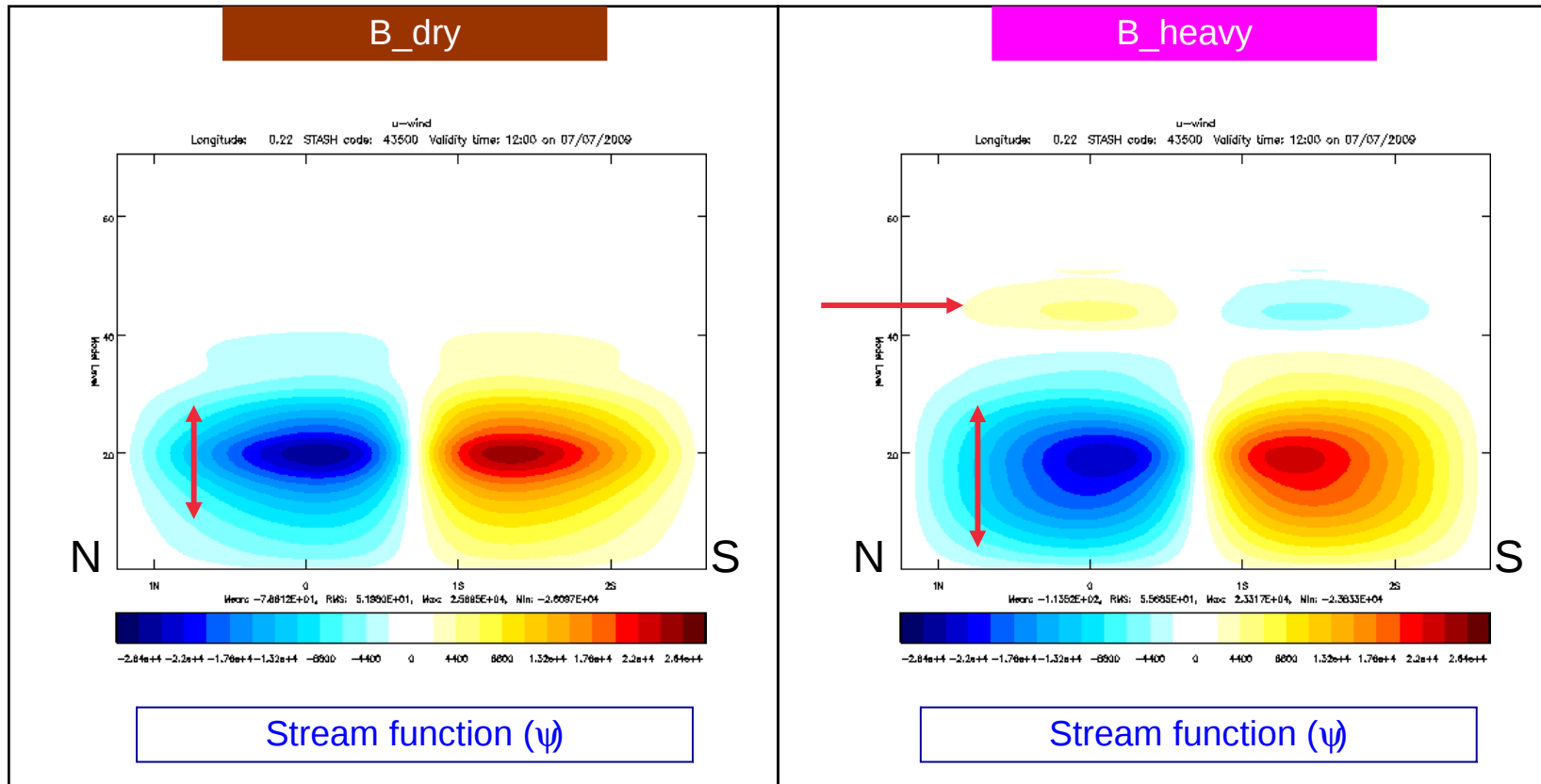


Single pseudo-obs experiments

Response to the assimilation of a pseudo

u-wind at level 20 (~850 hPa)

innovation = 2 m/s and observation error = 1 m/s

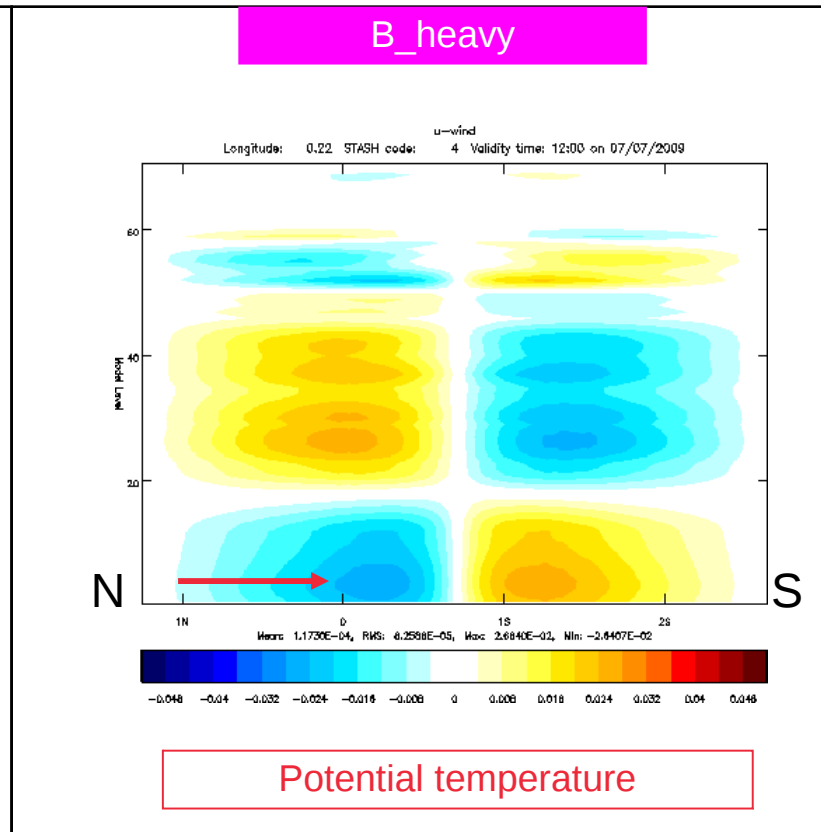
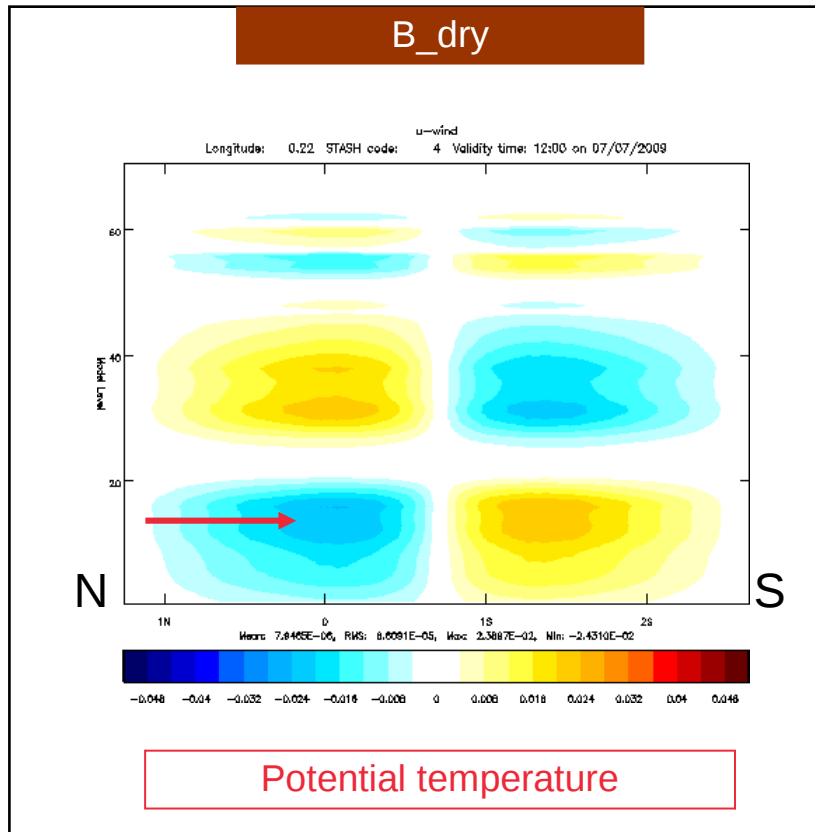


Single pseudo-obs experiments

Response to the assimilation of a pseudo

Temperature at level 20 (~850 hPa)

innovation = 2 K and observation error = 1 K





Summary and future work

- Statistical balance operators can also be improved over precipitation areas at meso- γ scale by using only precipitation profiles in training data.
- Background error covariance exhibit significant differences in precipitation areas compared to dry areas, most differences are proportional with precipitation intensity.

In precipitation areas, we found that...

- Variances are increased.
 - Vertical correlations are either broader or similar.
 - Horizontal length scales for mass fields are smaller. For wind fields: unexpected larger correlations in the low levels.
- Next Steps
 - Test different control variables (e.g. vorticity, divergence).
 - Investigate cross-covariances (e.g. humidity and divergence).
 - Implement B_{precip} approach in a 1DVAR for the retrieval of T and q from radar reflectivities.

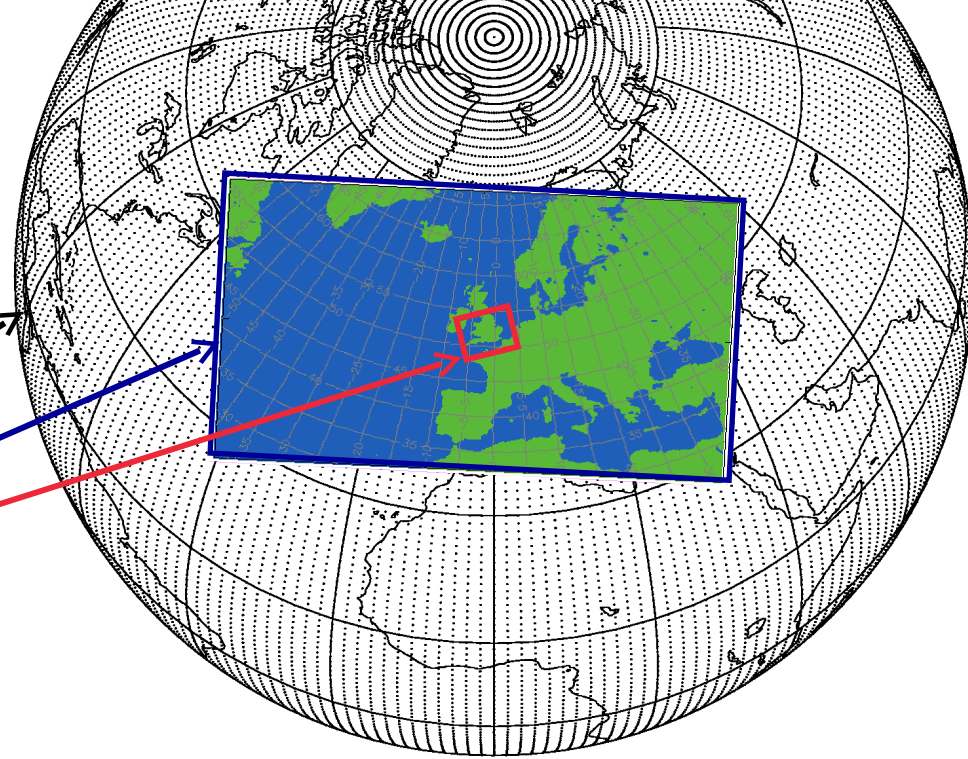


How to optimally treat large scale information in limited area EnDA/EPS ?

Thanks to: Neill Bowler, Mike Cullen, Stefano Migliorini, Sue Ballard and many others...

Introduction

- MOGREPS-G (60 km) - operational
- MOGREPS-R (18 km) - operational
- SUK ETKF (1.5 km) - research

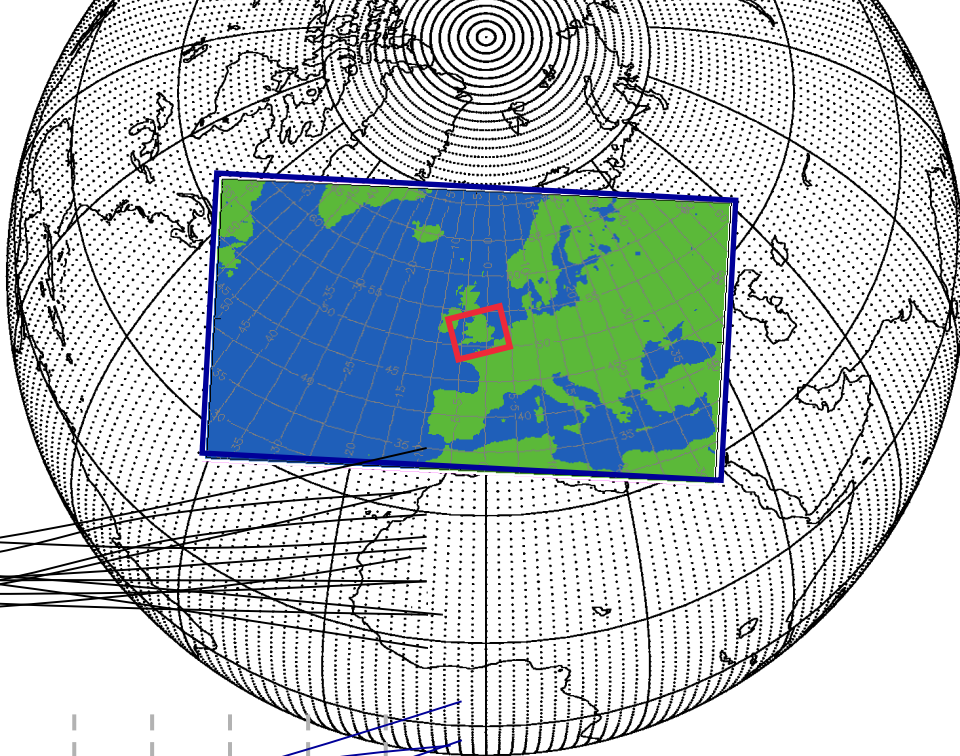


Purpose of this convective-permitting EPS

- Examine 1-h forecast error covariances for the benefit of a NWP-based nowcasting system in development (Bannister et al., 2011, Tellus)
- Predictability studies of very short-term weather events (Migliorini et al., 2011, Tellus)
- Test hybrid VAR DA at convective scale



ETKF 1.5km: The setup



MOGREPS-G
24 members

MOGREPS-R
24 members

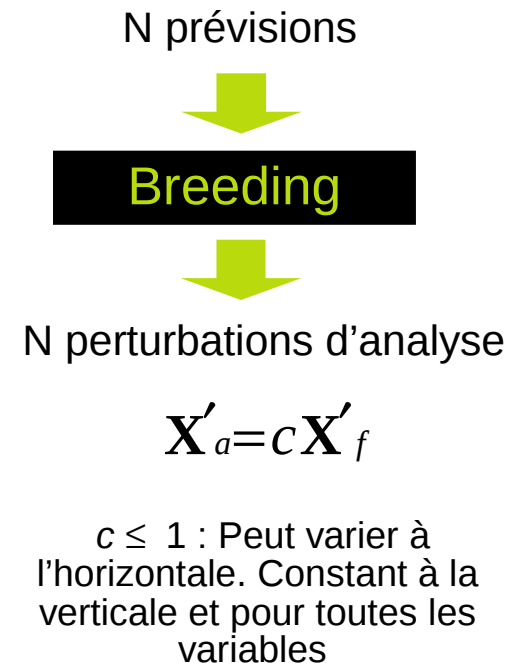
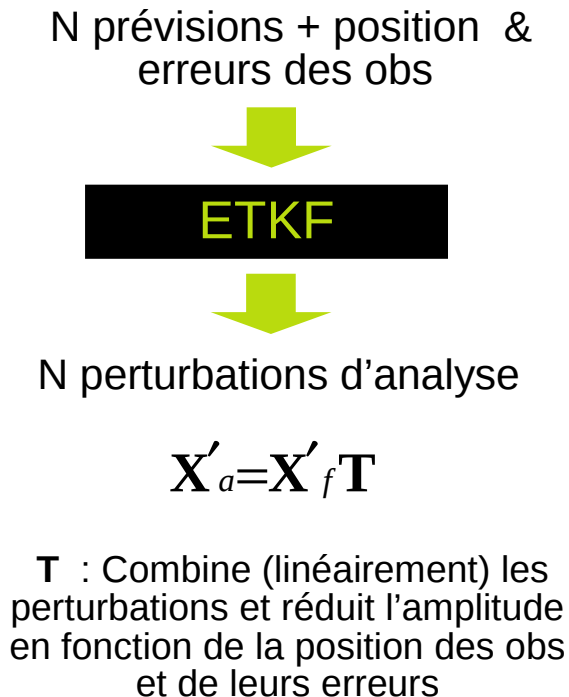
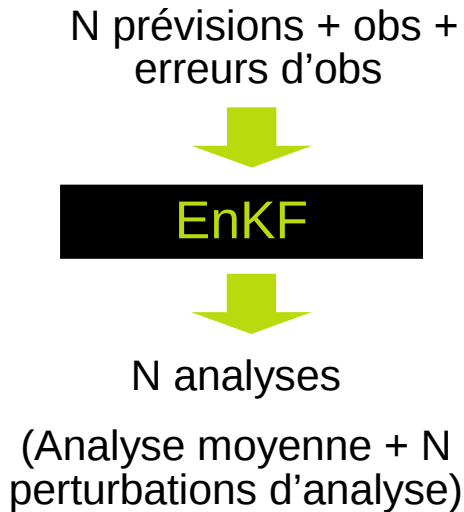
ETKF 1.5km
24 members

IC Perts + LBC

LBC



ETKF ???

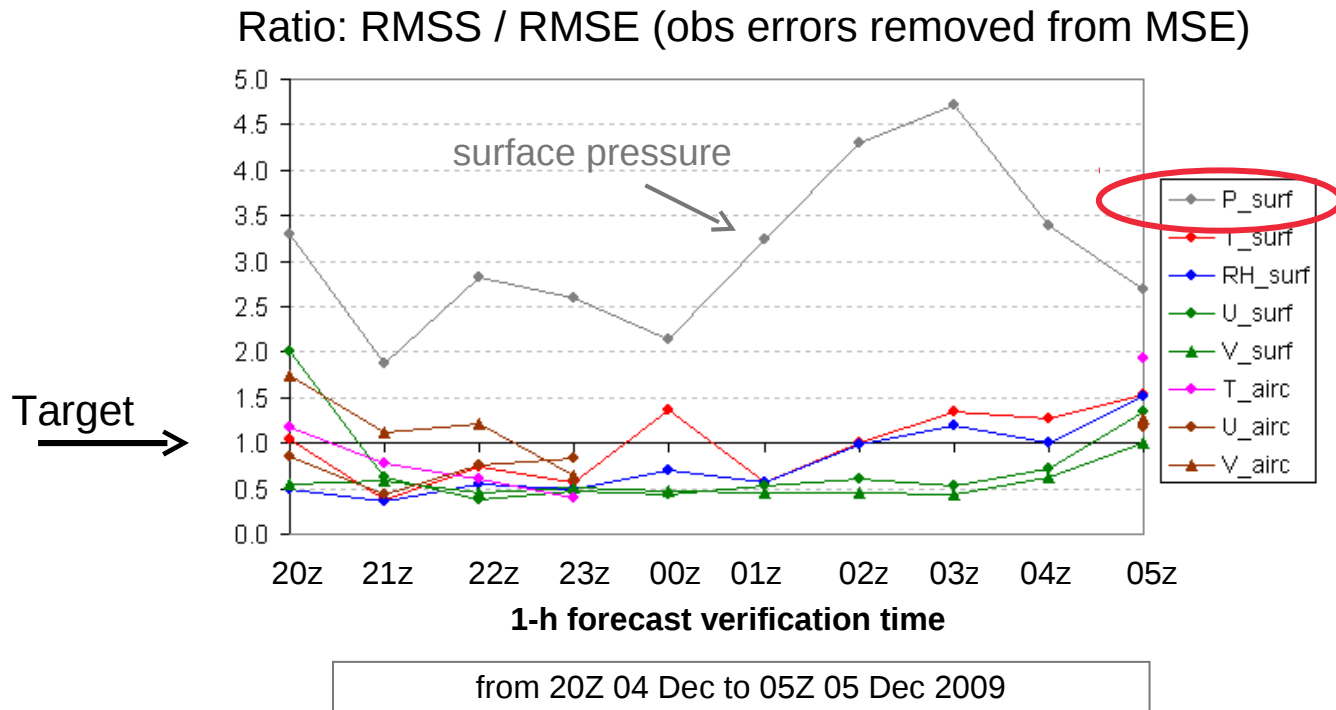


ETKF et Breeding dépendent d'un système d'assimilation pour générer l'analyse de contrôle (l'analyse qui sera perturbée par les N jeux de perturbations)

ETKF 1.5km

Performance assessments

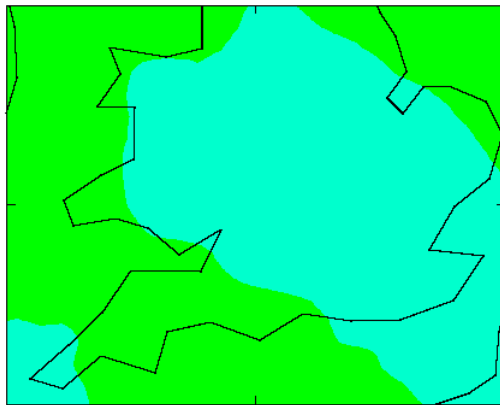
- 3 case studies so far (each case = 11 consecutive 1-h cycle of 1-h forecast only)
- Give satisfactory results except for surface pressure



Overestimation of p_{surf} spread

- Most of the overspread in p_{surf} was found to be attributed to an overestimation in the spread of p_{surf} by MOGREPS-R.
- However p_{surf} spread is increased in the 1.5km ETKF vs MOGREPS-R

MOGREPS-R : 12h fcst



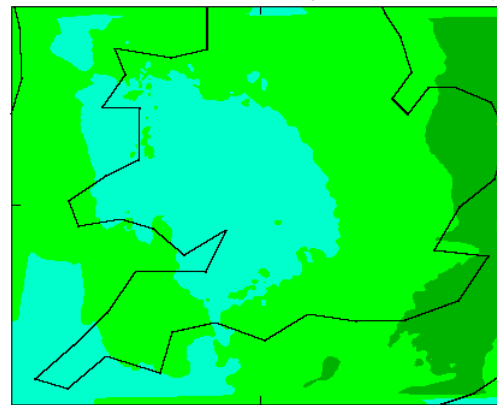
Mean: 1.2582E+00, RMS: 1.2606E+00, Max: 1.5229E+00, Min: 1.0851E+00



0 1 1.8 5

hPa

ETKF 1.5km : 1-h fcst



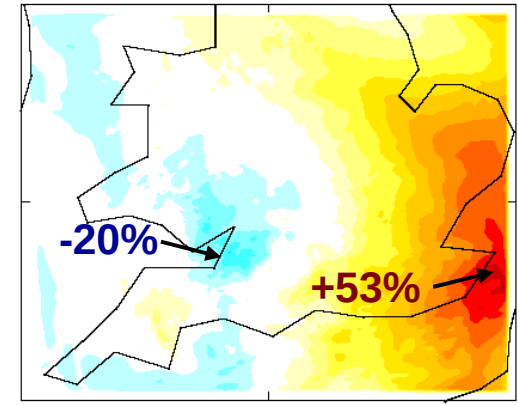
Mean: 1.3560E+00, RMS: 1.3647E+00, Max: 1.8834E+00, Min: 1.0040E+00



0 1 1.8 5

hPa

Relative difference (%)



Mean: 8.2522E+00, RMS: 1.6899E+01, Max: 5.2649E+01, Min: -2.0125E+01



-60 -30 0 25 60

%

Domain average = +9.1%

fields valid at 06z 05/12/2009

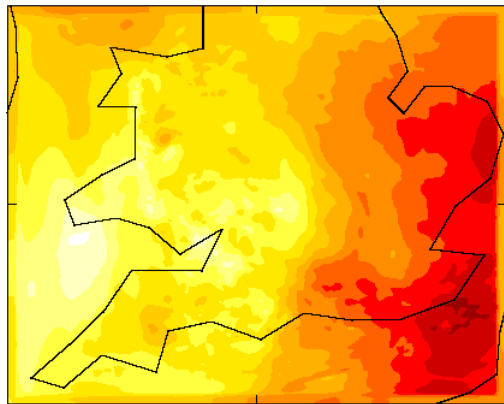
Overestimation of p_{surf} spread

- Why the p_{surf} spread is increased in the 1.5km ETKF ?

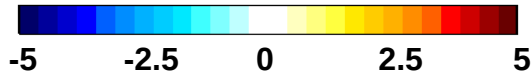
1-h p_{surf} perturbations for member #8

Contribution from LBC perturbations only

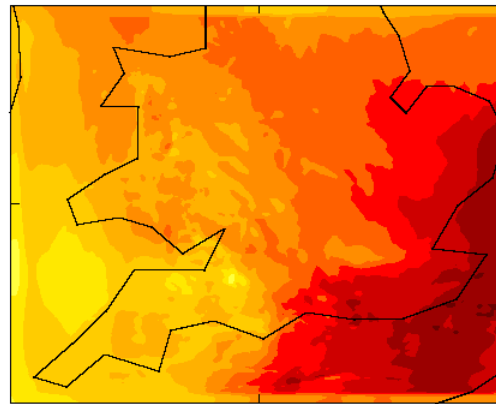
Contribution from IC perturbations only



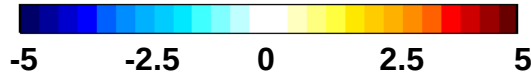
Mean: 2.2331E+00, RMS: 2.4039E+00, Max: 4.2885E+00, Min: 5.1152E-02



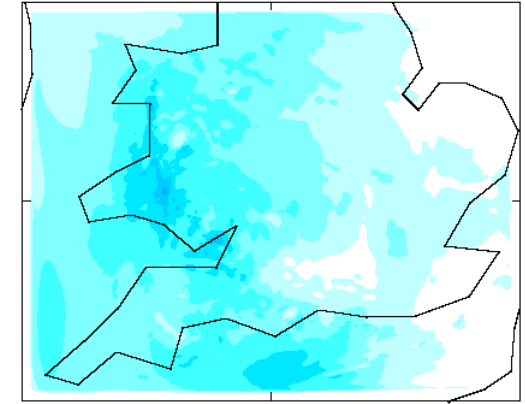
hPa



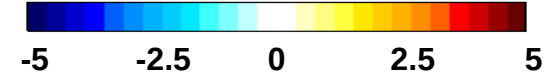
Mean: 2.9571E+00, RMS: 3.0402E+00, Max: 4.6920E+00, Min: 1.1728E+00



hPa



Mean: -7.5403E-01, RMS: 8.8387E-01, Max: 3.0273E-01, Min: -2.3842E+00



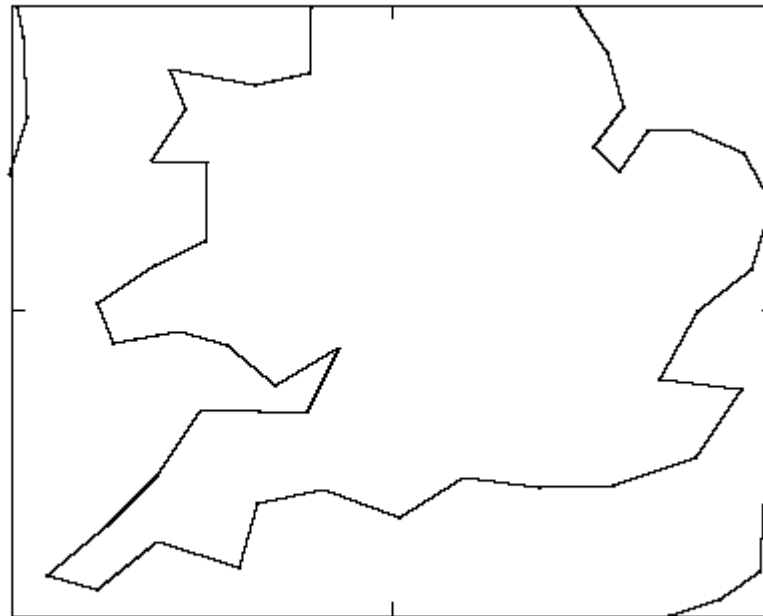
hPa

1h forecast valid at 06z 05/12/2009

Overestimation of p_{surf} spread

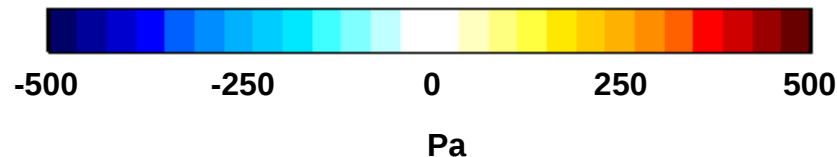
- Consequence of perturbing only the LBC

ETKF - pert08 lbc onlySurface pressure pert. (Pa)
Validity time: 04:30 on 05/12/09



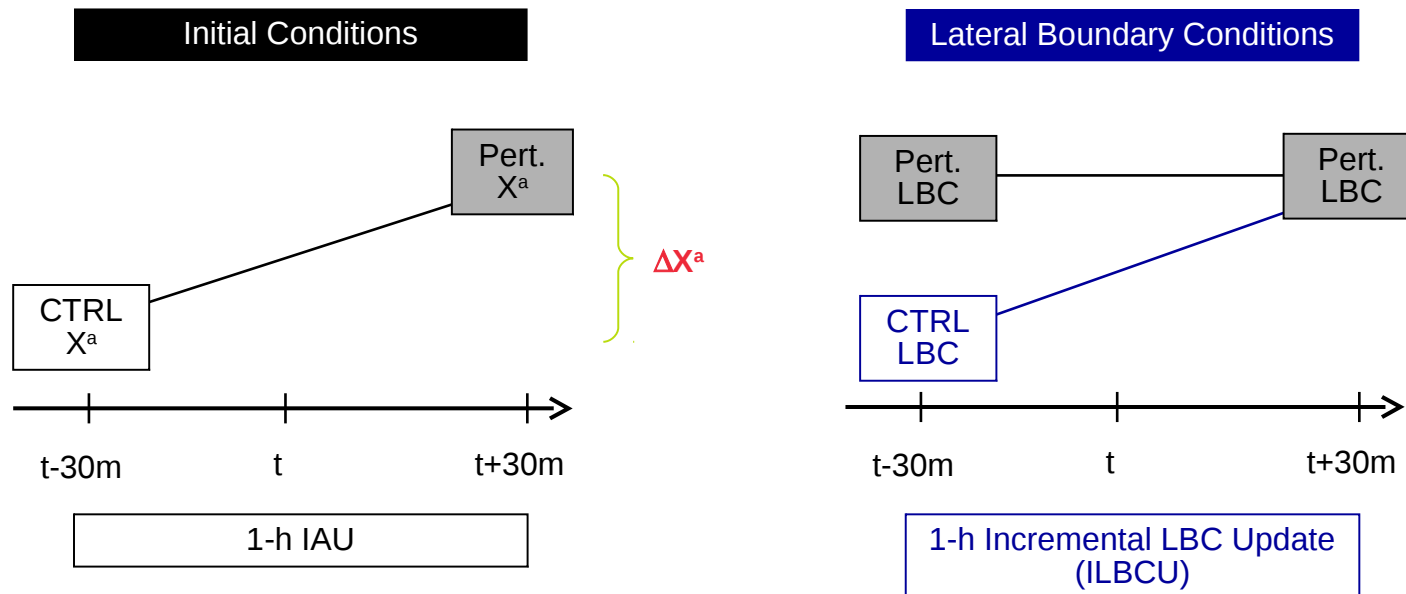
Discontinuities between IC and LBC can introduce significant spurious perturbations.

Mean: 0.0000E+00, RMS: 0.0000E+00, Max: 0.0000E+00, Min: 0.0000E+00



Sources of discontinuities at the LB (1)

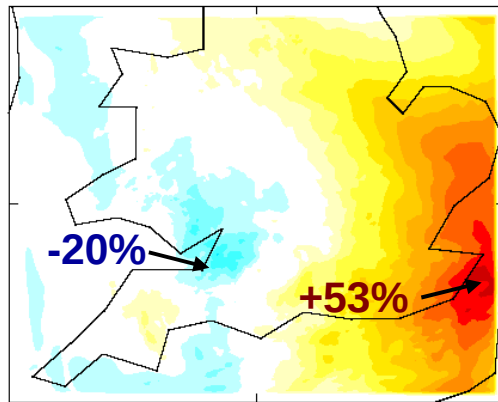
- The Incremental Analysis Update (IAU, i.e., how we add the IC perturbations)



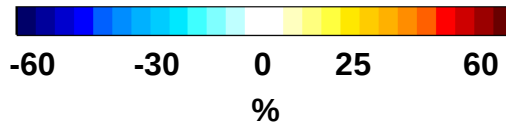
Sources of discontinuities at the LB (1)

- Relative difference in p_{surf} spread : 1.5km ETKF vs. MOGREPS-R

IAU only

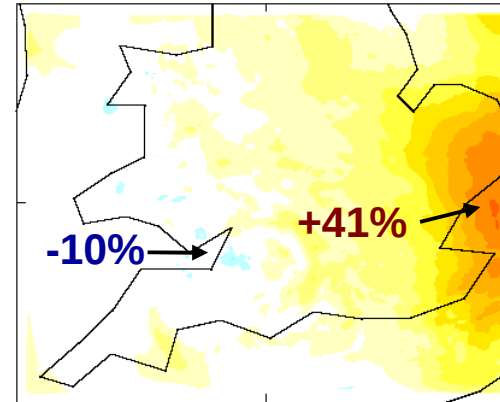


Mean: 8.2522E+00, RMS: 1.6899E+01, Max: 5.2649E+01, Min: -2.0125E+01

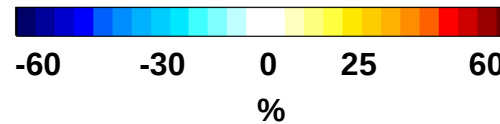


Domain average = +9.1%

IAU + ILBCU



Mean: 7.4320E+00, RMS: 1.1717E+01, Max: 4.0738E+01, Min: -1.0186E+01

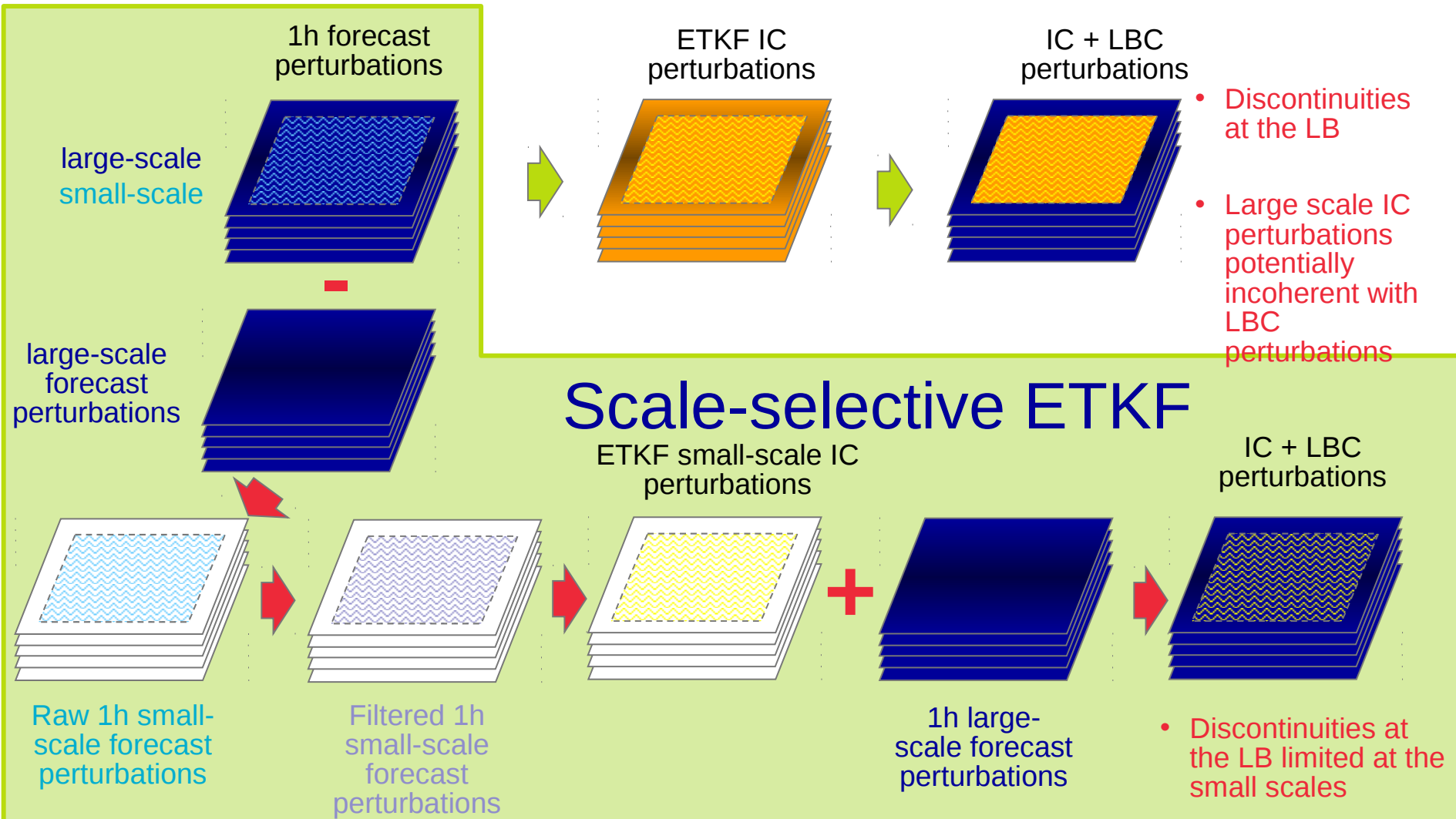


Domain average = +8.1%

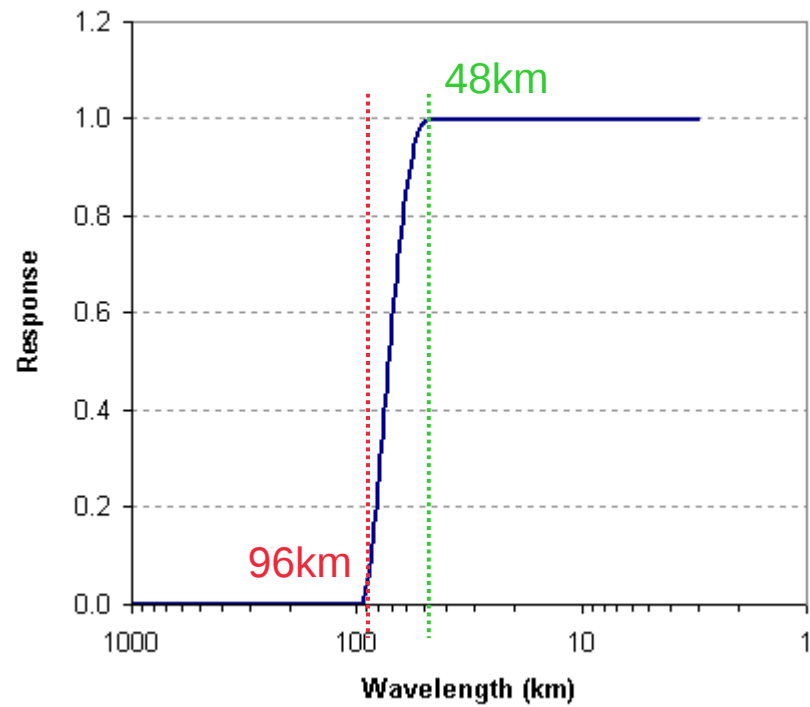
Other sources of discontinuities between IC and the LBC must be present

1h forecast valid at 06z 05/12/2009

Sources of discontinuities at the LB (2)



Filtering





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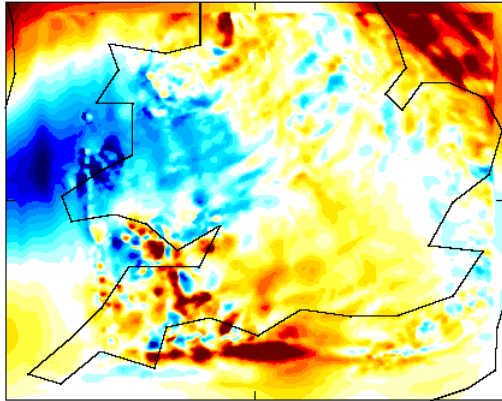
1-h forecast perturbations

Full

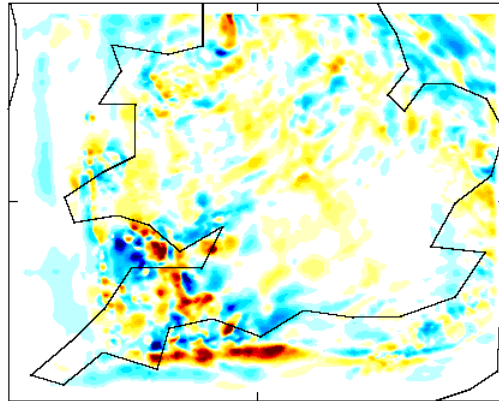
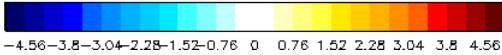
Raw small-scale (residual)

Filtered small-scale

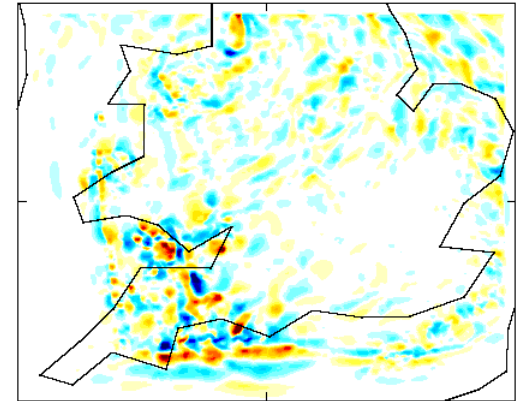
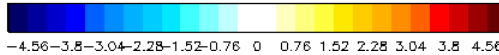
u-wind at 2km



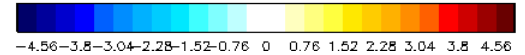
Mean: 5.0203E-01, RMS: 1.8934E+00, Max: 9.1189E+00, Min: -6.1481E+00



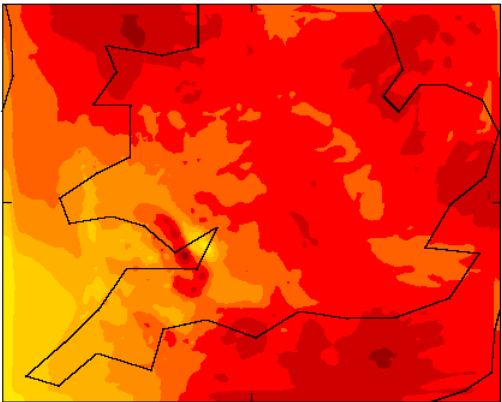
Mean: 6.7872E-03, RMS: 8.7277E-01, Max: 5.9133E+00, Min: -5.9606E-01



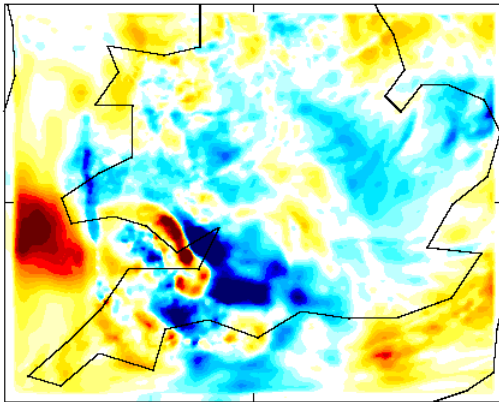
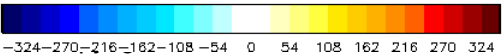
Mean: 4.6531E-03, RMS: 6.8740E-01, Max: 5.5655E+00, Min: -5.7321E+00



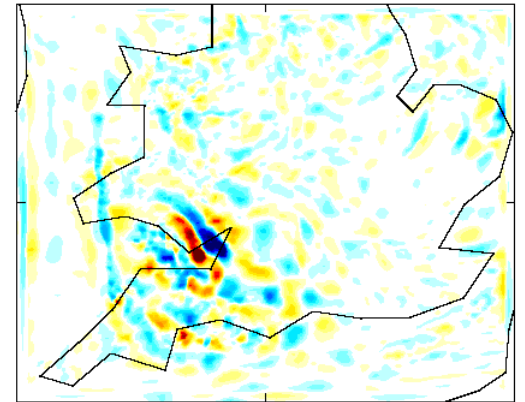
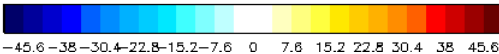
Surface pressure



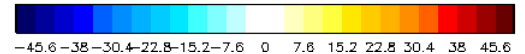
Mean: 2.4055E+02, RMS: 2.4320E+02, Max: 3.1056E+02, Min: 1.1156E+02



Mean: -5.6165E-02, RMS: 1.5156E+01, Max: 7.7272E+01, Min: -1.2420E+02



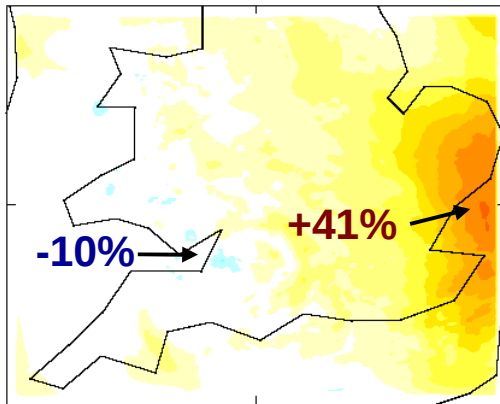
Mean: -3.5631E-02, RMS: 6.4901E+00, Max: 7.8236E+01, Min: -6.3901E+01



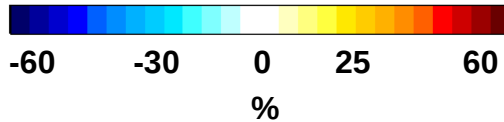
Scale-selective ETKF

- Relative difference in p_{surf} spread : 1.5km EPS vs. MOGREPS-R

1.5km ETKF

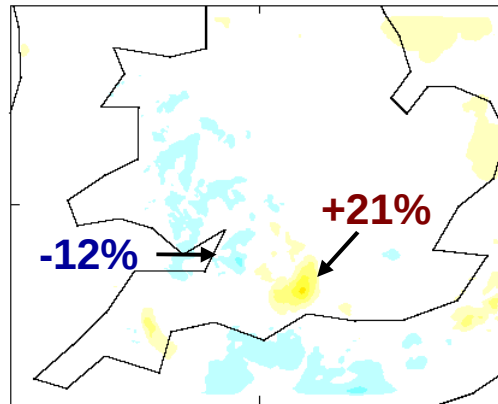


Mean: 7.4320E+00, RMS: 1.1717E+01, Max: 4.0738E+01, Min: -1.0186E+01

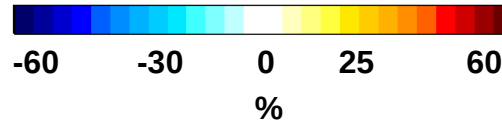


Domain average = +8.1%

Scale-selective 1.5km ETKF

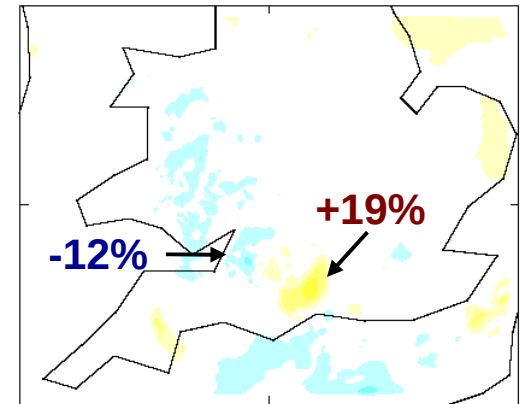


Mean: -1.9941E-01, RMS: 3.1145E+00, Max: 2.1205E+01, Min: -1.2200E+01

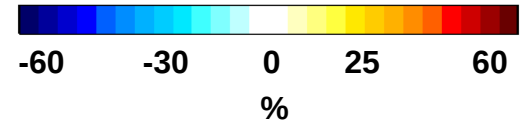


= -0.2%

Downscaled perturbations from MOGREPS-R



Mean: -1.6971E-01, RMS: 3.0999E+00, Max: 1.9610E+01, Min: -1.2429E+01

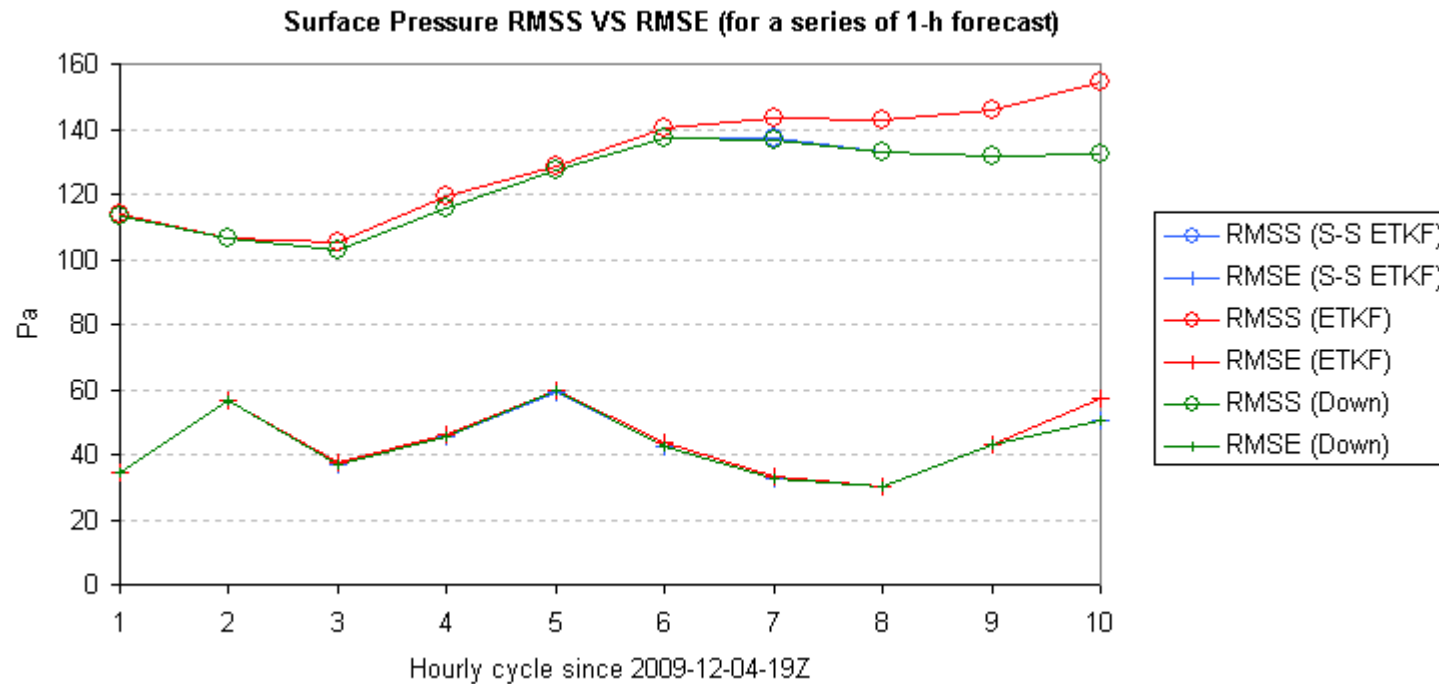


= -0.2%

1h forecast valid at 06z 05/12/2009

Scale-selective ETKF

- 1-h p_{surf} spread vs RMSE* in observation space



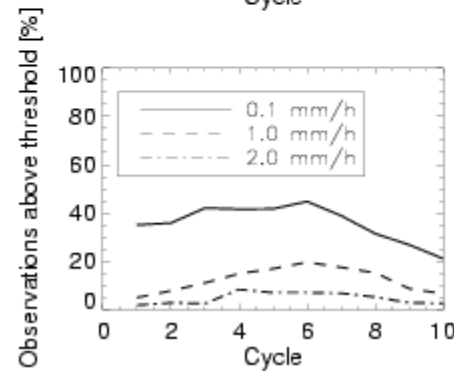
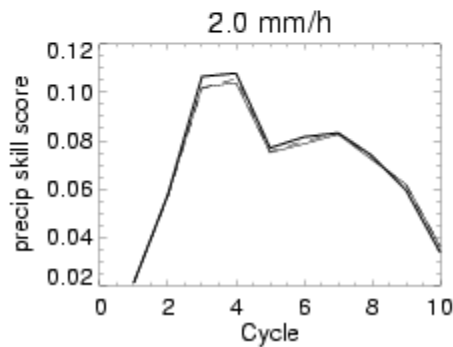
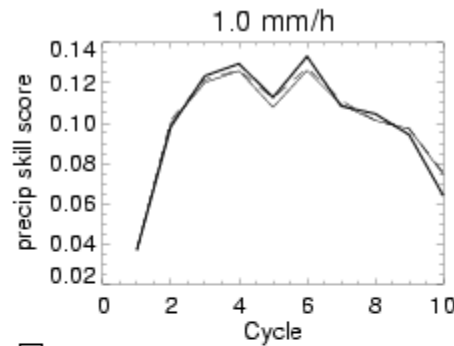
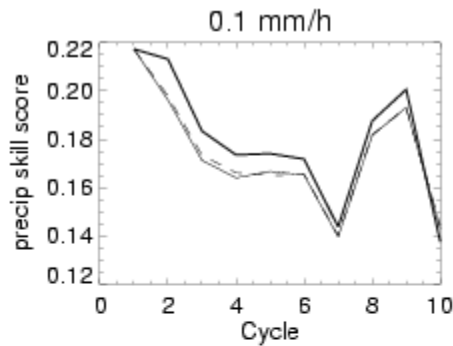
* The observation error was removed from the MSE



Scale-selective ETKF

- Verification of 1-h precipitation rate – Brier Score

↓
A lower value is a better EPS forecast



— ETKF
— S-S ETKF
- - - Downscaling



Summary and discussion

- The current ETKF approach for limited area generates discontinuity between IC and LBC perturbations

This is likely to be true for all current limited area EnDA approach

- In our small domain, discontinuities were found to introduce significant spurious perturbations in the pressure field.

This is likely to be less important in larger domains.

- Results from the scale-selective ETKF suggest that preventing discontinuities at the resolved scales of the driving EPS can remove these spurious perturbations.
- Scale-selective ETKF improves slightly some other variables w.r.t. to the ETKF but... so far no clear benefit was found with respect to downscaled perturbations.



Summary and discussion

The experiment conducted here is far from being optimal...

- Large scales = scales resolved by the driving EPS
Where to set the line between large-scale and small-scale ?
- The continuity of the spectra of the resulting IC perturbations was not inspected.
- No localization was applied on the small scale forecast perturbations
- An arbitrary fixed inflation factor was applied on the small scale IC perturbations
Designing a variable inflation factor will not be trivial
- 1-h IAU was used. Probably too long for the small scale perturbations.

My main concern with the scale-selective ETKF

- The small scale IC perturbations are constructed without the knowledge of the large scale perturbations
The small scale IC perturbations could potentially be incoherent with the large scale component.



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Questions

VAR covariance model: MetO vs. EC

Steps from model variable space to covariance model space

1

2

3

4

Met Office	
\mathbf{x}'	$u', v', \theta', \rho', p', q'$
Parameter transform $\mathbf{v}_p = \mathbf{T}_p \mathbf{x}'$	Global and LAM ψ', χ', p'_u, r' $p'_u = p'_h - \mathbf{F} f(\psi)$
Vertical transform \mathbf{T}_v	auto covariances only $\mathbf{B}_v(i, j) = \mathbf{E} \Lambda \mathbf{E}^T$ $\mathbf{v}_p(i, j, k) \rightarrow \mathbf{v}_v(i, j, m)$
Horizontal transform \mathbf{T}_h	$\mathbf{B}_h(r) = \mathbf{B}_h(0) \left(1 + \frac{r}{s}\right) e^{r/s}$ LAM $\mathbf{v}_v(i, j, m) \rightarrow \mathbf{v}_h(l, n, m)$ Global

EC	
\mathbf{x}'	u', v', T', p'_s, q'
Parameter transform $\mathbf{v}_p = \mathbf{T}_p \mathbf{x}'$	$\psi', \chi'_u, T'_u, p'_{s_u}, q'$ Global $(T'_u, p'_{s_u}) = (T', p'_s) - \mathbf{F} f(\psi)$ $\chi'_u = \psi - \mathbf{G} \psi$
Horizontal transform \mathbf{T}_h	$\mathbf{v}_p(i, j, k) \rightarrow \mathbf{v}_h(l, n, k)$
Vertical transform \mathbf{T}_v	$\mathbf{B}_v(l, n)$ Global: auto covariances LAM: auto and cross covariances



Outline

- Description of a 1.5 km ETKF-based limited area EPS for research purpose
- Issues related to discontinuities at the large scale between IC and LBC perturbations
 - * In this presentation the term “large scale” will refer to the scales resolved by the parent (or driving) EPS
- How to improve the representation of the large scale IC perturbations ? Some results using a scale-selective ETKF

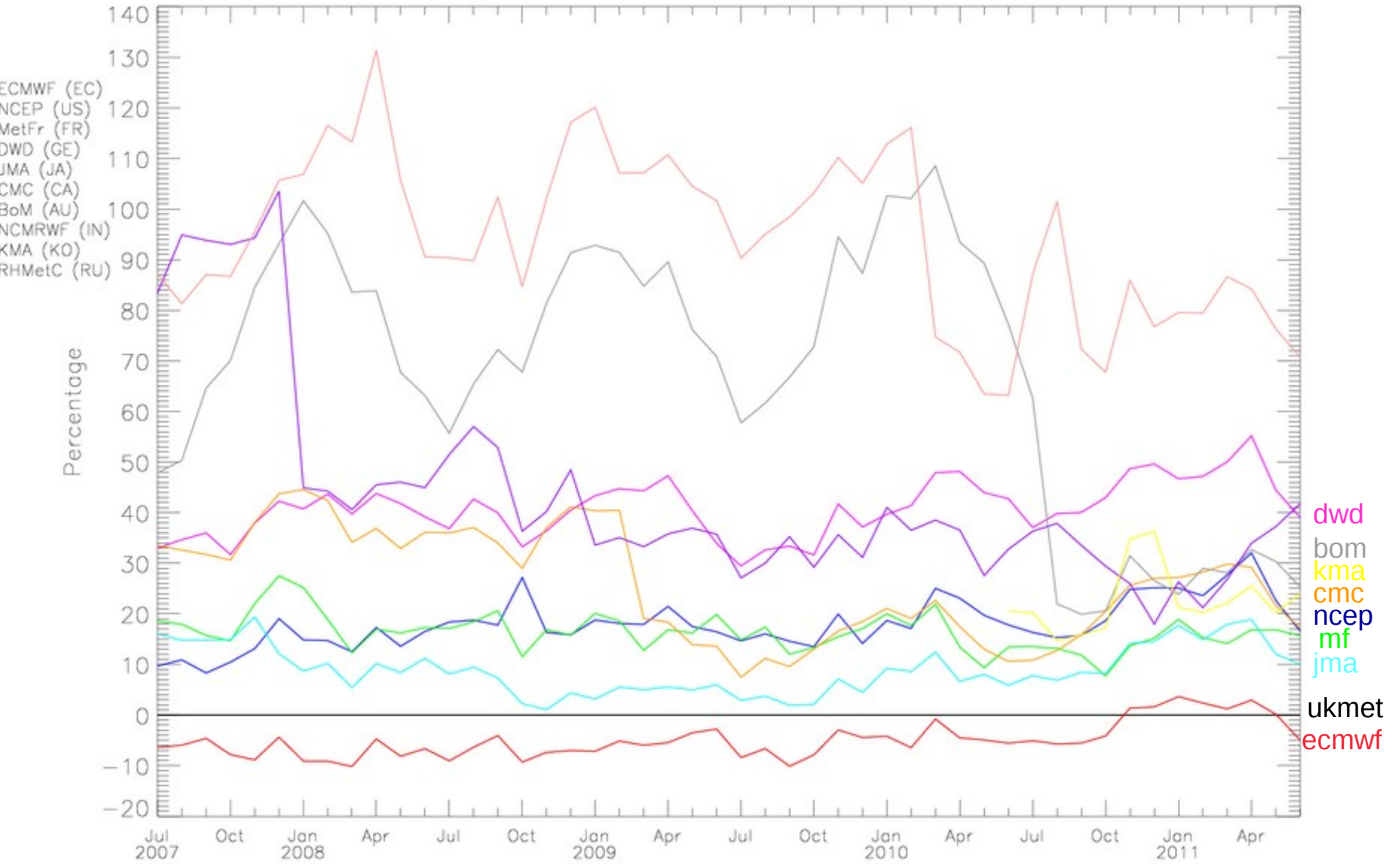


ETKF 1.5km: The setup

- Control analysis from 3DVAR SUK 1.5km 1-h cycle with cloud and latent heat nudging and UK 4km LBC
- IC perturbations* are produced by the ETKF using +1h forecast perturbations in observation space for all the assimilated observations in 3DVAR
 - Surface obs, Aircrafts
 - Radio-sondes
 - GPS, radiances
 - + *Radar derived surface rain rate (not assimilated in VAR)* For more info see Sefano Migliorini's poster
- No localizations
- Variable inflation factor derived from surface obs (u, v, T) and aircraft data (u, v, T)
- No representation of model errors
- LBC taken from MOGREPS-R (24km version)

* for the first cycle IC perturbations are taken from MOGREPS-G as in MOGREPS-R

Ranking relative to Met Office Weighted NWP Index Components Area:NH



Ce que les Britanniques pensent du Met Office ?

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[News](#) > [UK news](#) > [Met Office](#)

Met Office forecasts storm warnings over its accuracy

They are among the most respected, scientific and accurate forecasters in the world. Yet to the British public they are a joke.

Tim Adams visits the Met Office's HQ in Exeter to meet the people for whom the outlook is always gloomy



Tim Adams
The Observer, Sunday 21 February 2010
[Article history](#)

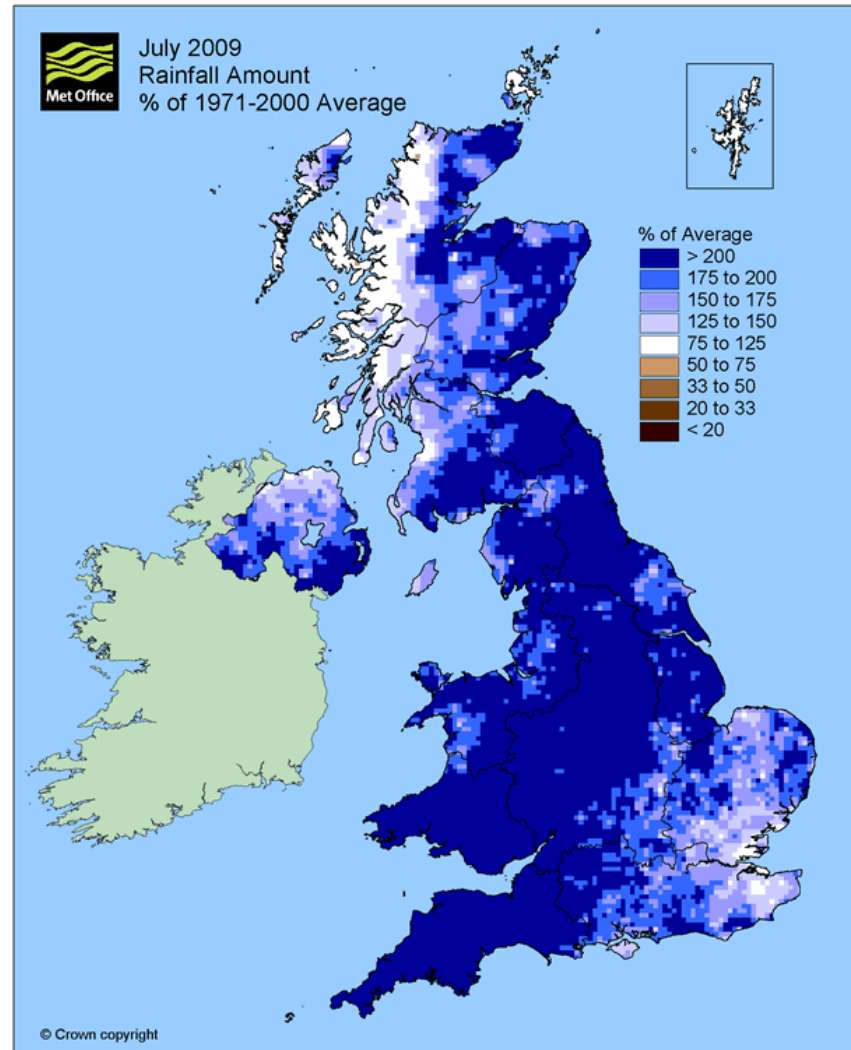


Ah! les prévisions saisonnières ... (1)

Avril 2009

« The coming summer is 'odds on for a barbecue summer' »

Juillet 2009 >





Ah! les prévisions saisonnières ... (2)

Octobre 2009 :

**« Winter temperature
are likely to be near or
above average »**

Janvier 2010 >

