

A mass-conserving semi-implicit semi-Lagrangian scheme for the shallow water equations on the sphere

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Overview

- **Motivations/Aims**
- **Coupling conservative scheme to momentum equations**
- **Other aspects of the SW model**
- **Results**
- **Conclusions**

Motivations (1)

- Mass conservation considered important for climate simulations
- **Inherent conservation more accurate than *ad hoc* corrections**
- **Standard semi-Lagrangian (SL) schemes non-conservative**

Motivations (2)

Currently the Unified Model (the New Dynamics)

- Uses *Eulerian flux* form of continuity equation
- Mixed SL/Eulerian approach
- Inconsistent treatment of equations

Aims of the work

Primary Aim = Develop a global shallow-water model that:

- Inherently conserves mass;
- Stable (unconditional (on Δt) stability);
- Consistent treatment of all equations;
- At least as accurate as present SISL scheme;
- Not a major departure from our existing model (**evolution** not **revolution**)

Continuous Continuity equations

Standard (Lagrangian) shallow-water continuity equation:

$$\frac{D\Phi}{Dt} + \Phi \nabla \cdot \mathbf{u} = 0$$

We use the more primitive integral (conservative) form:

$$\frac{D}{Dt} \left(\int_{\delta\mathcal{A}} \Phi d\mathcal{A} \right) = 0$$

Now apply SL scheme but key to stability is to couple directly to momentum equations

Reference “profile”

Define a reference depth $\Phi^{\text{ref}} = \text{constant}$ and rewrite as:

$$\frac{D}{Dt} \left[\int_{\delta\mathcal{A}} (\Phi - \Phi^{\text{ref}}) d\mathcal{A} \right] + \frac{D}{Dt} \left(\int_{\delta\mathcal{A}} \Phi^{\text{ref}} d\mathcal{A} \right) = 0$$

or as

$$\frac{D}{Dt} \left[\int_{\delta\mathcal{A}} (\Phi - \Phi^{\text{ref}}) d\mathcal{A} \right] + \Phi^{\text{ref}} \left(\int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A} \right) = 0$$

Key aspect is that we have now pulled off a **divergence** term to be treated **semi-implicitly**.

Temporal discretization (1)

Integrating along the trajectory gives the *exact* result

$$\left[\int_{\delta\mathcal{A}} (\Phi - \Phi^{\text{ref}}) d\mathcal{A} \right]_A^{n+1} - \left[\int_{\delta\mathcal{A}} (\Phi - \Phi^{\text{ref}}) d\mathcal{A} \right]_D^n + \Delta t \Phi^{\text{ref}} \int_{\text{traj}} \left(\int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A} \right) dt = 0$$

Temporal discretization (2)

Approximating the time integral along trajectory using weighted averages and grouping terms, gives:

$$\left\{ \int_{\delta\mathcal{A}} \left[\left(\Phi - \Phi^{\text{ref}} \right) + \alpha \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right] d\mathcal{A} \right\}_A^{n+1} = \left\{ \int_{\delta\mathcal{A}} \left[\left(\Phi - \Phi^{\text{ref}} \right) - \beta \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right] d\mathcal{A} \right\}_D^n$$

Temporal discretization (3)

Using a control volume approach $\left\{ \int_{\delta\mathcal{A}} \Phi d\mathcal{A} \right\}_A^n \simeq \Phi_A^n \Delta\mathcal{A}$ ($\Delta\mathcal{A} \equiv$ arrival area):

$$\left[\left(\Phi - \Phi^{\text{ref}} \right) + \alpha \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right]_A^{n+1} = \frac{1}{\Delta\mathcal{A}} \left\{ \int_{\delta\mathcal{A}} \left[\left(\Phi - \Phi^{\text{ref}} \right) - \beta \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right] d\mathcal{A} \right\}_D^n$$

Conservation assured if RHS integration is conservative

(here use **SLICE**)

Compare with Non-conservative SL implementation

Starting point:

$$\frac{D \left(\Phi - \Phi^{\text{ref}} \right)}{Dt} + \Phi^{\text{ref}} \nabla \cdot \mathbf{u} = - \left(\Phi - \Phi^{\text{ref}} \right) \nabla \cdot \mathbf{u}$$

Integrate along trajectory and approximate time integrals:

$$\left[\left(\Phi - \Phi^{\text{ref}} \right) + \alpha \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right]_A^{n+1} = \left[\left(\Phi - \Phi^{\text{ref}} \right) - \beta \Delta t \Phi \nabla \cdot \mathbf{u} \right]_D^n - \alpha \Delta t \left[\left(\Phi - \Phi^{\text{ref}} \right) \nabla \cdot \mathbf{u} \right]_A^{n+1}$$

Other aspects of model

- **Velocity handled via (shallow) rotation matrix**

$$(\mathbf{u} - \alpha \Delta t \Psi)_A^{n+1} = \left[\Lambda (\mathbf{u} + \beta \Delta t \Psi)_{D_L}^n \right]$$

- **Coriolis terms handled as in (Thuburn & Staniforth, 2004), eg:**

$$-2 (\boldsymbol{\Omega}_r \times \mathbf{u})_u = \frac{1}{a \cos \phi \Delta \lambda} \left\langle \langle \tilde{v} \rangle^\phi \frac{f}{\bar{\Phi}} \right\rangle^\lambda$$

where

$$\tilde{v} \equiv a \cos \phi \Delta \lambda \bar{\Phi}^\phi v$$

is a mass flux variable

$$f \equiv 2\Omega \sin \phi$$

Departure points calculations

$$\frac{D\mathbf{x}}{Dt} = \mathbf{u}$$

discretized as

$$(\mathbf{x} + \beta_x \Delta t \mathbf{u})_D^n = (\mathbf{x} - \alpha_x \Delta t \mathbf{u})_A^{n+1}$$

and solved as

$$X_D^{(k)} = -\gamma \Delta t \left[\alpha_x U_A^{n+1} + \beta_x U^n \left(\lambda_D^{(k-1)}, \phi_D^{(k-1)} \right) \right]$$

$$Y_D^{(k)} = -\gamma \Delta t \left[\alpha_x V_A^{n+1} + \beta_x V^n \left(\lambda_D^{(k-1)}, \phi_D^{(k-1)} \right) \right]$$

where

$$\begin{pmatrix} U \\ V \end{pmatrix} = \Lambda \begin{pmatrix} u \\ v \end{pmatrix}$$

Unified implementation (conservative or not)

For non-(or) conservative options, the discrete equations are:

$$\left[\left(\Phi - \Phi^{\text{ref}} \right) + \alpha \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right]^{n+1} = R_{\Phi} \quad (1)$$

$$\left[\mathbf{u} - \alpha \Delta t \Psi \right]^{n+1} = \Lambda (\mathbf{u} + \beta \Delta t \Psi)_{D_L}^n \quad (2)$$

where,

$$R_{\Phi} = \begin{cases} \frac{1}{\Delta \mathcal{A}} \left\{ \int_{\delta \mathcal{A}} \left[\left(\Phi - \Phi^{\text{ref}} \right) - \beta \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right] d\mathcal{A} \right\}_D^n, & \text{for SLICE} \\ \left[\left(\Phi - \Phi^{\text{ref}} \right) - \beta \Delta t \Phi \nabla \cdot \mathbf{u} \right]_D^n - \alpha \Delta t \left[\left(\Phi - \Phi^{\text{ref}} \right) \nabla \cdot \mathbf{u} \right]_A^*, & \text{for SL} \end{cases}$$

The Helmholtz problem

Irrespective of the options (conservative or not), the discrete equations are:

$$\begin{aligned} [\mathbf{u} - \alpha \Delta t \Psi]^{n+1} &= R_u \\ \left[\left(\Phi - \Phi^{\text{ref}} \right) + \alpha \Delta t \Phi^{\text{ref}} \nabla \cdot \mathbf{u} \right]^{n+1} &= R_\Phi \end{aligned}$$

which are combined to obtain the Helmholtz problem for the geopotential perturbation $\Phi' = \Phi - \Phi^{\text{ref}}$,

$$h_1 \delta_\lambda (h_2 \delta_\lambda \Phi') + h_1 \delta_\phi (h_3 \delta_\phi \Phi') - \Phi' = RHS(R_u, R_\Phi) \quad (3)$$

where $h_{j=1,3} = h_j \left(\alpha, \Delta t, a, \cos(\phi), \Phi^{\text{ref}} \right)$ are invariant coefficients for a fixed grid in time. Eq. (3) is solved using the BiCGstab (Van der Vorst, SIAM J. Sci. Stat. Comput., 1992).

The iterative solution procedure

Do time-step loop:

- given $(u, v, \Phi)^n$ at level n

Do outer-loop iteration (departure loop):

- compute (λ_D^n, ϕ_D^n) using $(u, v)^n$ and latest $(u, v)^{n+1}$
- evaluate the departure terms R_D^n

Do inner-loop iteration (Coriolis loop):

- evaluate Coriolis and nonlinear terms $R^* = R^{n+1}$
- solve Helmholtz problem for Φ^{n+1}
- update $(u, v)^{n+1}$

Enddo

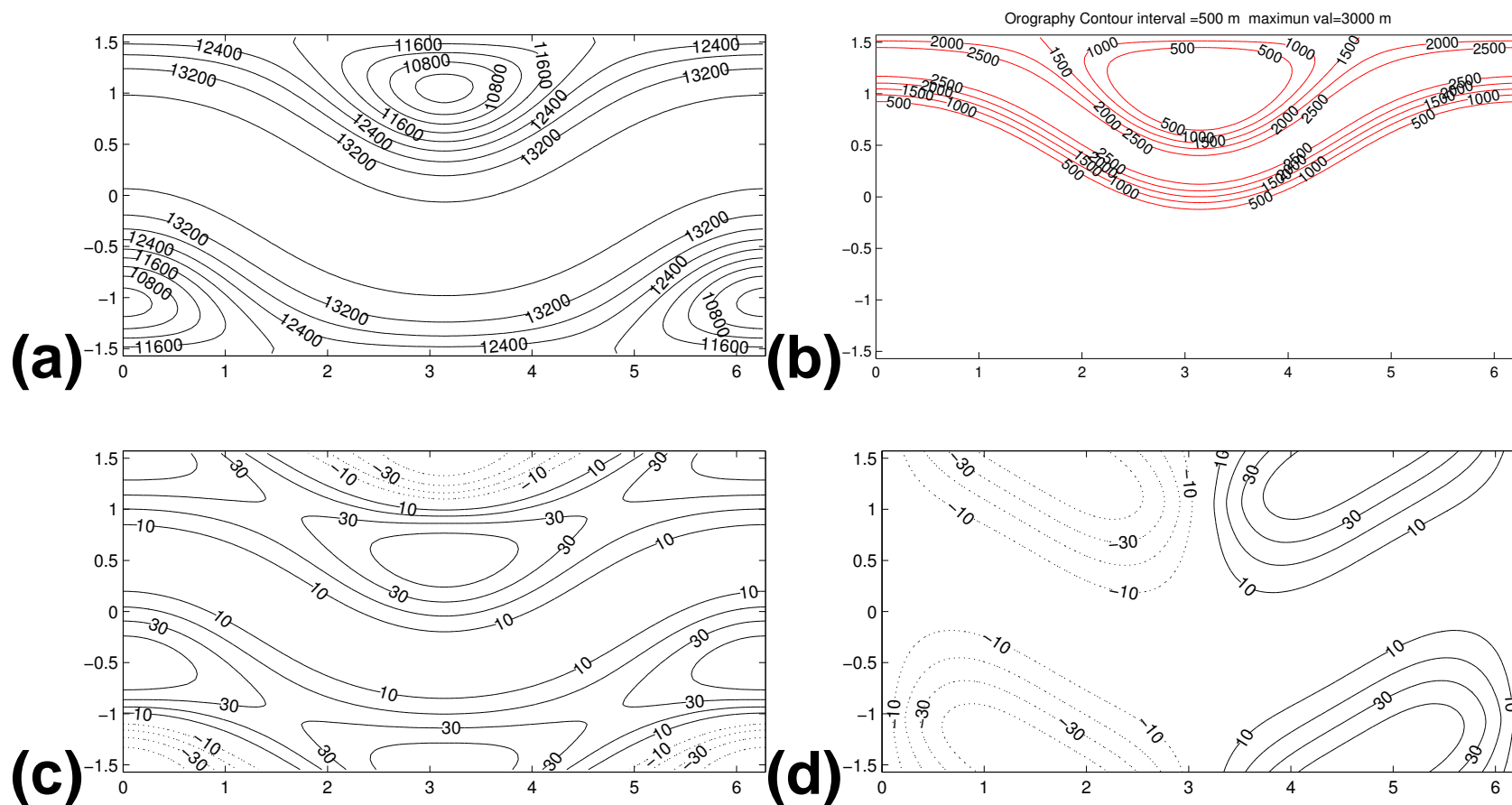
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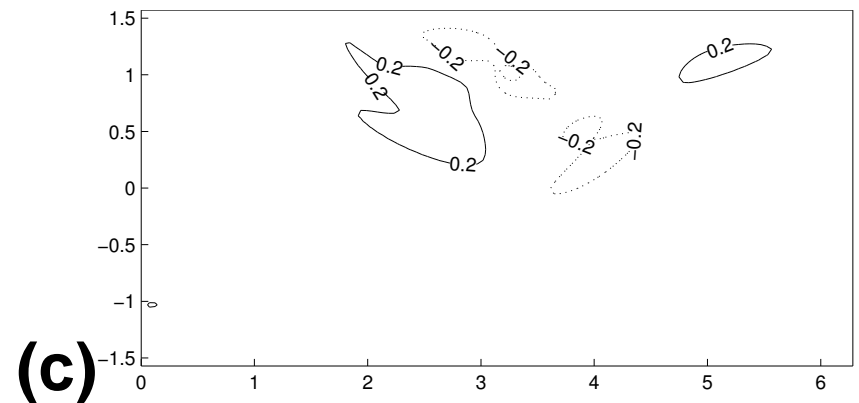
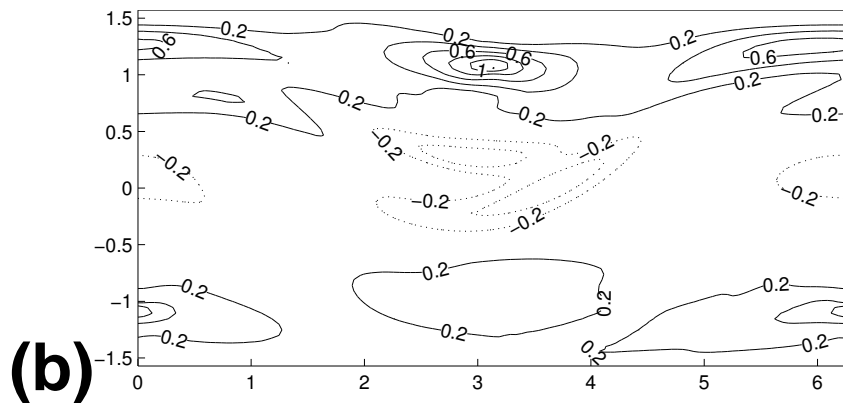
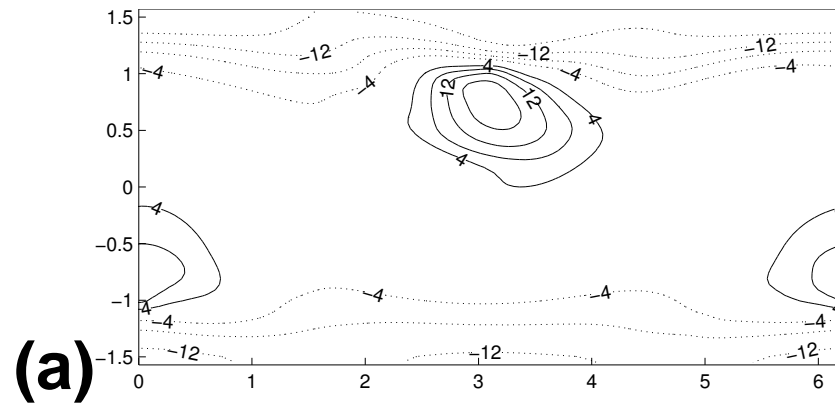
Test cases

1. **Stationary jets over zonal orography (Staniforth & White, 2007)**
2. **Exact unsteady flow (Lauter et al., 2005)**
3. **Flow over an isolated mountain; test case 5 of (Williamson et al., 1992)**

Stationary jets over zonal orography - Initial Fields



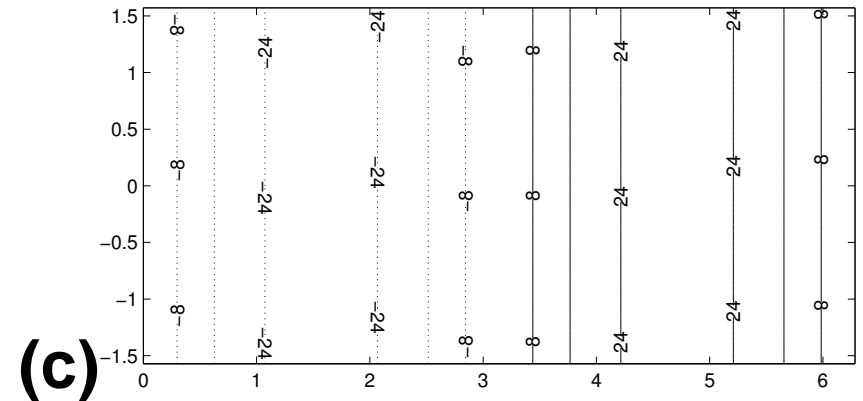
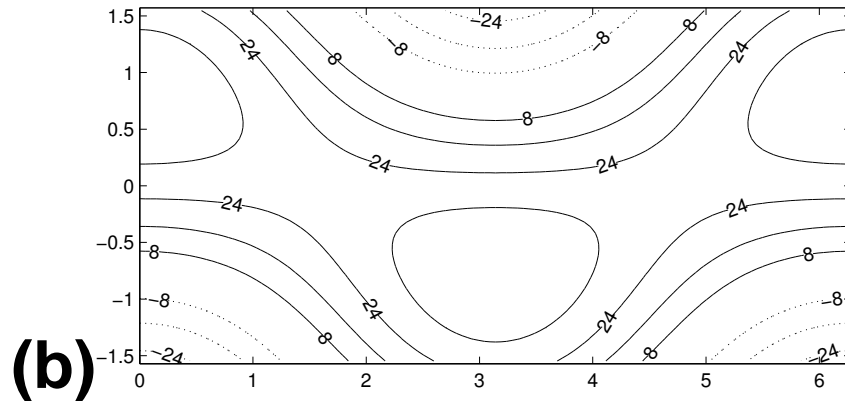
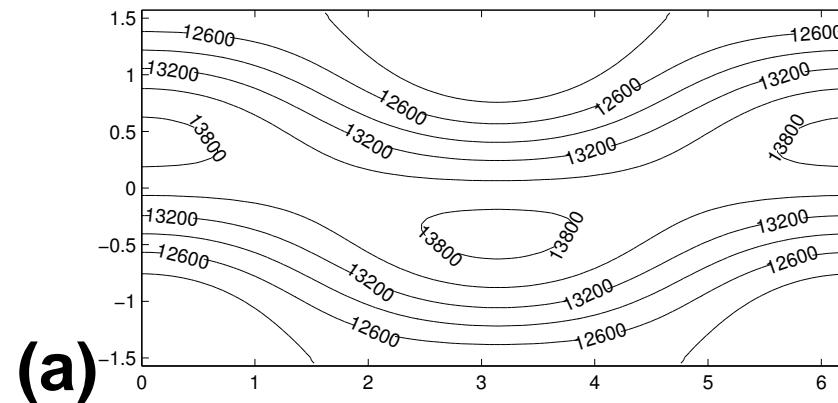
Stationary jets - Error Fields after 5 days



Resolution 128x64; $\Delta t = 1\text{h}$;

(a) $\left(\Phi + \Phi^S\right) / g$ [4 m]; **(b) & (c)** u & v [0.2 m s^{-1}]

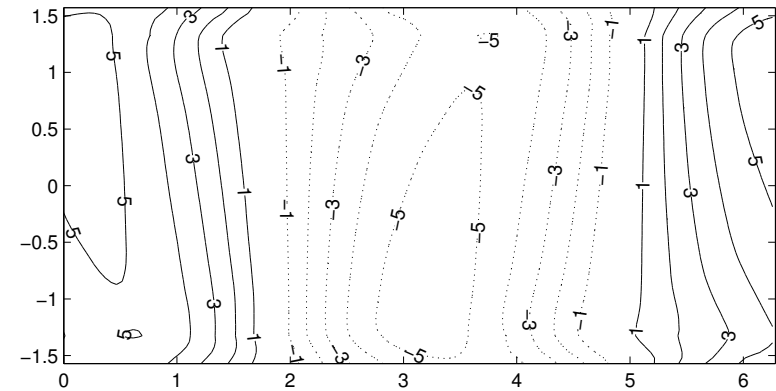
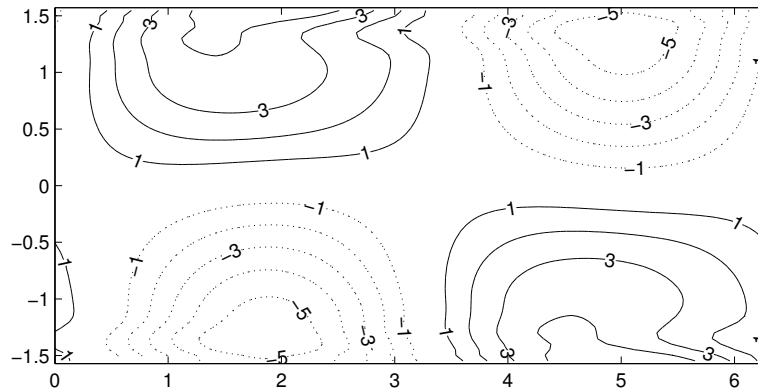
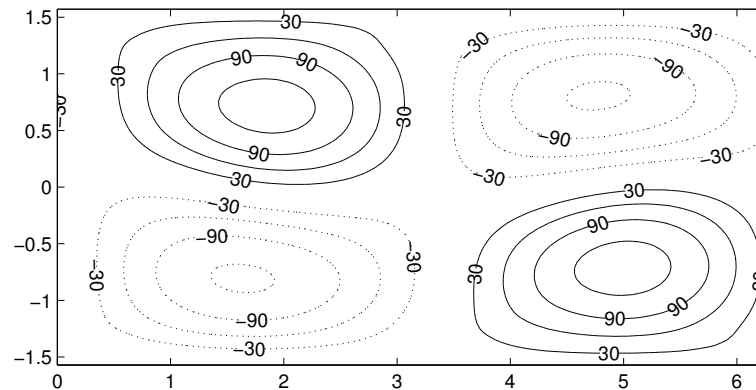
Exact Unsteady Flow - Initial Fields



$$\text{(a)} \quad (\Phi + \Phi^S) / g \text{ [400 m];}$$

$$\text{(b)} \quad u \text{ [10 m s}^{-1}\text{]}; \text{ (c)} \quad v \text{ [10 m s}^{-1}\text{]}.$$

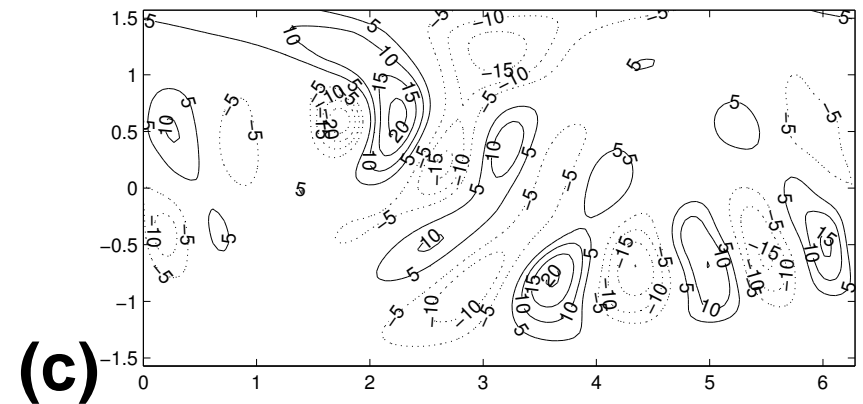
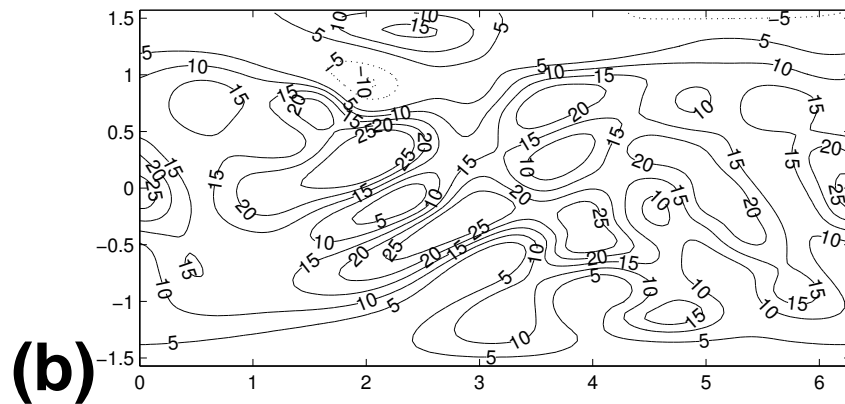
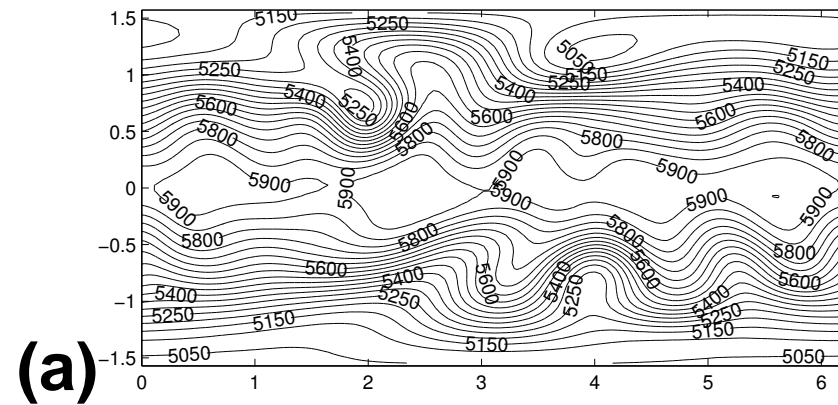
Exact Unsteady Flow - Error Fields 5 Days



$$\Delta t = 1 \text{ h}; (I = 128, J = 64).$$

Contour intervals: 30 m for $(\Phi + \Phi^S) / g$; 1 m s⁻¹ for u and v .

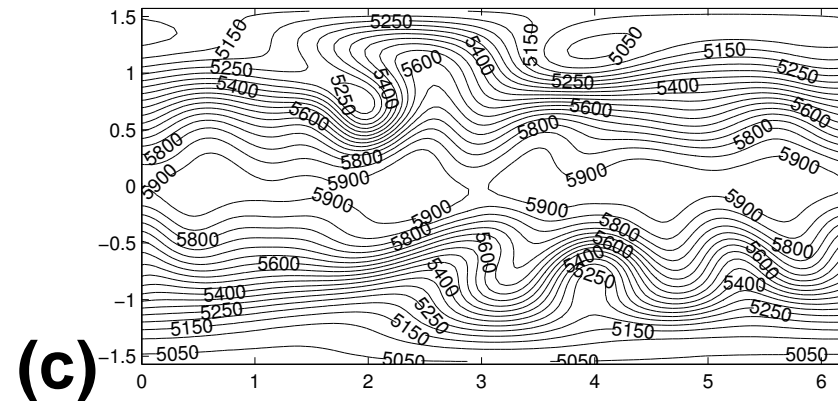
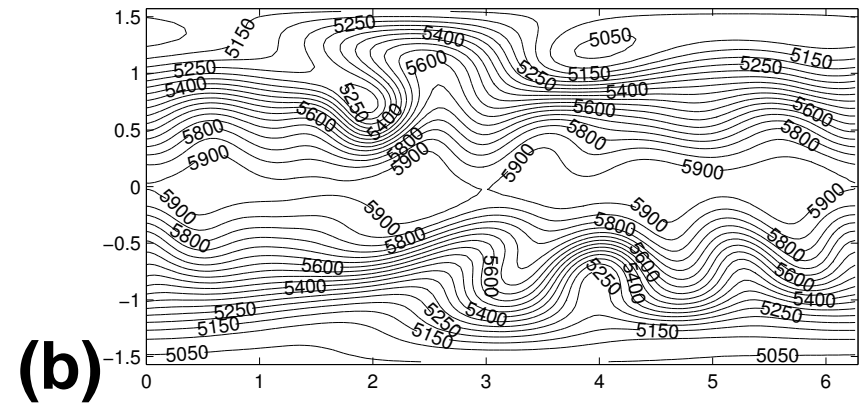
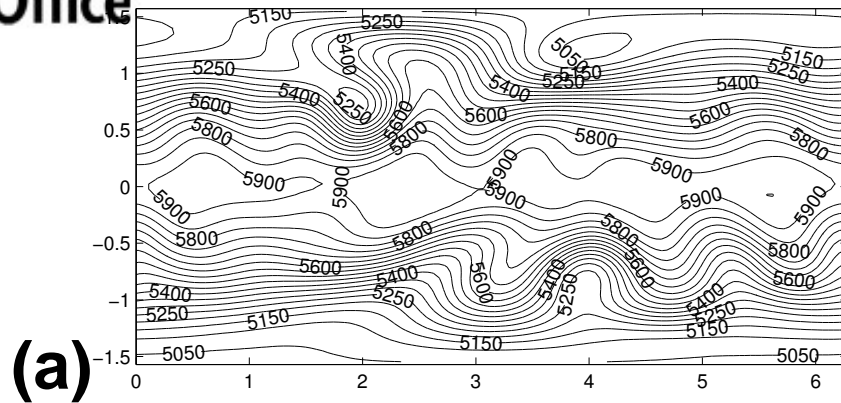
Williamson Test Case 5 - Fields 15 Days



Mass-conserving $\Delta t = 600 \text{ s}; (128 \times 64)$: (a) $(\Phi + \Phi^S) / g [50 \text{ m}]$;

(b) $u [5 \text{ m s}^{-1}]$; (c) $v [5 \text{ m s}^{-1}]$.

Williamson Test Case 5 - Height Fields 15 Days



a) Mass-conserving $\Delta t = 600$ s; b) $\Delta t = 6000$ s;

c) Non-conserving model $\Delta t = 6000$ s.

Comparison with SISL - L2 Error Norms

Stationary jets over zonal orography

$I \times J$	$\Delta t(\text{mn})$	SLICE	SL	SLICE	SL
		$L_2(h)$	$L_2(h)$	$L_2(v)$	$L_2(v)$
64 × 32	12	0.165E-02	0.178E-02	0.333E-01	0.363E-01
128 × 64	6	0.406E-03	0.464E-03	0.839E-02	0.946E-02
256 × 128	3	0.989E-04	0.116E-03	0.208E-02	0.238E-02

Comparison with SISL - L2 Error Norms

Exact Unsteady Flow

$I \times J$	Δt (mn)	SLICE	SL	SLICE	SL
		$L_2(h)$	$L_2(h)$	$L_2(\mathbf{v})$	$L_2(\mathbf{v})$
64 × 32	12	0.176E-02	0.179E-02	0.428E-01	0.429E-01
128 × 64	6	0.447E-03	0.447E-03	0.114E-01	0.114E-01
256 × 128	3	0.113E-03	0.111E-03	0.289E-02	0.285E-02

Reducing the dependency on the reference state

- The generalised approach of using a reference profile to couple SLICE to a SISL treatment of the momentum equations has **a weak dependency on the reference state**.
- There is an alternative approach that **reduces this dependency**.
- A concomitant benefit of the alternative approach is the **improved consistency between mass field and conserved scalar fields**.

The shallow water continuity equation is written in conservative Lagrangian form as

$$\frac{D}{Dt} \left(\int_{\delta\mathcal{A}} \Phi d\mathcal{A} \right) = 0, \quad (4)$$

Consider Φ^{ref} as constant reference value for Φ . Substitution of $\Phi = \Phi - \Phi^{ref} + \Phi^{ref}$ in (4) gives:

$$\frac{D}{Dt} \left[\int_{\delta\mathcal{A}} \left(\Phi - \Phi^{ref} \right) d\mathcal{A} \right] = -\Phi^{ref} \int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A}. \quad (5)$$

Eq. (5) also implies the evolution of the material area element $d\mathcal{A}$,

$$\frac{D \left(\int_{\delta\mathcal{A}} d\mathcal{A} \right)}{Dt} = \int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A}, \quad (6)$$

A CSISL discretization of (5) is:

$$\left\{ \int_{\delta\mathcal{A}} (\Phi - \Phi^{ref}) d\mathcal{A} \right\}_A^{n+1} - \left\{ \int_{\delta\mathcal{A}} (\Phi - \Phi^{ref}) d\mathcal{A} \right\}_D^n = -\alpha\Delta t\Phi^{ref} \left(\int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A} \right)_A^{n+1} - \beta\Delta t\Phi^{ref} \left(\int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A} \right)_D^n. \quad (7)$$

A discrete form of (6) is implied by (7) and can be obtained by setting $\Phi \equiv 0$ therein as:

$$\left(\int_{\delta\mathcal{A}} d\mathcal{A} \right)_A^{n+1} - \left(\int_{\delta\mathcal{A}} d\mathcal{A} \right)_D^n = \alpha\Delta t \left(\int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A} \right)_A^{n+1} + \beta\Delta t \left(\int_{\delta\mathcal{A}} \nabla \cdot \mathbf{u} d\mathcal{A} \right)_D^n.$$

This equation is not forced to hold and therefore in general will not do so \Rightarrow Therefore the numerical solution is dependent on Φ^{ref} .

Dependence on the reference value, test 1 results

L_2 norms after 5 days; The mid-latitude jets problem;
Resolution 128x64; $\Delta t = 60$ mins;

RSS (Reference-State Scheme)

Φ^{ref}	$L_2(h)$	$L_2(\mathbf{v})$
Φ_1^{ref}	0.405E-03	0.859E-02
Φ_2^{ref}	0.406E-03	0.839E-02
Φ_3^{ref}	0.445E-03	0.934E-02

Reference-State-Free Scheme (RSFS) (1)

Specifically, the discrete SISL form of (4) is

$$\left(\int_{\delta\mathcal{A}} \Phi d\mathcal{A} \right)_A^{n+1} = \left(\int_{\delta\mathcal{A}} \Phi d\mathcal{A} \right)_D^n, \quad (8)$$

i.e., (7) with $\Phi^{ref} \equiv 0$. This can be rewritten as

$$\Phi_A^{n+1} = \frac{1}{\Delta\mathcal{A}} \left(\int_{\delta\mathcal{A}} \Phi d\mathcal{A} \right)_D^n. \quad (9)$$

Reference-State-Free Scheme (RSFS) (2)

Then, to maintain the same implicit form as (1), add the term $-\Phi^{ref} + \alpha\Delta t\Phi^{ref} (\nabla \cdot \mathbf{u})_A^{n+1}$ on both sides of (9) to obtain:

$$\left[\left(\Phi - \Phi^{ref} \right) + \alpha\Delta t\Phi^{ref} \nabla \cdot \mathbf{u} \right]_A^{n+1} = R_{\Phi 2}^n, \quad (10)$$

where

$$R_{\Phi 2}^n \equiv \frac{1}{\Delta\mathcal{A}} \left(\int_{\delta\mathcal{A}} \Phi d\mathcal{A} \right)_D^n - \Phi^{ref} + \alpha\Delta t\Phi^{ref} (\nabla \cdot \mathbf{u})_A^{n+1}. \quad (11)$$

Unified approach non/or-conservative scheme

For all the schemes, the discrete equations are:

$$\begin{aligned} [\mathbf{u} - \alpha\Delta t\Psi]^{n+1} &= R_u \\ \left[\left(\Phi - \Phi^{\text{ref}} \right) + \alpha\Delta t\Phi^{\text{ref}}\nabla \cdot \mathbf{u} \right]^{n+1} &= R_\Phi^m \end{aligned}$$

where,

$$m = SL, RSS, RSFS1, RSFS2$$

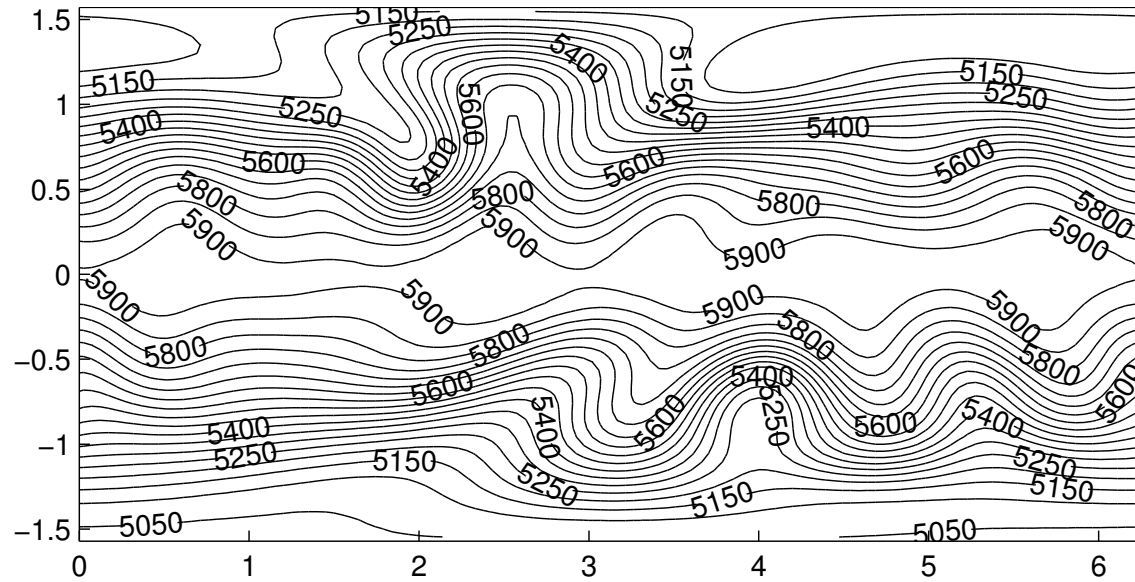
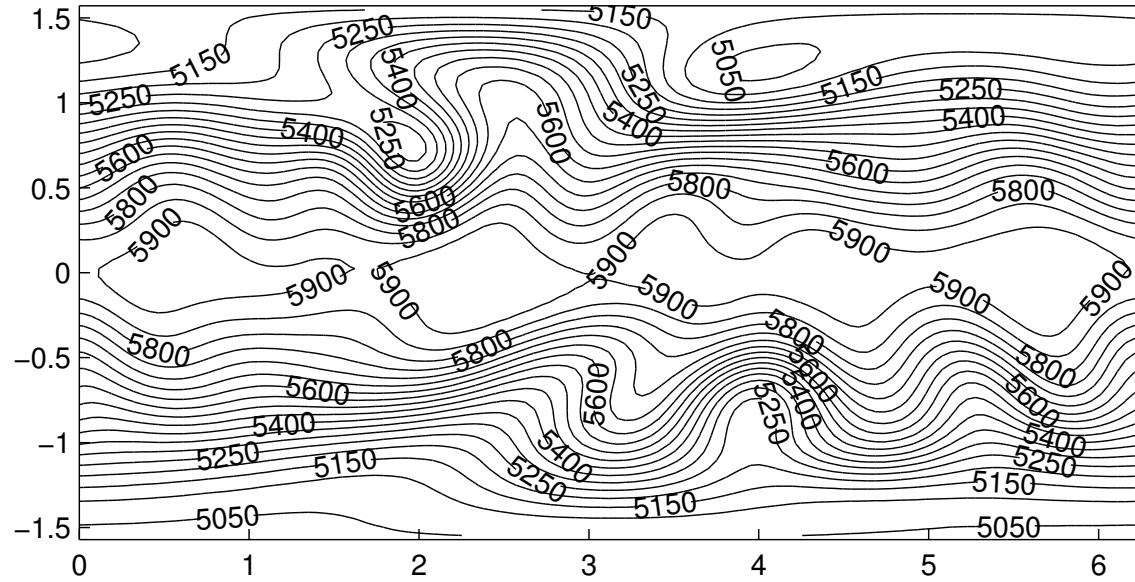
Reference-State-Free Scheme (RSFS), test 1 results

L_2 norms after 5 days; The mid-latitude jets problem;
Resolution 128x64; $\Delta t = 60$ mins.

RSFS1 (Reference-State-Free Scheme with (λ, ϕ) -remapping)

Φ^{ref}	$L_2(h)$	$L_2(\mathbf{v})$
Φ_1^{ref}	0.842E-03	0.207E-01
Φ_2^{ref}	0.845E-03	0.207E-01
Φ_3^{ref}	0.845E-03	0.207E-01

Test 5 (Williamson et al.); Solution after 15 days; Resolution 128x64; (a) $\Delta t = 600s$; (b) $\Delta t = 6000s$





Met Office

An alternative remapping coordinate (FFS2 scheme)

The implied divergence averaged along the trajectory for SLICE, \overline{D}^{SLICE} is:

$$\overline{D}^{SLICE} \equiv \frac{2}{\Delta t} \left(\frac{\delta \mathcal{A}_A - \delta \mathcal{A}_D^{SLICE}}{\delta \mathcal{A}_A + \delta \mathcal{A}_D^{SLICE}} \right), \quad (12)$$

$\delta \mathcal{A}_D^{SLICE}$ can be written as:

$$\delta \mathcal{A}_D^{SLICE} = \delta \mathcal{A}_A \left[1 + \Delta t \overline{D}^{exact} + O(\Delta t^2) \right]$$

Substituting this into (12) shows that **the implied trajectory-average divergence is only first-order accurate in time.**

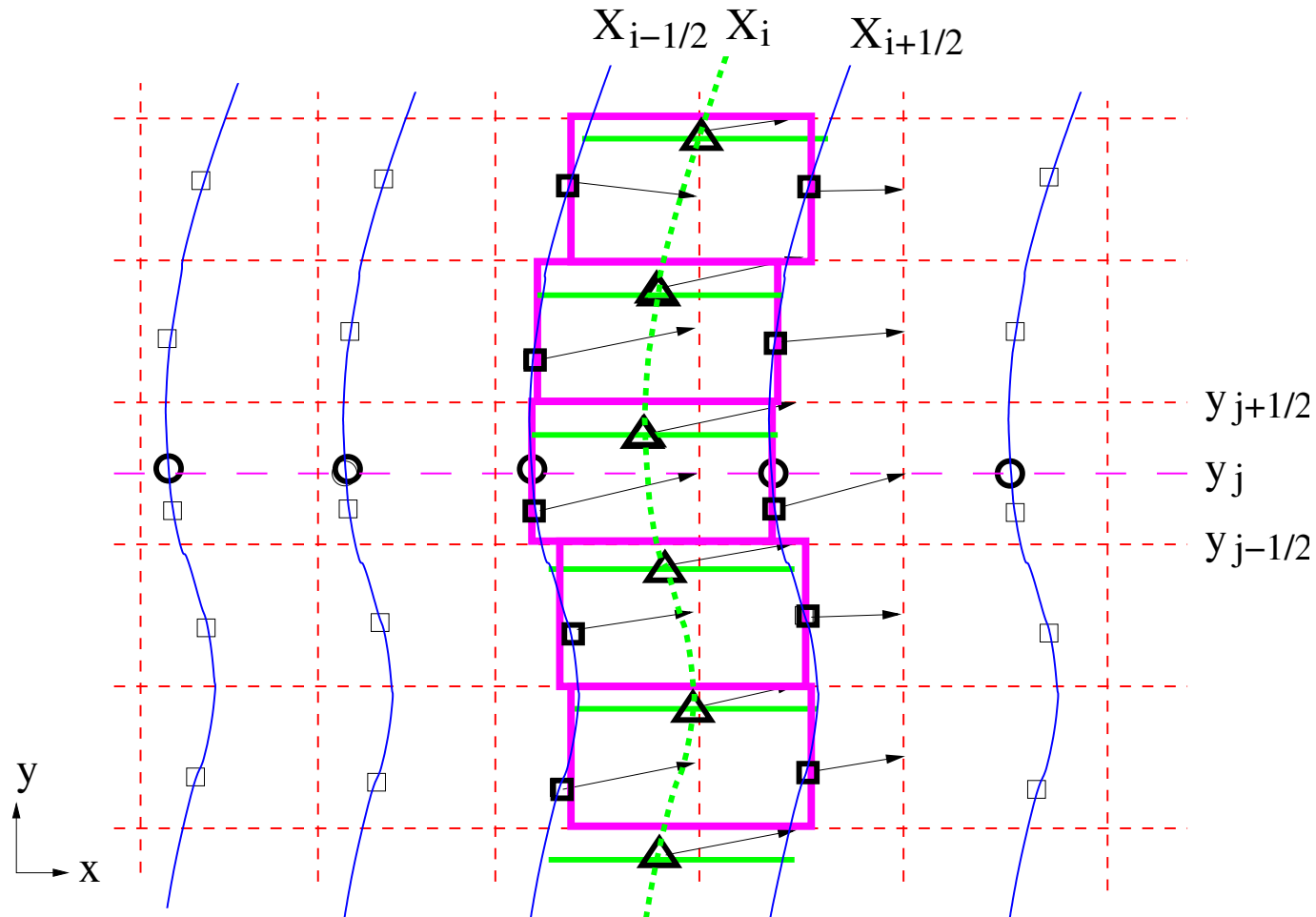
It is possible to restore the second-order accuracy in time, by:

- **computing an estimate for the departure areas whose implied trajectory-average divergence is second-order accurate, and**
- **modifying SLICE to allow the required departure cell areas to be imposed.**
 - **A satisfactory solution is to make an iterative improvement to the implied SLICE departure areas:**

$$\delta A_D^{SLICE} = \delta A_A - \frac{\Delta t}{2} \left(D_A \delta A_A + \{D_D \delta A_D\}^{SLICE} \right). \quad (13)$$

- **A modification is required to SLICE to force it to see the accurate estimate of δA_D^{SLICE} .**

This achieved by changing the **second remapping coordinate**, as the area integrated.



Schematic for C-SLICE

FFS2 results

L_2 norms after 5 days; The mid-latitude jets problem.
Resolution 128x64; $\Delta t = 60$ mins.

RSFS2 [((λ , Area)-remapping]

Φ^{ref}	$L_2(h)$	$L_2(\mathbf{v})$
Φ_1^{ref}	0.437E-03	0.103E-01
Φ_2^{ref}	0.437E-03	0.103E-01
Φ_3^{ref}	0.437E-03	0.103E-01

RSFS1 [((λ , ϕ)-remapping]

Φ^{ref}	$L_2(h)$	$L_2(\mathbf{v})$
Φ_1^{ref}	0.842E-03	0.207E-01
Φ_2^{ref}	0.845E-03	0.207E-01
Φ_3^{ref}	0.845E-03	0.207E-01

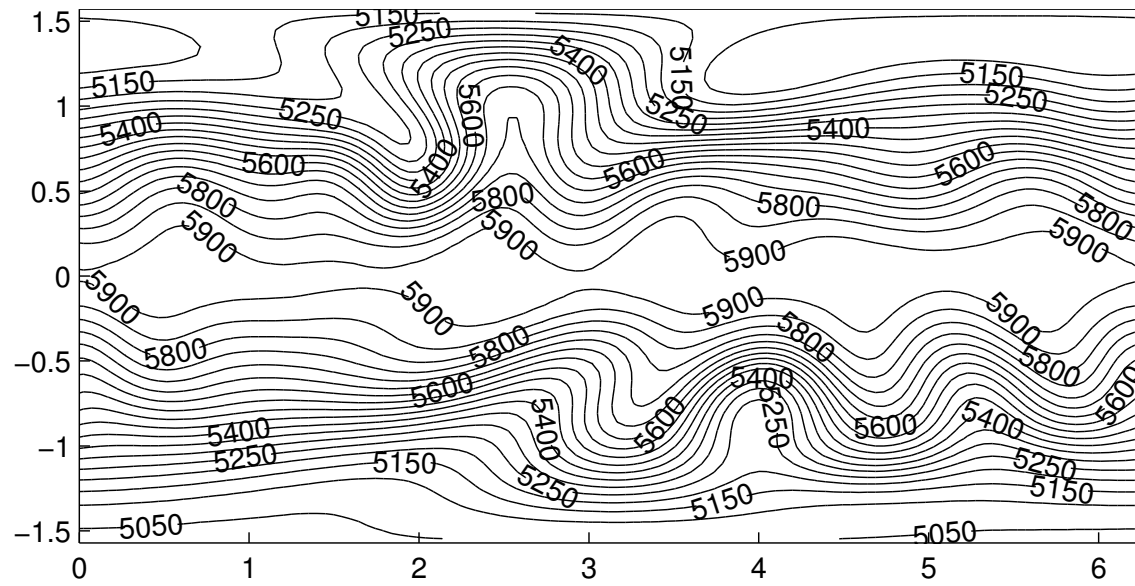


Figure 1: Test 5, 15 days, scheme SFS1, $\Delta t = 6000s$

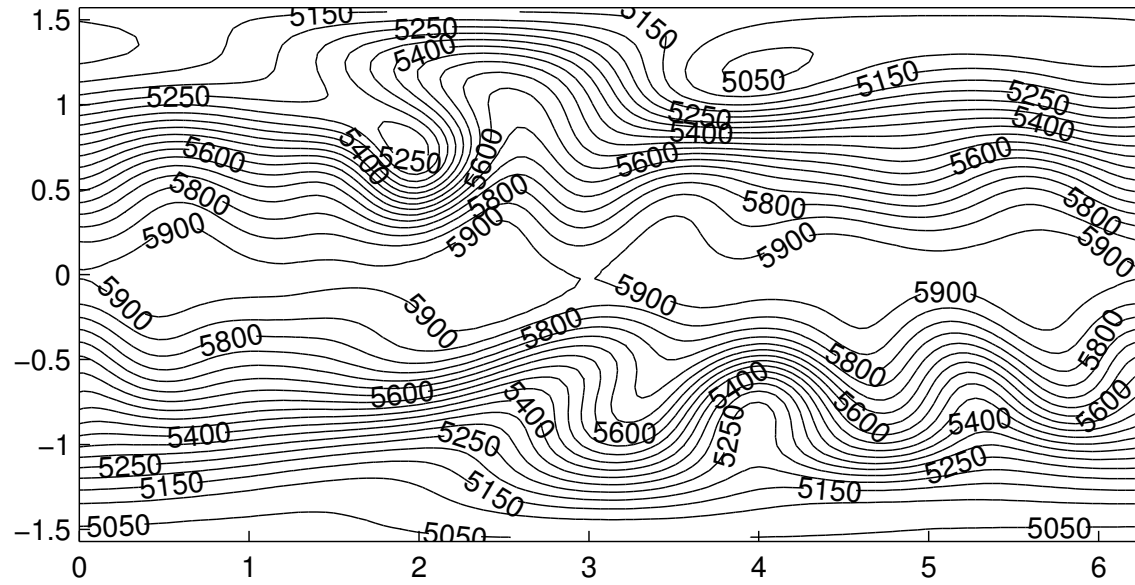
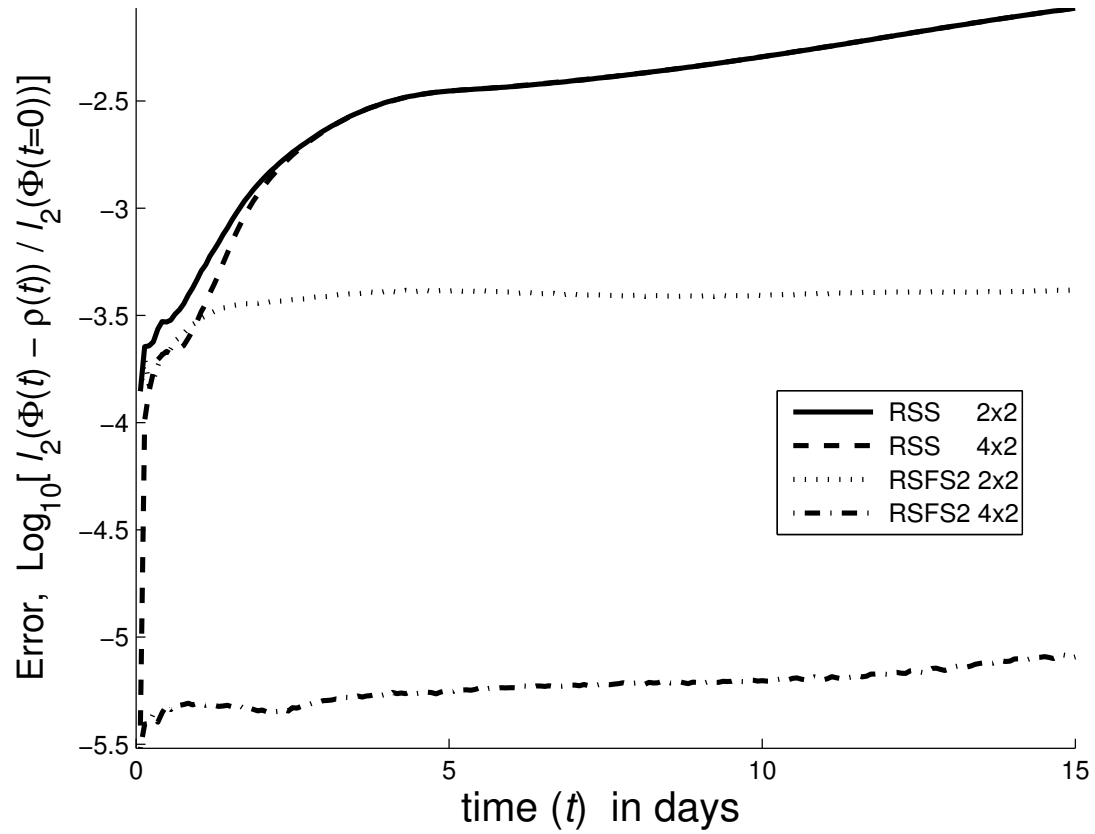


Figure 2: Test 5, 15 days, scheme SFS2, $\Delta t = 6000s$

Consistency of scalar transport: Test 5 of Williamson et al.



Conclusions

- Simple framework for implicit coupling of SLICE (or any Conservative scheme) to momentum equations
- Improved handling of Coriolis terms on C-grid.
- Performs at least as well as standard SL continuity scheme but with mass exactly conserved
- Semi-implicit semi-Lagrangian (No restriction on CFL number ==> Large time steps)
- Inherently conserves mass using a high-order remapping scheme (SLICE)
- Option to reduce the dependence on the reference profile
- Option to preserve a constant for a non-divergent flow
- Optimal generalization to three dimensions
- Option to eliminate the mass-wind inconsistency problem

Some Related Publications

M. Zerroukat, N. Wood, and A. Staniforth, An improved version of SLICE for conservative monotonic remapping on C-grid, [Q. J. R. Meteorol. Soc.](#), vol. 135, 541-546 (2009).

M. Zerroukat, N. Wood, A. Staniforth, A. A. White and J. Thuburn, An inherently mass conserving semi-implicit semi-Lagrangian discretisation of the shallow water equations on the sphere, [Q. J. R. Meteorol. Soc.](#), (in press).

J. Thuburn, M. Zerroukat, N. Wood and A. Staniforth, Coupling a mass conserving semi-Lagrangian scheme (SLICE) to a semi-implicit discretization of the shallow water equations: minimizing the dependence on a reference atmosphere; [Q. J. R. Meteorol. Soc.](#), (submitted).